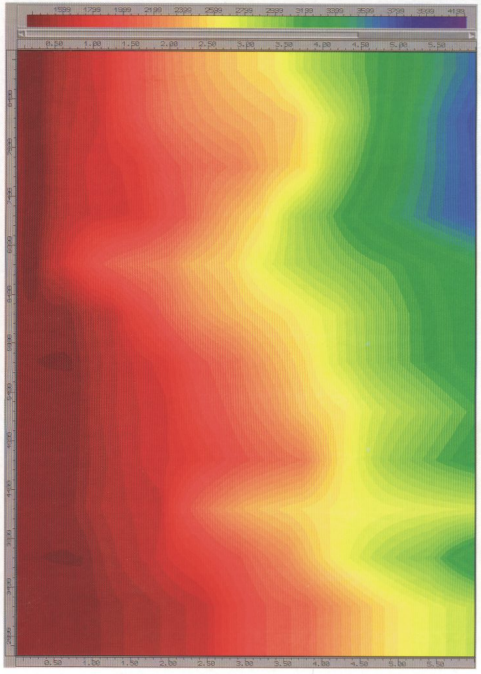


# Introduction to frequency domain waveform inversion: theory and applications

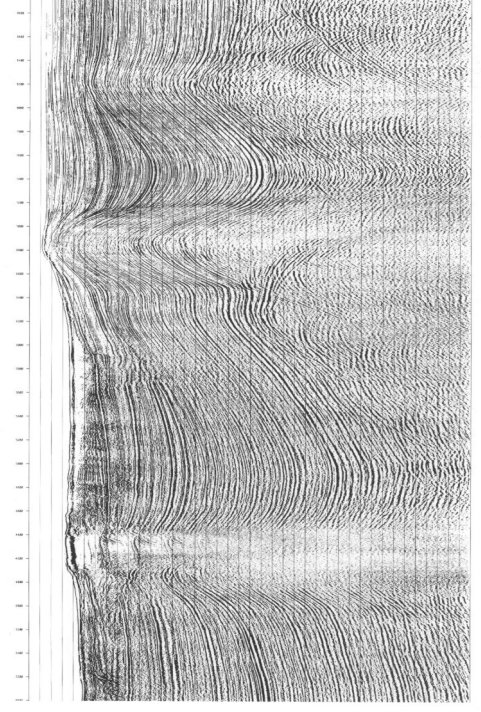
R. Gerhard Pratt, Rie Kamei  
University of Western Ontario

# Imaging with Exploration Seismic Data – the Fundamentals

- Common practice in reflection processing:
  - The “model” and the “image” have distinct spectral characteristics
  - Each are derived from distinct aspects of the data



Seismic velocity model



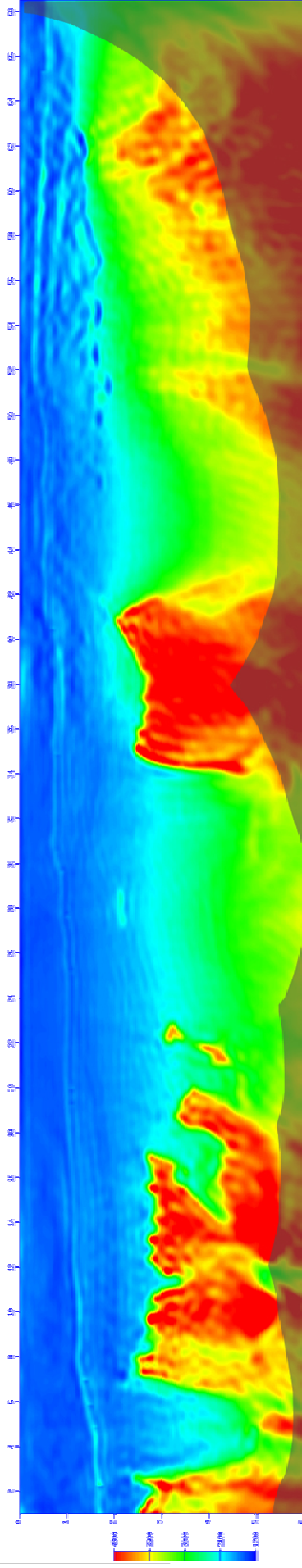
Seismic reflectivity image

Textbook example (Yilmaz 2003)

# The Basic Idea

- Common practice in reflection processing:
  - The “model” and the “image” have distinct spectral characteristics
  - Each are derived from distinct aspects of the data
- The approach of waveform tomography:
  - No distinction between model and image
  - All available information contributes to spectrum of the result
  - Overlap in methodology (“migration-like”)
  - Significant change in philosophy

# Waveform tomography of long offset seismic data



(Pratt and Brenders, EAGE workshop, 2004)

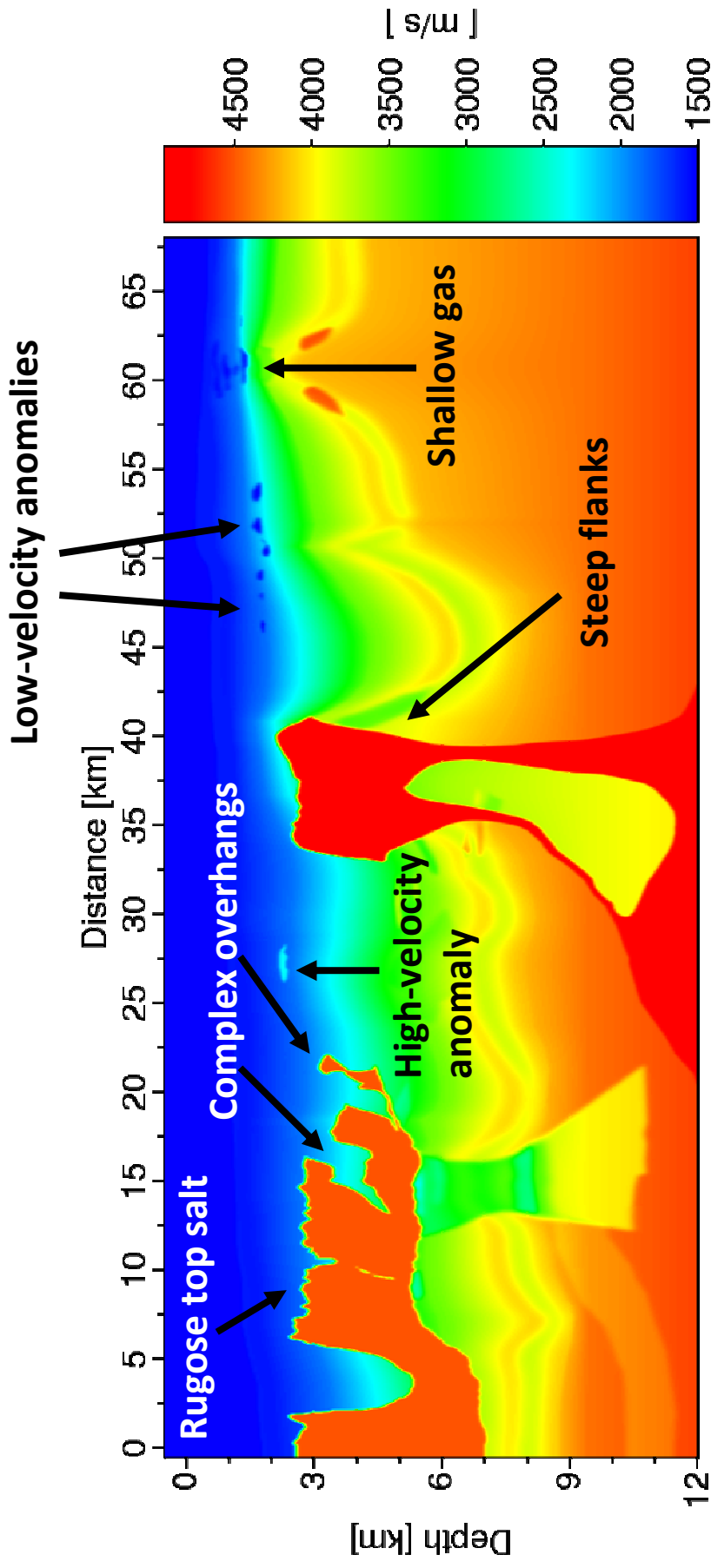
# Outline

- Long offset seismic data
  - An imaging opportunity
- **Waveform Tomography**
  - Theoretical aspects
  - Making it work
- **Nankai Trough Subduction Zone Example**
  - Geological implications
  - The Challenge of High Resolution

# Outline

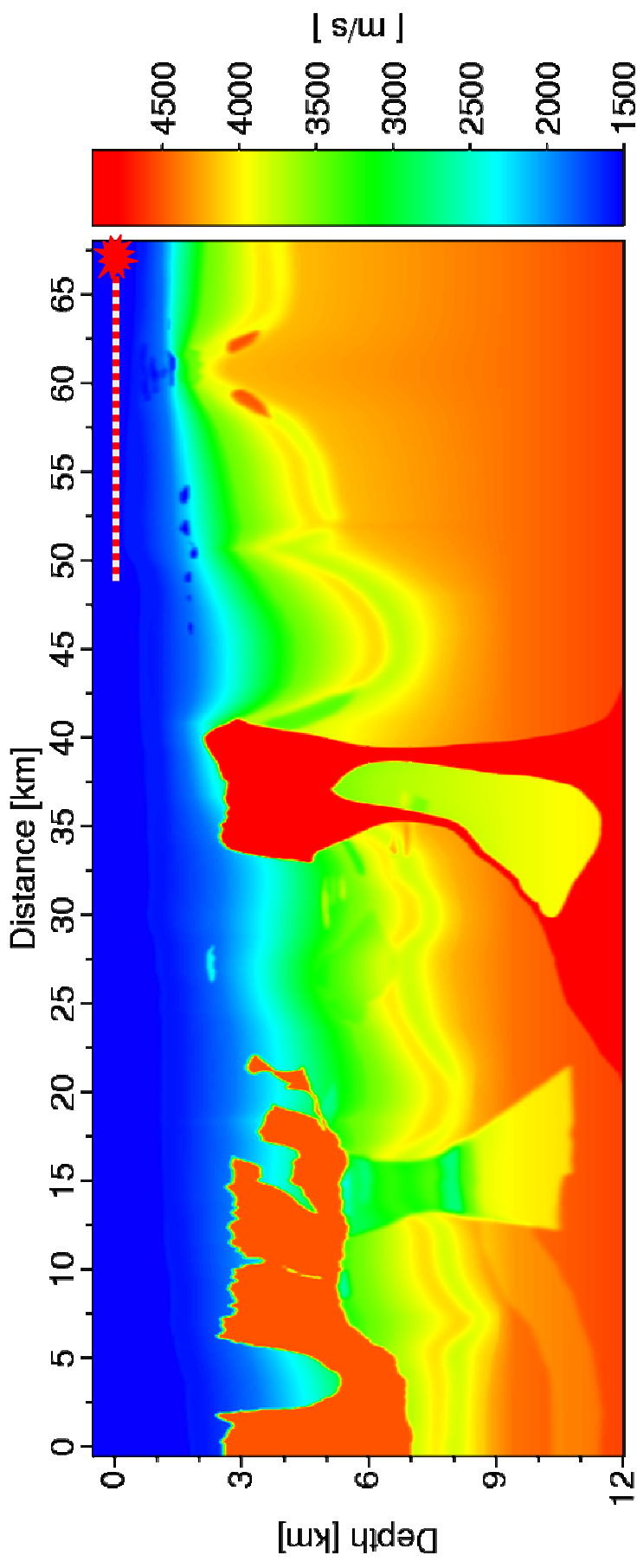
- Long offset seismic data
  - An imaging opportunity
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# 2004 BP Velocity Benchmark: Model



- 67 km wide, 12 km deep
- Built on a 6.25 m x 6.25 m grid
- Range of imaging challenges

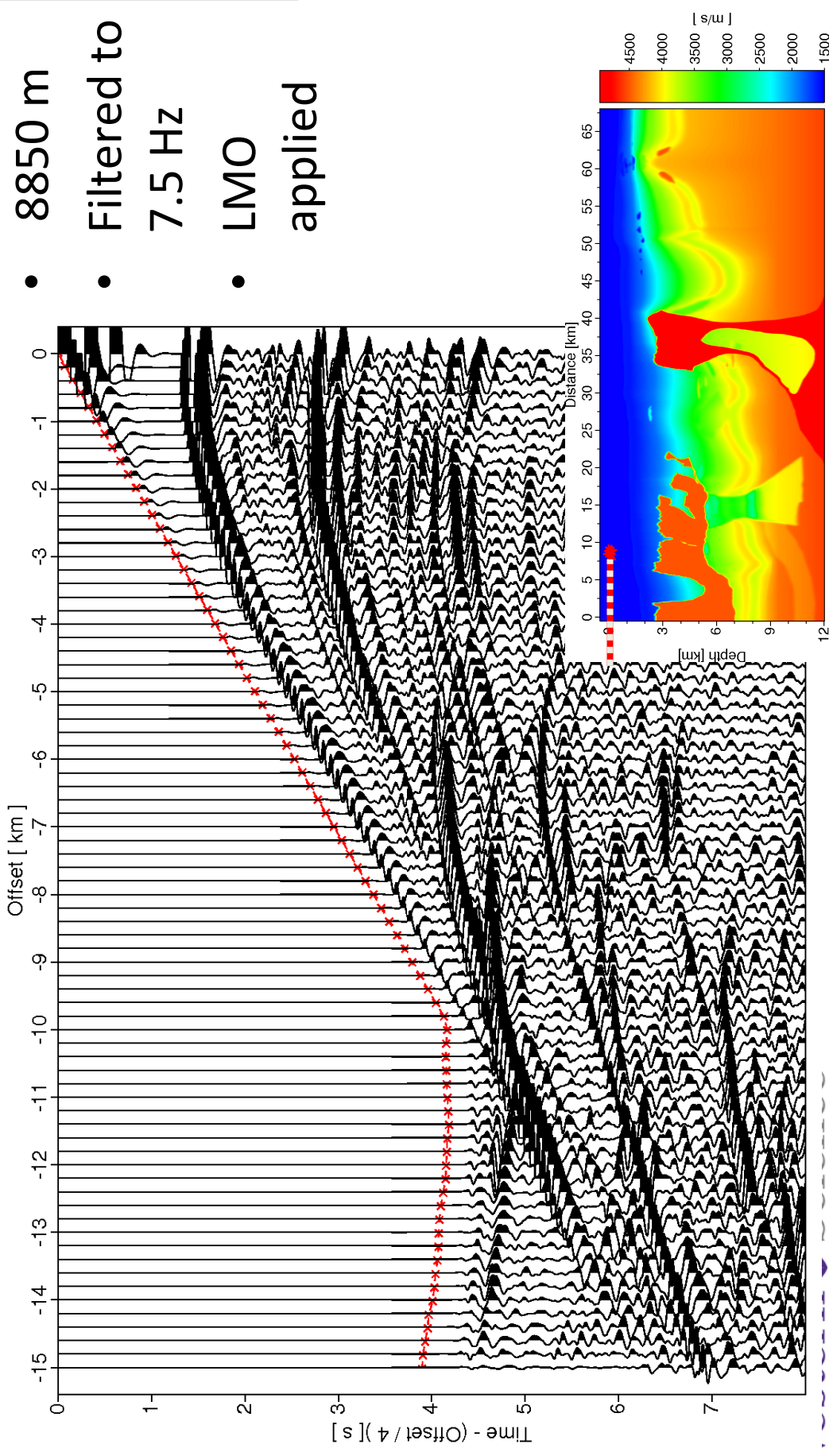
# 2004 BP Velocity Benchmark: Synthetic Data



- 2-D finite-difference acoustic wave equation
- Shot spacing: 50 m, 1348 shots total, 1201 receivers per shot
- 15 km offset streamer data provided

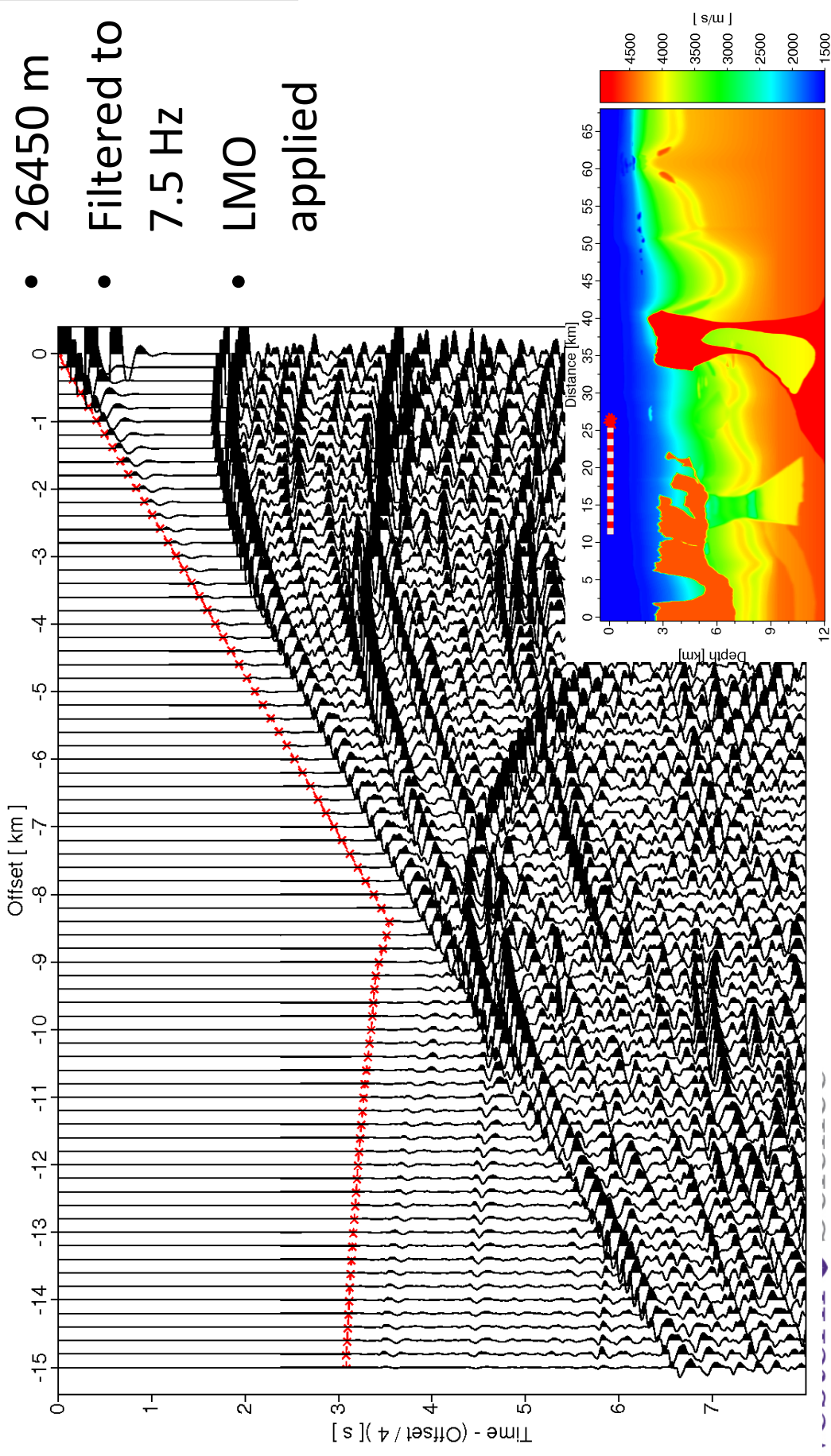


# 2004 BP Velocity Benchmark: Synthetic Data

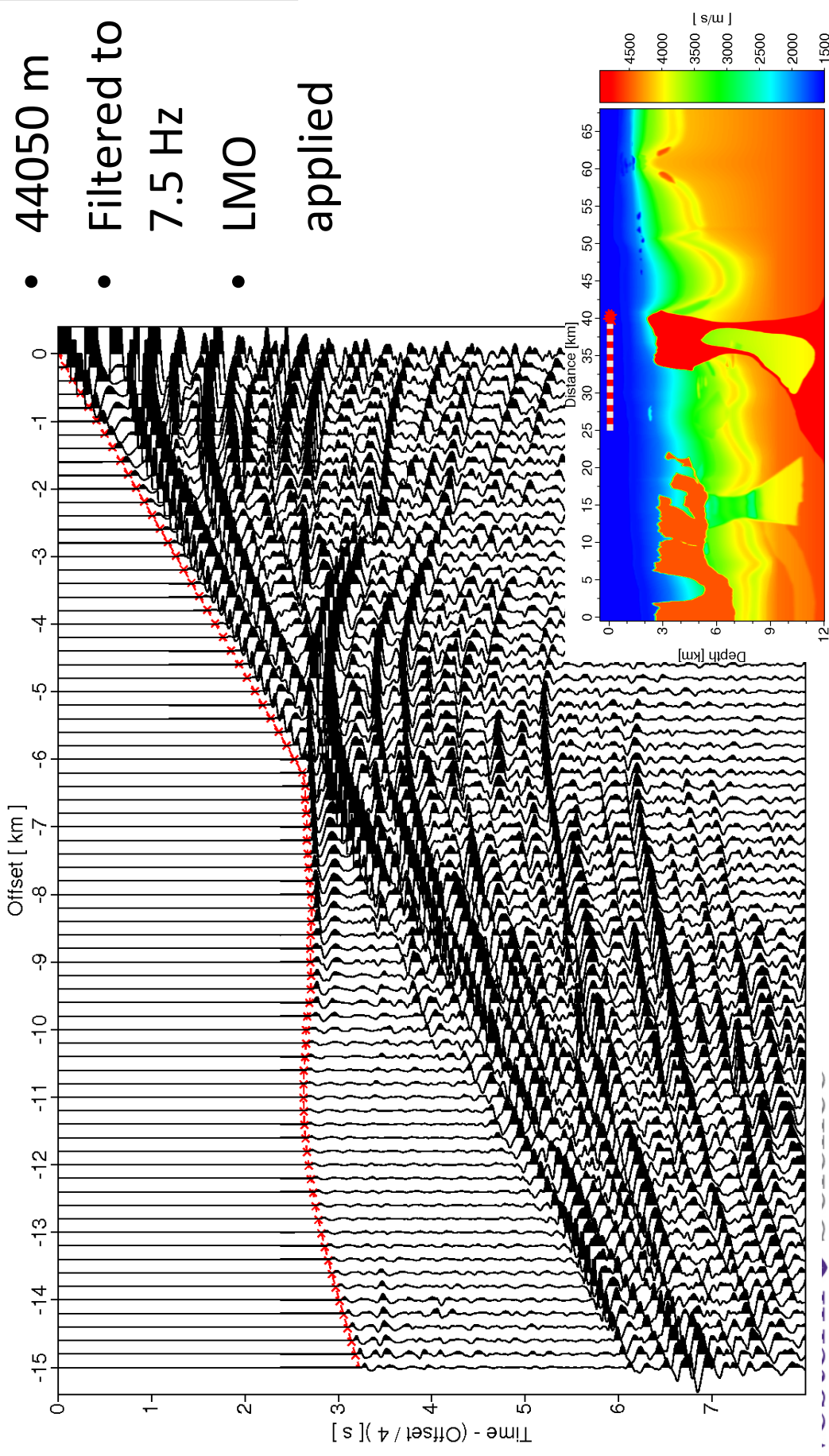


# 2004 BP Velocity Benchmark: Synthetic Data

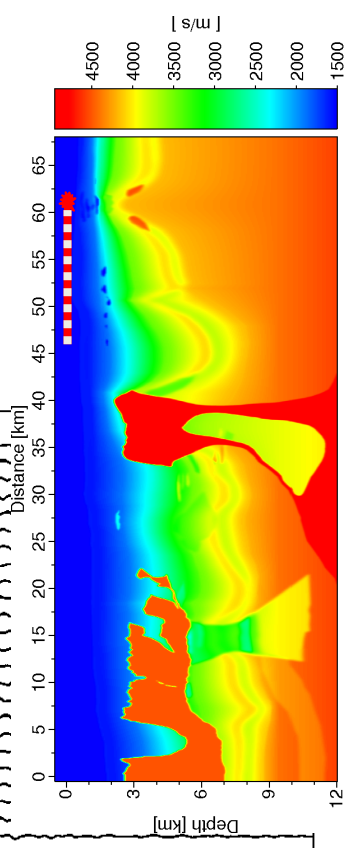
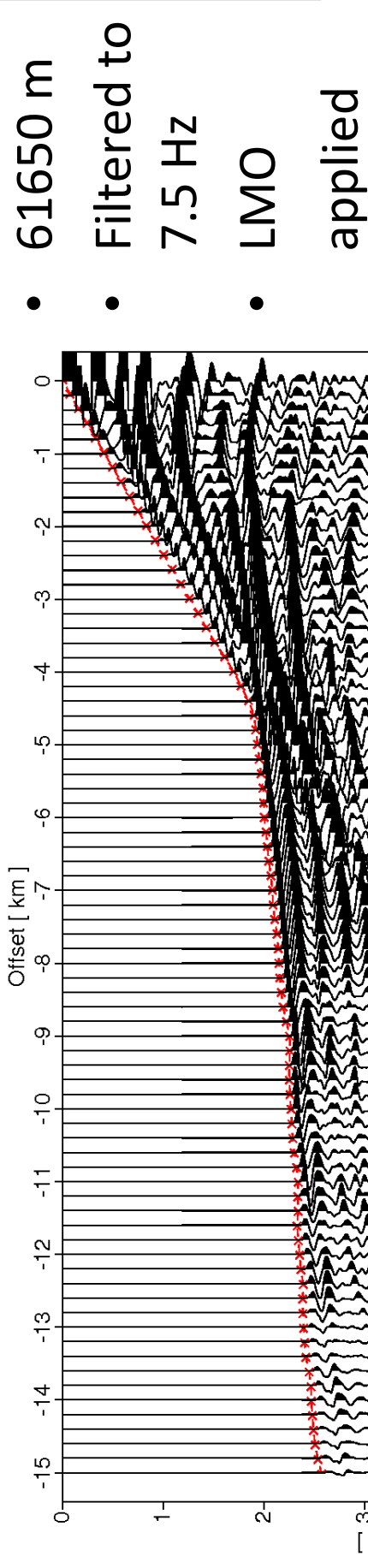
- 26450 m
- Filtered to 7.5 Hz
- LMO applied



# 2004 BP Velocity Benchmark: Synthetic Data



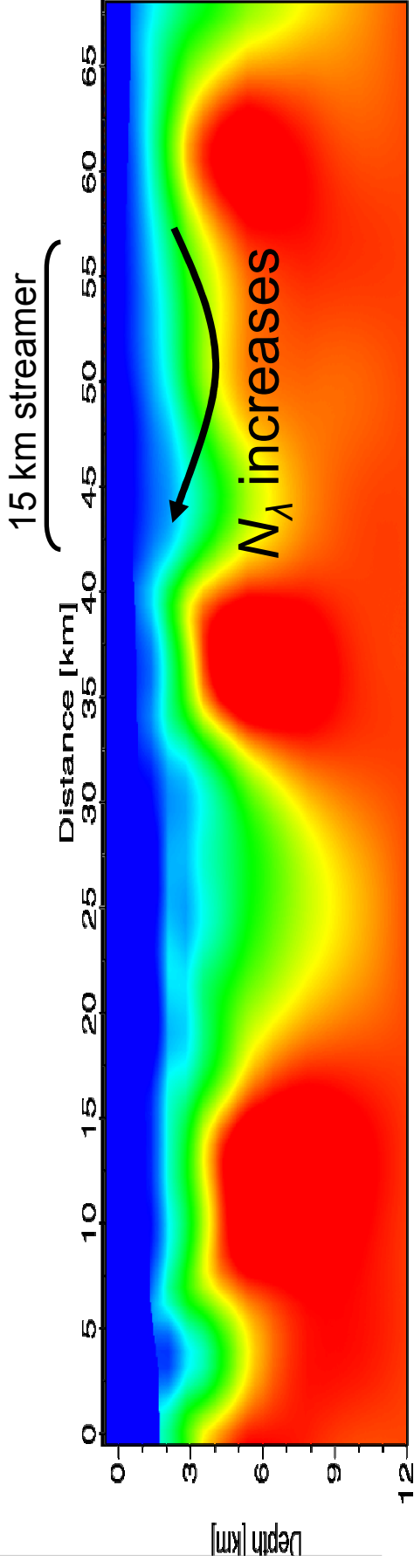
# 2004 BP Velocity Benchmark: Synthetic Data



# Outline

- Long offset seismic data
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# Forward modelling of wide angle propagation



Frequency 3 Hz

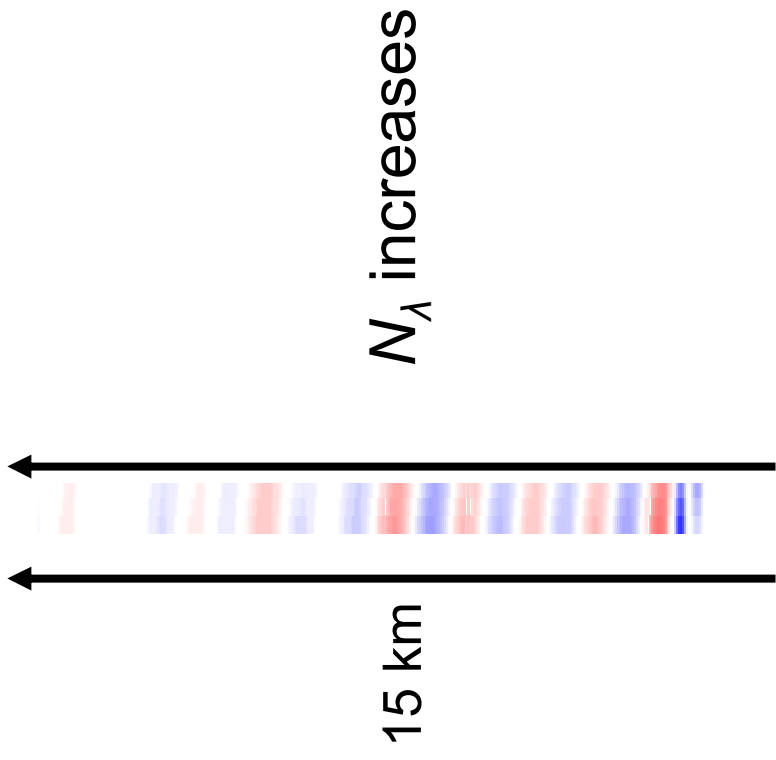
Wavelength  $\lambda \approx 1$  km

- 2D Acoustic, isotropic “2-way” wave equation
- Frequency-domain finite differences (implicit)

Animations courtesy Drew Breeders

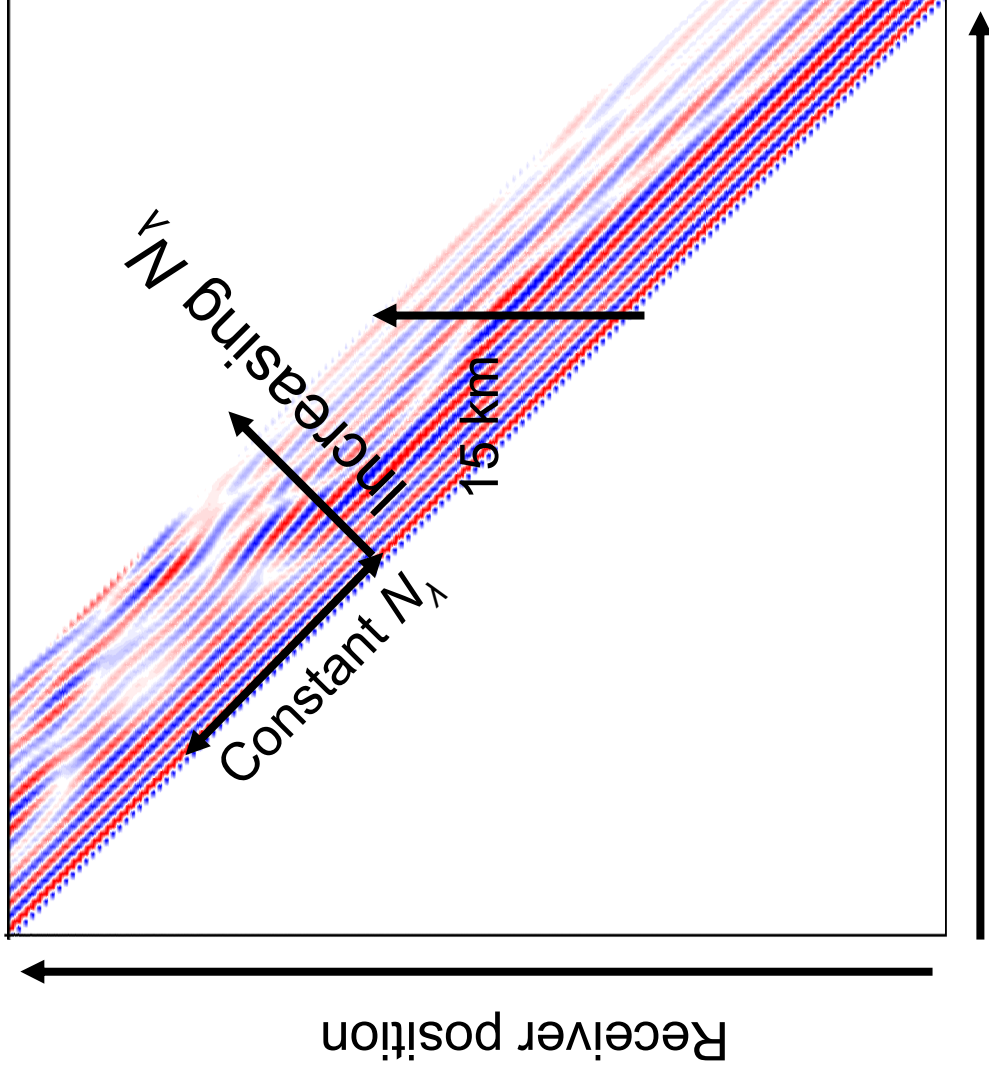
Western Science

# Forward propagated wavefield at receivers



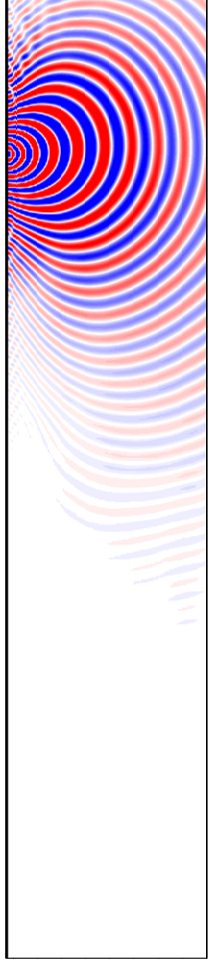
# Forward propagated wavefield at receivers:

All source locations





# Predicting data perturbations



Frequency 3 Hz

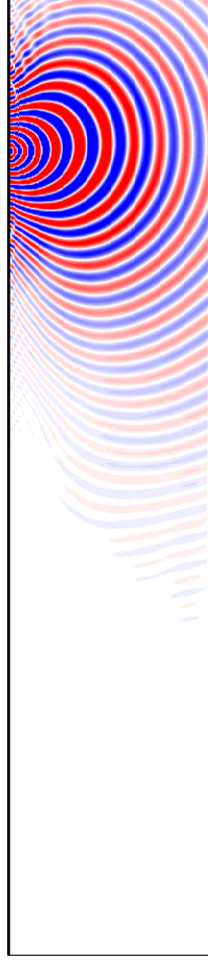
Wavelength  $\lambda \approx 1$  km

$$\delta d(\mathbf{r}_s, \mathbf{r}_g) = -\omega^2 \int_V G(\mathbf{r}, \mathbf{r}_s) \delta s^2(\mathbf{r}) G(\mathbf{r}, \mathbf{r}_g) d\mathbf{r}$$

# Predicting data perturbations

Frequency 3 Hz

Wavelength  $\lambda \approx 1$  km



Virtual source  
excitation

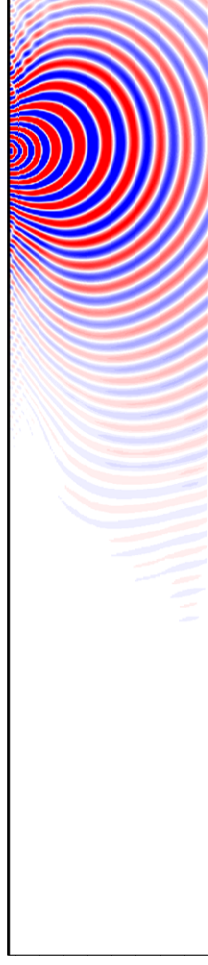


$$\delta d(\mathbf{r}_s, \mathbf{r}_g) = -\omega^2 \int_V G(\mathbf{r}, \mathbf{r}_s) \delta s^2(\mathbf{r}) G(\mathbf{r}, \mathbf{r}_g) d\mathbf{r}$$

# Predicting data perturbations

Frequency 3 Hz

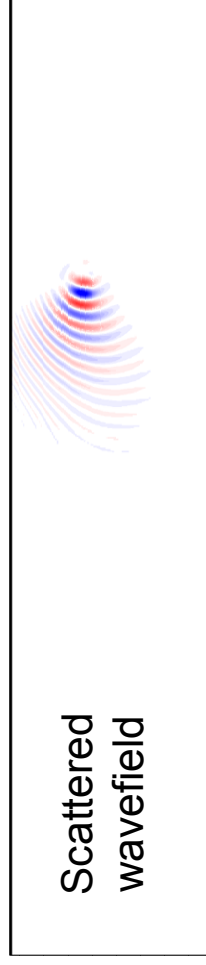
Wavelength  $\lambda \approx 1$  km



Virtual source  
excitation



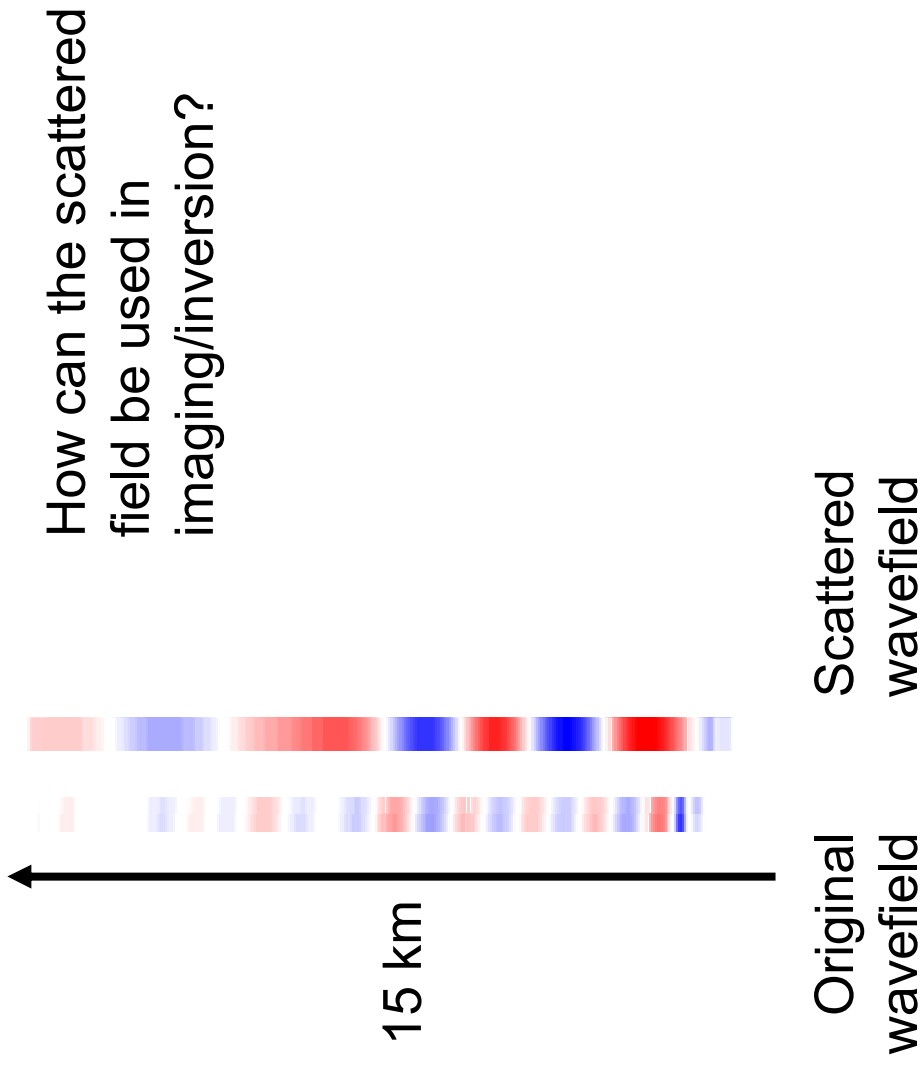
15 km streamer



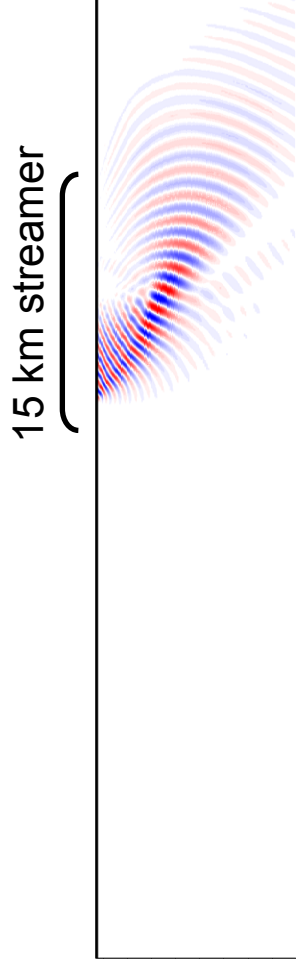
Scattered  
wavefield

$$\delta d(\mathbf{r}_s, \mathbf{r}_g) = -\omega^2 \int_V G(\mathbf{r}, \mathbf{r}_s) \delta s^2(\mathbf{r}) G(\mathbf{r}, \mathbf{r}_g) d\mathbf{r}$$

# Scattered wavefield at receivers



# The backpropagated wavefield



*Frequency* 3 Hz

*Wavelength*  $\lambda \approx 1$  km

Time reverse scattered wavefield, and propagate this as if it were a new source

This generates a partial focussing at the perturbation

How can we turn this intuitive idea into an image?

## Adjoint calculation

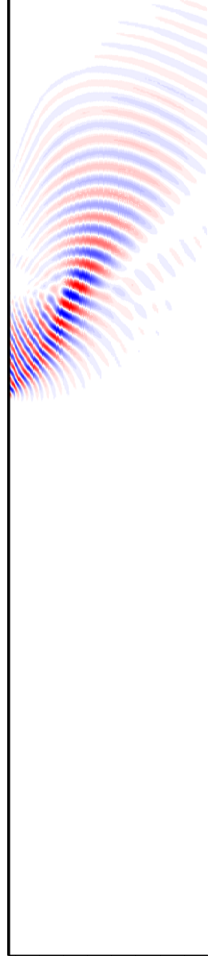
Born forward equation

$$\delta d(\mathbf{r}_s, \mathbf{r}_g) \approx -\omega^2 \int_V \delta s(\mathbf{r}) G(\mathbf{r}, \mathbf{r}_s) G(\mathbf{r}, \mathbf{r}_g) d\mathbf{r}$$

Gradient (update image)

$$\delta s(\mathbf{r}) = -\omega^2 \int_D G(\mathbf{r}, \mathbf{r}_s) \delta d(\mathbf{r}_s, \mathbf{r}_g) G(\mathbf{r}, \mathbf{r}_g) dD$$

# Adjoint calculation: Multiplication of forward and backpropagated wavefields



Frequency 3 Hz

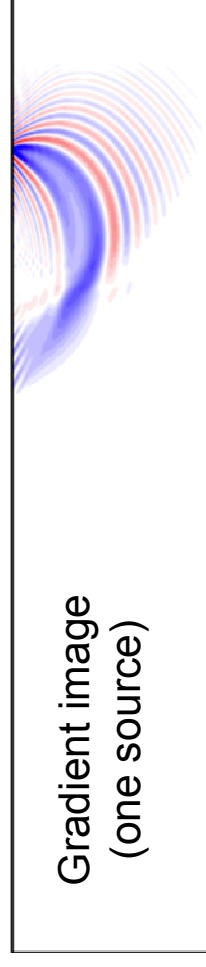
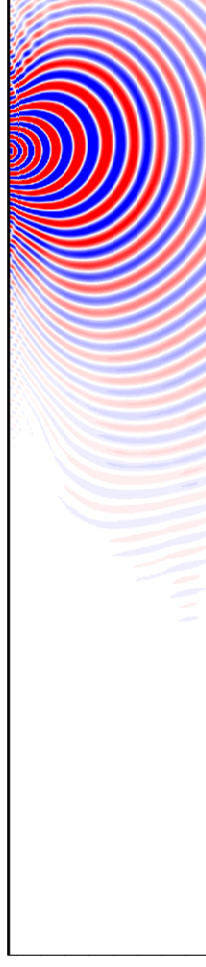
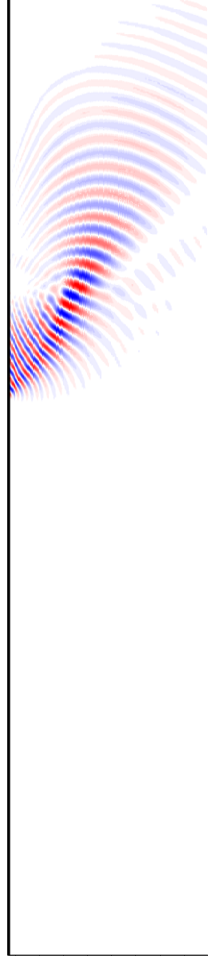
Wavelength  $\lambda \approx 1$  km

$$\widehat{\delta s^2}(\mathbf{r}) = -\omega^2 \int_{\mathcal{D}} G(\mathbf{r}, \mathbf{r}_s) \delta d(\mathbf{r}_s, \mathbf{r}_g) G(\mathbf{r}, \mathbf{r}_g) d\mathcal{D}$$

# Adjoint calculation: Multiplication of forward and backpropagated wavefields

Frequency 3 Hz

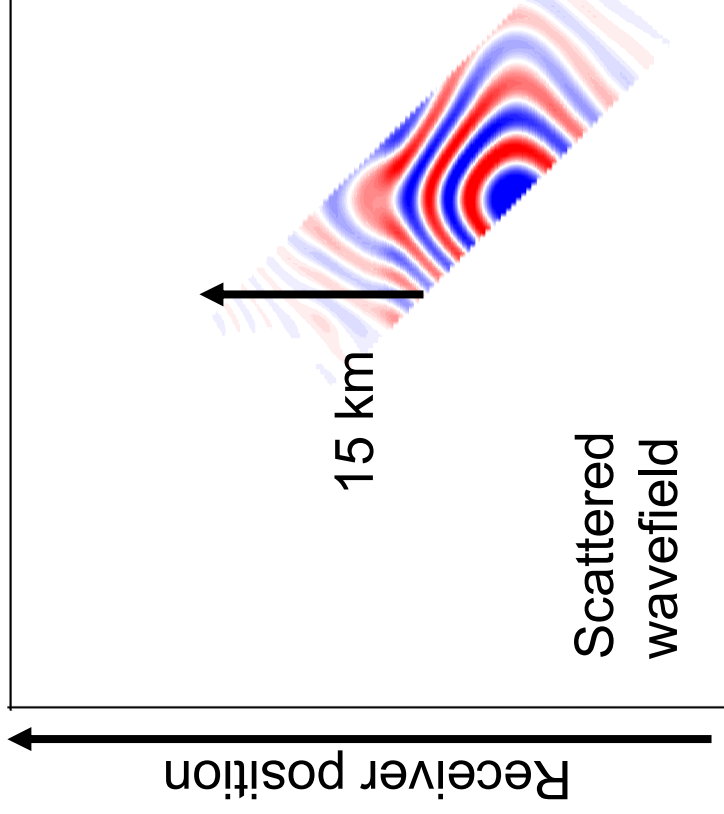
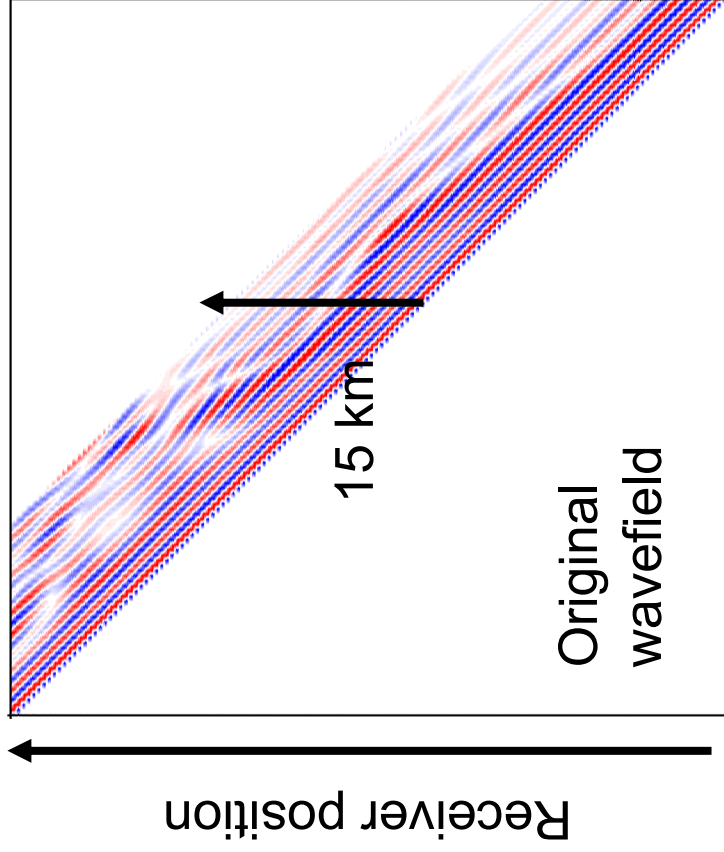
Wavelength  $\lambda \approx 1$  km



$$\widehat{\delta s^2(\mathbf{r})} = -\omega^2 \int_{\mathbf{D}} G(\mathbf{r}, \mathbf{r}_s) \delta d(\mathbf{r}_s, \mathbf{r}_g) G(\mathbf{r}, \mathbf{r}_g) d\mathbf{D}$$



# Scattered wavefield at receivers: All source locations



# Adjoint calculation, many sources

Frequency 3 Hz

Wavelength  $\lambda \approx 1$  km



$$\widehat{\delta s^2}(\mathbf{r}) = -\omega^2 \int_{\mathcal{D}} G(\mathbf{r}, \mathbf{r}_s) \delta d(\mathbf{r}_s, \mathbf{r}_g) G(\mathbf{r}, \mathbf{r}_g) d\mathcal{D}$$

# Waveform tomography: An iterative process

Frequency 3 Hz

Wavelength  $\lambda \approx 1$  km



$$\widehat{\delta s^2}(\mathbf{r}) = -\omega^2 \int_{\mathbf{D}} G(\mathbf{r}, \mathbf{r}_s) \delta d(\mathbf{r}_s, \mathbf{r}_g) G(\mathbf{r}, \mathbf{r}_g) d\mathbf{D}$$

# Outline

- Long offset seismic data
  - An imaging opportunity
- **Waveform Tomography**
  - Theoretical aspects
  - **Making it work**
- Nankai Trough Subduction Zone Example
  - Geological implications
  - The Challenge of High Resolution

# Practical Aspects

Waveform inversion:

- High resolution / highly non-linear / computationally costly

Seek “robustness” and computational efficiency

- Low frequencies – robust spectral components
- Phase – robust spectral attributes
- Early arrivals – robust temporal wavefield

Leads to consideration of

- Phase-only objective function
- Acoustic formulation

# Practical Aspects

FWI – the “Grand vision” of Tarantola et al:

- Models would predict the full wavefield
- Models would respect full elastic wave propagation physics
- Models would reveal full information on all elastic parameters (including insight into posteriori distributions)

The “phase we are going through”

- Bootstrap from kinematic starting models (velocity analysis, traveltimes)
- Acoustic formulation – phase-only
- Acoustic formulation – phase & amplitude
- Elastic formulation – amplitude and phase



# Practical Aspects

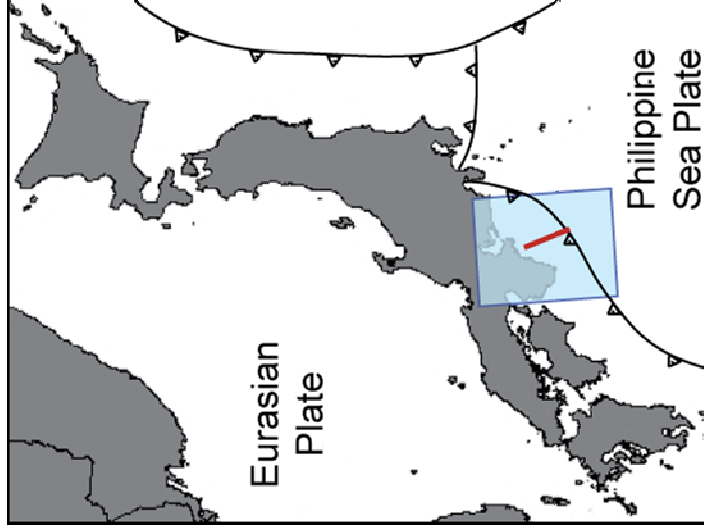
- Bootstrap from kinematic starting models (velocity analysis, traveltimes)
- Acoustic formulation – phase-only
- Acoustic formulation – phase & amplitude
- Elastic formulation – amplitude and phase



# Data sets

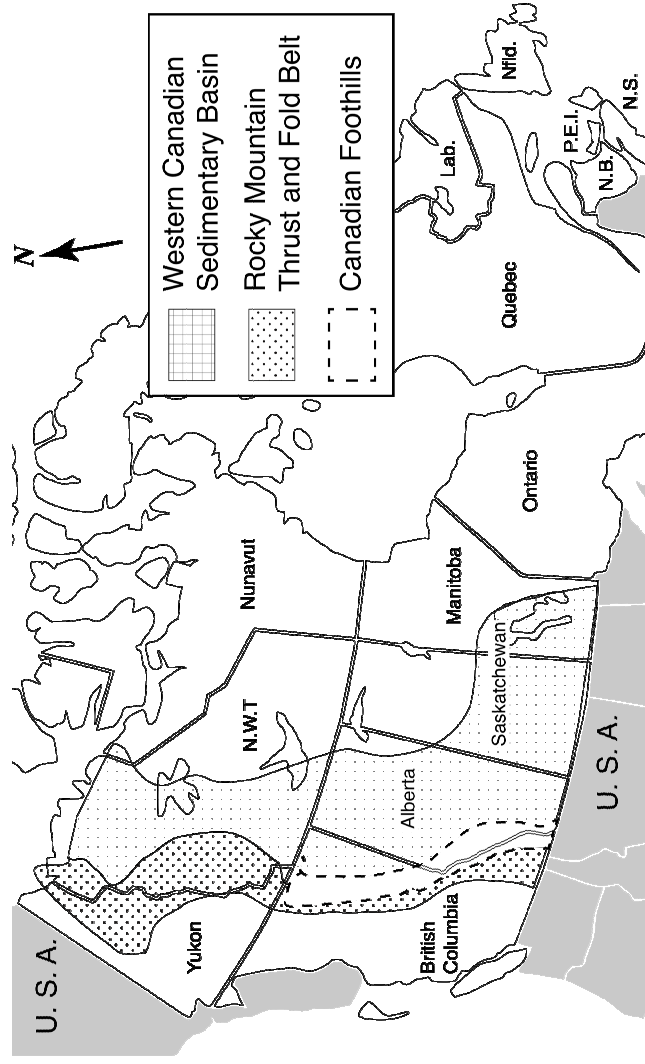
Marine

**Nankai subduction zone**  
Crustal scale imaging



Land

**Canadian Foothills**  
Hydrocarbon exploration



Kamei and Pratt, 2010, SEG

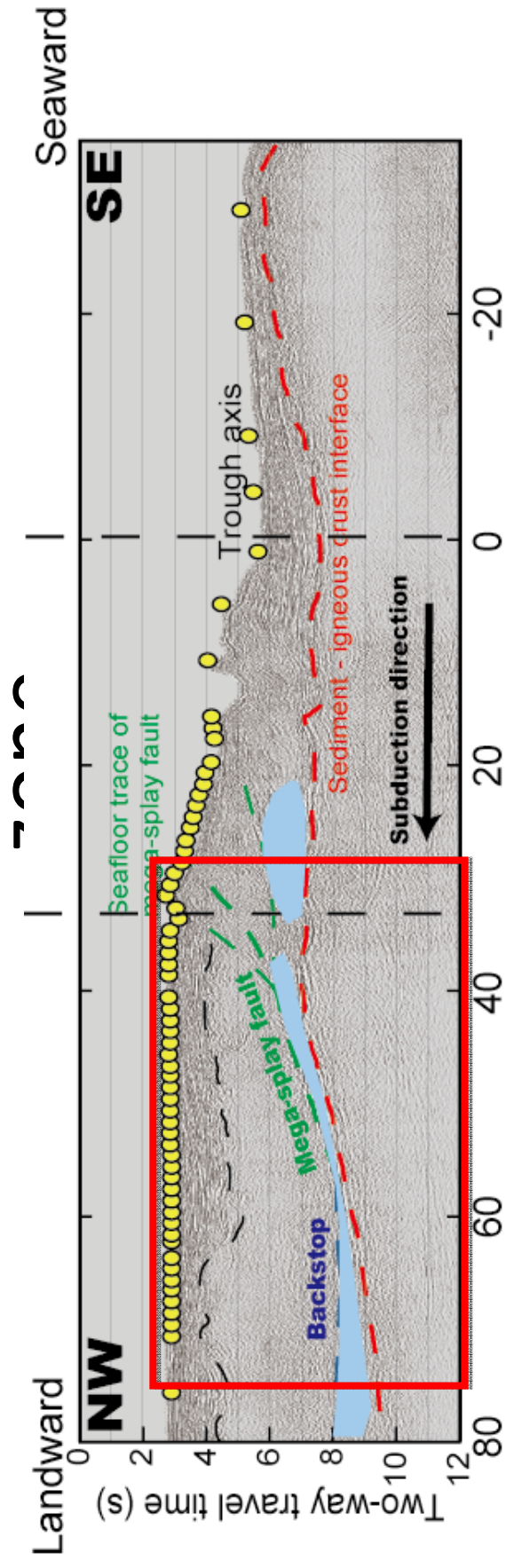
Wameter et al., 2011, IHPST, Under review

Brenders et al., 2009, 2010, SEG

Brenders, 2011, PhD Thesis



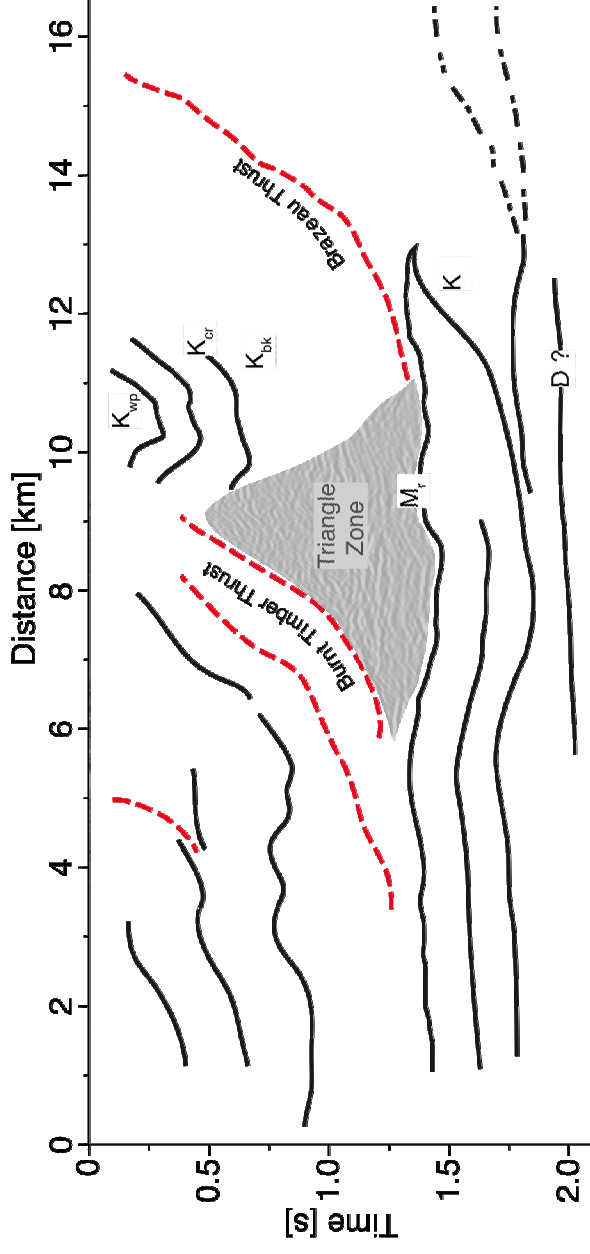
# Marine data set: Nankai subduction



(Courtesy of T. Tsuji)

# of OBSs	54 (2 km water depth)
OBS interval	1 km
# of shots	285
Shot interval	200 m (10 m towed depth)
Max offset	60 km
Min available freq	2 Hz

# Land data set: Canadian foothills



# of channels	666 (live)
Receiver Interval	25 m
# of shots	168
Shot interval	100 m
Max offset	16.6 km
Min available freq	4 Hz

# Waveform inversion strategy

- Acoustic forward modeling and P-wave imaging
- Data preprocessing
  - Bandpass filter
  - Top and bottom mute
  - Amplitude scaling
  - Choice of frequency and damping schedule
  - Extraction of Laplace-Fourier components from the data

# Waveform inversion strategy

$$2E = \sum_{\omega} \delta \mathbf{d}^T \delta \mathbf{d},$$

- Minimize  $L_2$  norm:
  - by conjugate gradient method
- Multiscale approach I: Data preconditioning
  - Selection of offset range(s)
  - Sequential inversion of a set of frequencies
  - Laplace-Fourier domain (Time damping)
- Multiscale approach II: Image preconditioning
  - Wavenumber filtering and/or smoothing

# Waveform inversion strategy

$$2E = \sum_{\omega} \delta \mathbf{d}^T \delta \mathbf{d},$$

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  - Selection of offset range(s)
  - Sequential inversion of a set of frequencies
  - Laplace-Fourier domain (Time damping)
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  - Wavenumber filtering and/or smoothing

# Objective functions



$$2E = \sum_{\omega} \delta \mathbf{d}^T \delta \mathbf{d},$$

The way in which we define the data residuals is critical

# Discussion on objective functions....

## Conventional residual

$$\delta d_j = u_j - d_j$$

predicted   observed

## Phase-only residual

$$\begin{aligned} \delta d_j &= \Im \left[ \ln \left( \frac{u_j}{d_j} \right) \right] \\ &= \arg(u_j) - \arg(d_j) \end{aligned}$$

# Discussion on objective functions....

## Conventional residual

$$\delta d_j = u_j - d_j$$

predicted
observed

$$10^{-5} < |\delta d_j| \approx |d_j| < 10$$

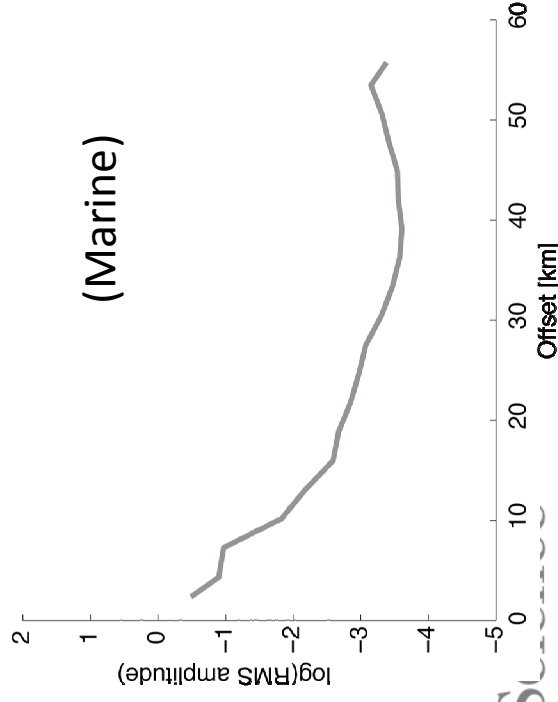
## Phase-only residual

$$\delta d_j = \Im \left[ \ln \left( \frac{u_j}{d_j} \right) \right]$$

$$= \arg(u_j) - \arg(d_j)$$

$$-\pi < |\delta d_j| < \pi$$

Dynamic range





# Discussion on objective functions....

## Conventional residual

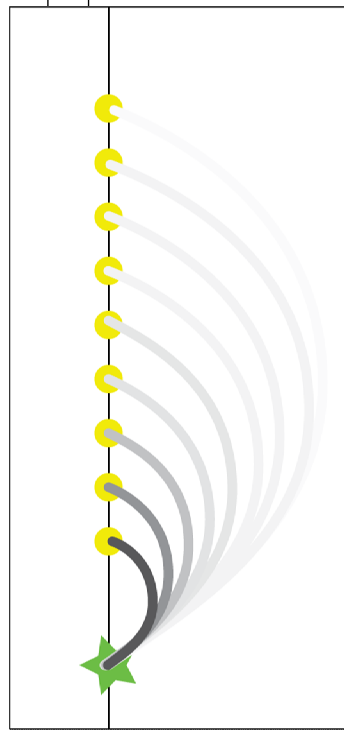
$$\delta d_j = u_j - d_j$$

predicted
observed

Dynamic range

$$10^{-5} < |\delta d_j| \approx |d_j| < 10$$

Effective offset



narrow

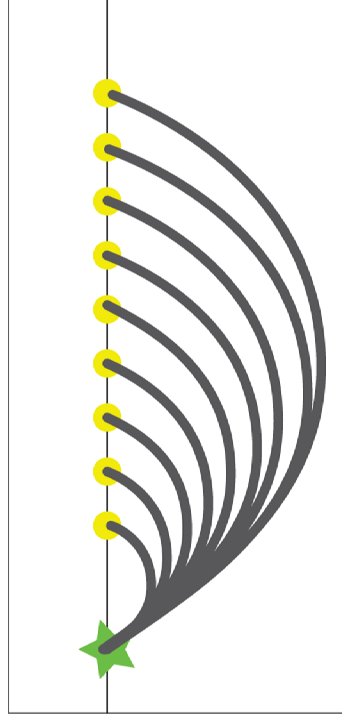
## Phase-only residual

$$\delta d_j = \Im \left[ \ln \left( \frac{u_j}{d_j} \right) \right]$$

$$= \arg(u_j) - \arg(d_j)$$

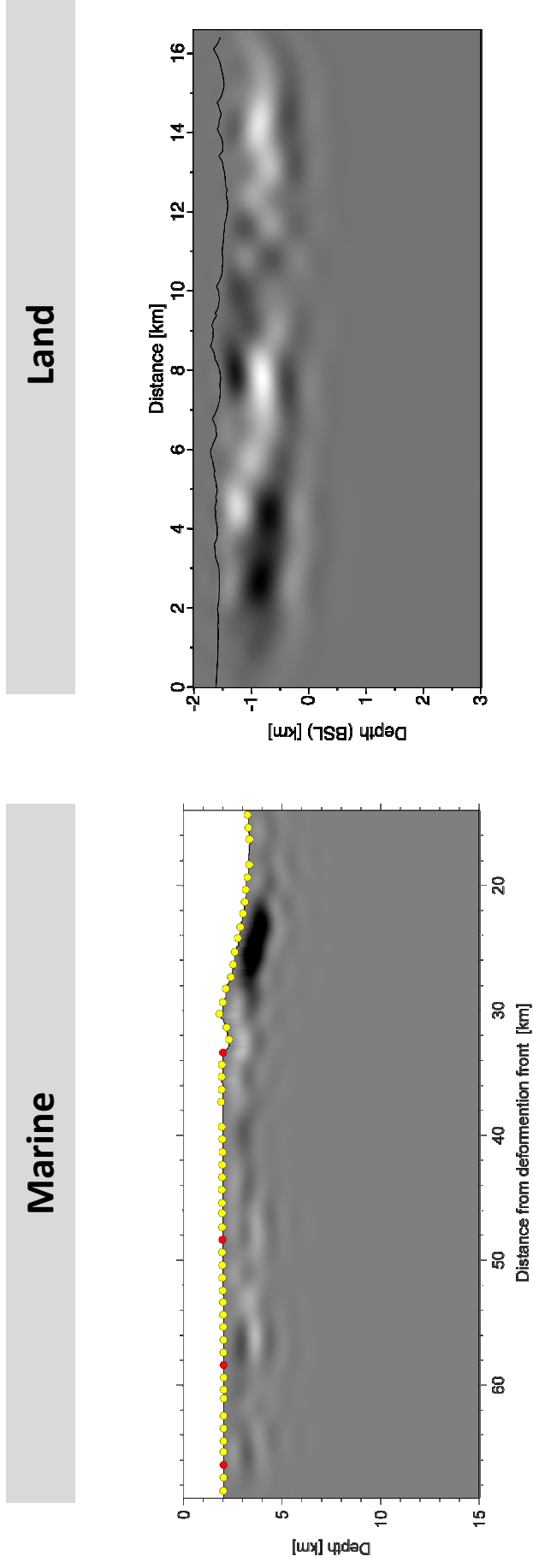
$$-\pi < |\delta d_j| < \pi$$

wide (entire)



# Gradient: Conventional objective function

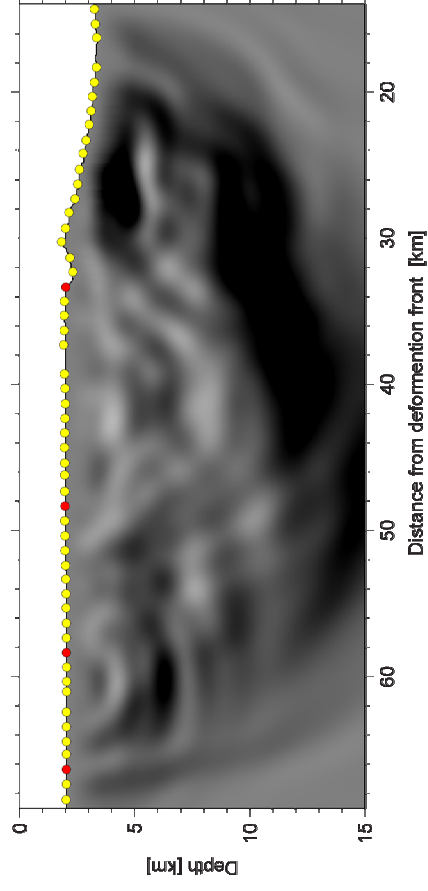
$$\delta d_j = u_j - d_j$$



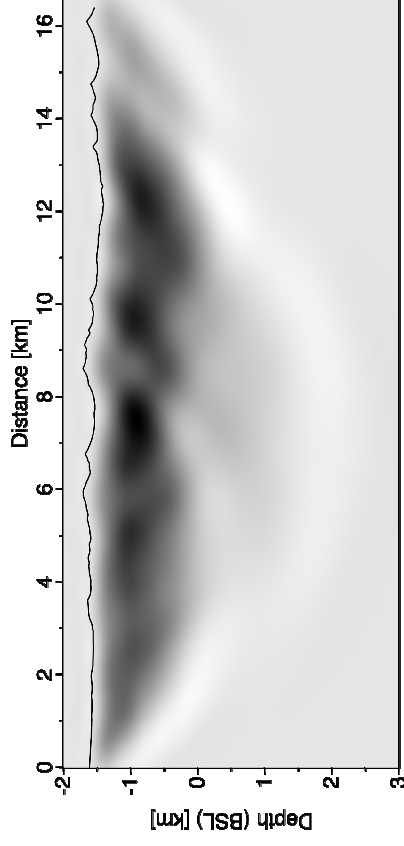
# Gradient: Phase-only objective function

$$\delta d_j = \Im \left[ \ln \left( \frac{u_j}{d_j} \right) \right]$$

Marine



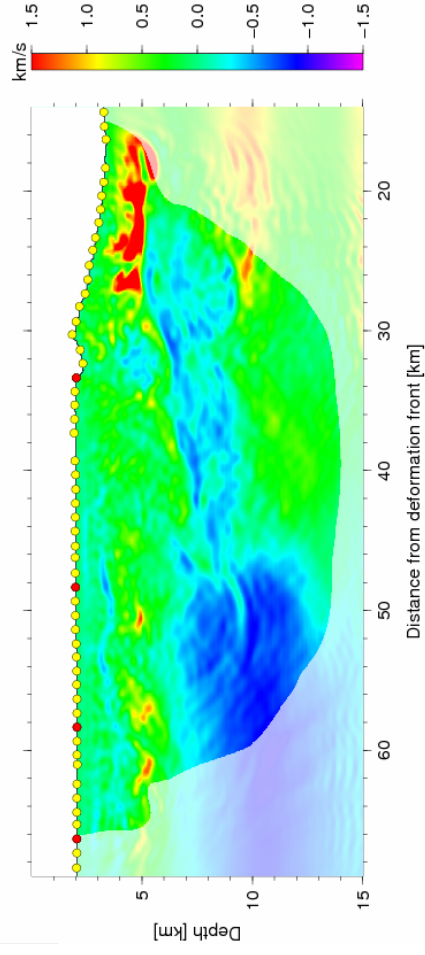
Land



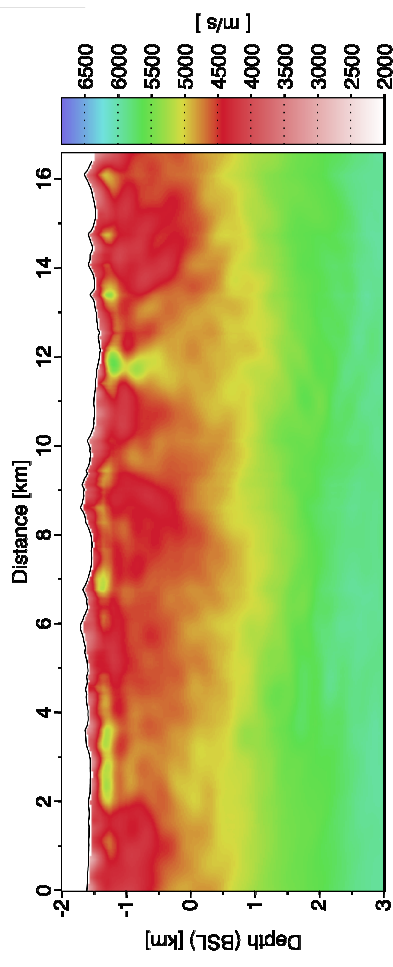
# Results: Conventional objective function

$$\delta d_j = u_j - d_j$$

Marine (2.25 – 8.5 Hz)



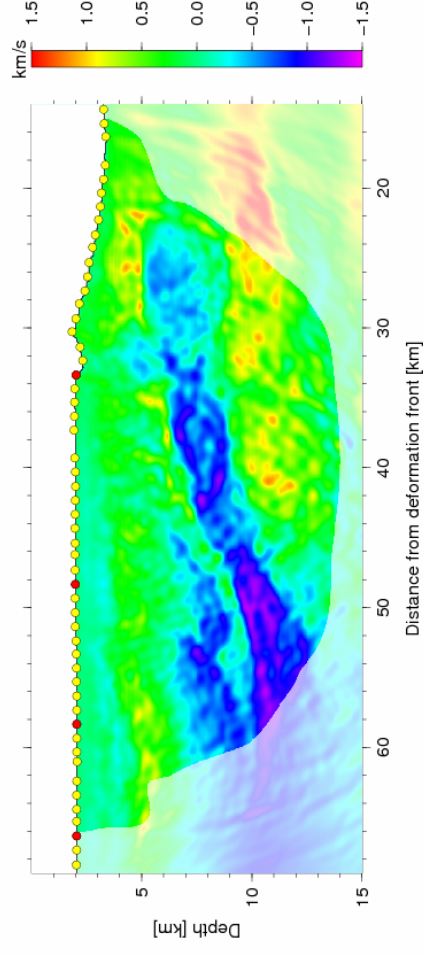
Land (4.0 – 12.0 Hz)



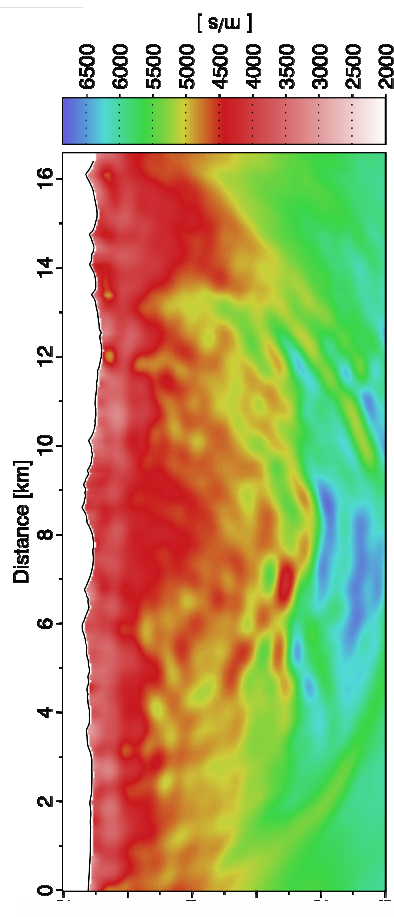
# Results: Phase-only

$$\delta d_j = \Im \left[ \ln \left( \frac{u_j}{d_j} \right) \right]$$

Marine (2.25 – 8.5 Hz)



Land (4.0 – 12.0 Hz)



“Phase-only” – are we giving up on amplitudes?

The “phase” information is much richer than  
simply arrival time information

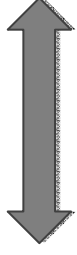
Complex wavefields with overlapping events  
have complex phase behaviour

**Relative** amplitude of events in the data govern  
the phase behaviour

# Laplace-Fourier approach

**Time damping**  
(Sirgue and Pratt, 2004)

**Laplace-Fourier approach**  
(Shin and Cha, 2009)



$$u(t) \rightarrow u(t) \exp(-t/\tau)$$

$$\tau = 1/s$$

$$u(t) \rightarrow u(t) \exp(-st)$$

$$\int u(t) \exp(-t/\tau) \exp(-i\omega t) dt$$

$$= \int u(t) \exp[-i(\omega + i/\tau)t] dt$$

$$= u(\omega + i/\tau)$$



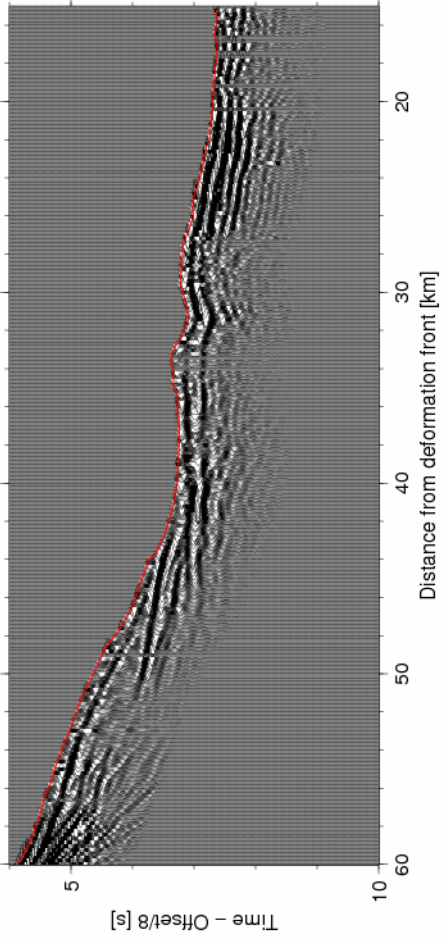
$$u(\omega) \rightarrow u(\omega + i/\tau)$$

$$u(\omega) \rightarrow u(\omega + is)$$

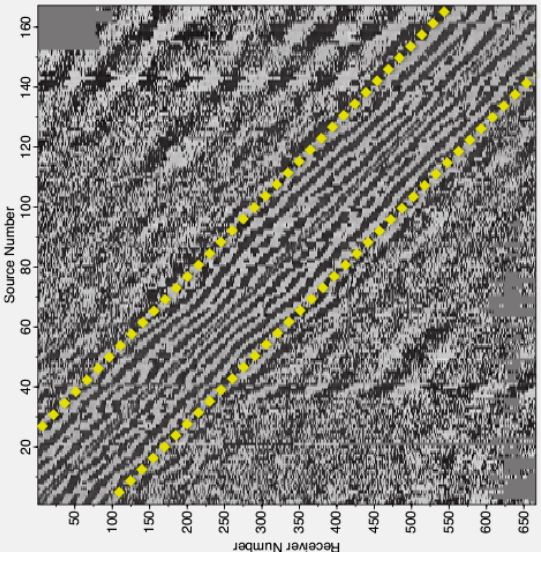
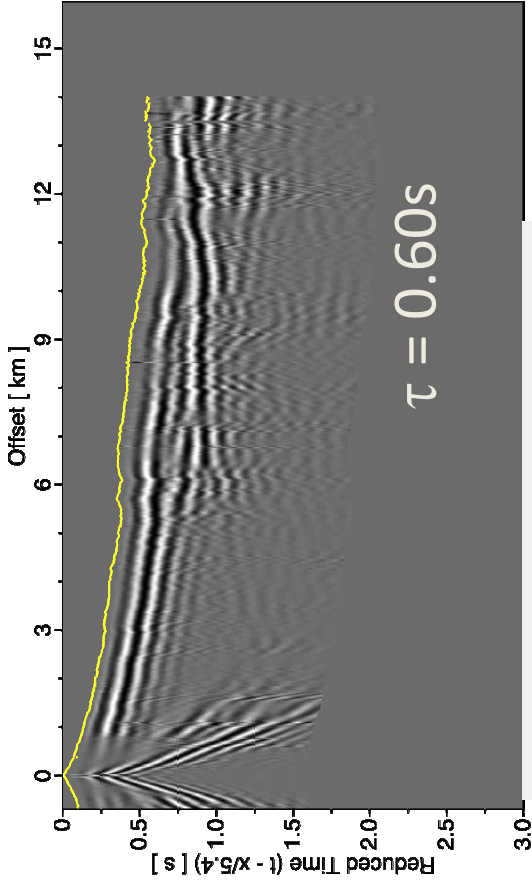
□

# Examples of time damping; damping 1

Marine



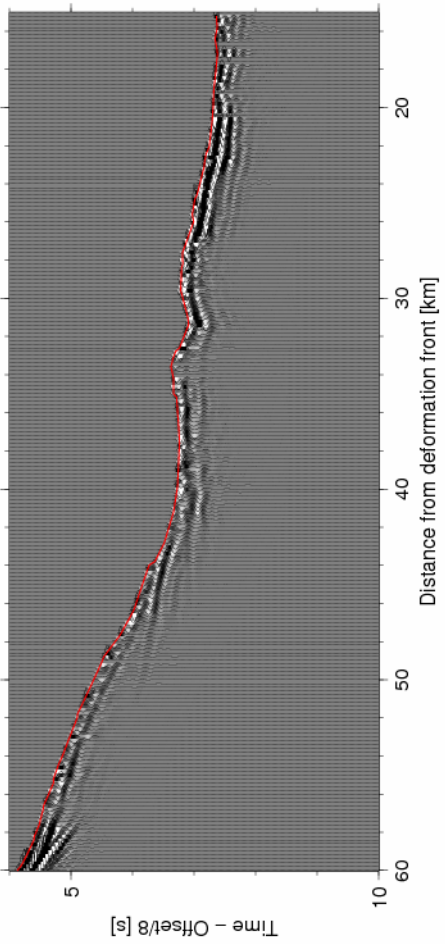
Land



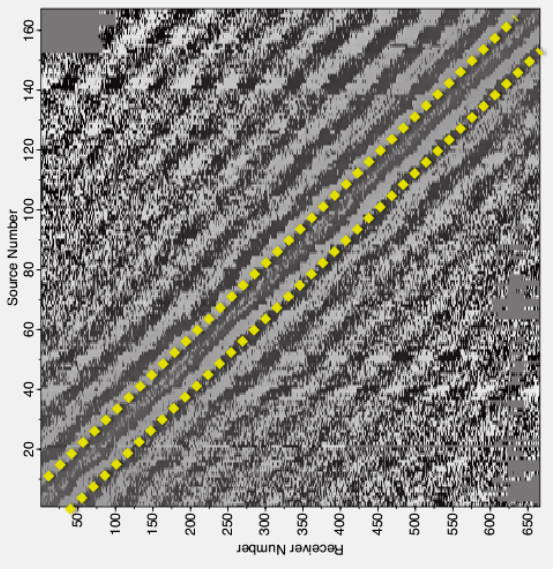
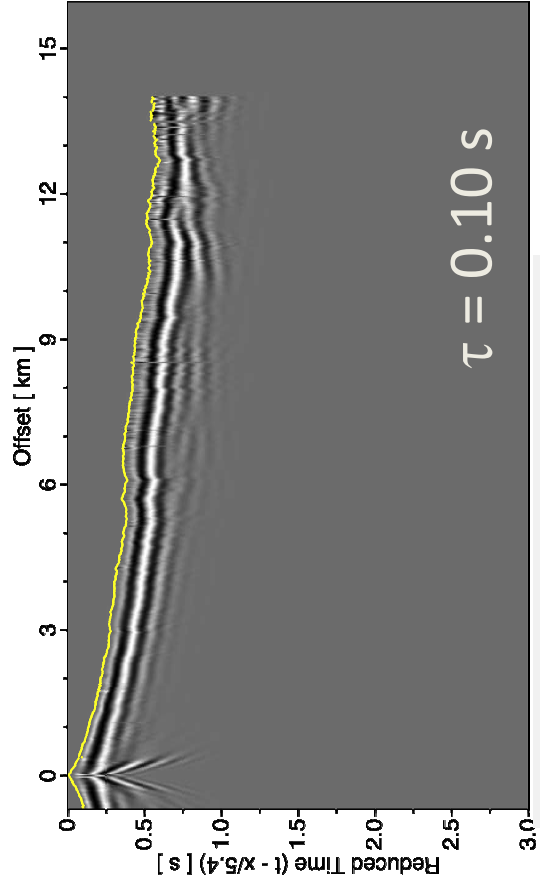


# Examples of time damping; damping 2

Marine

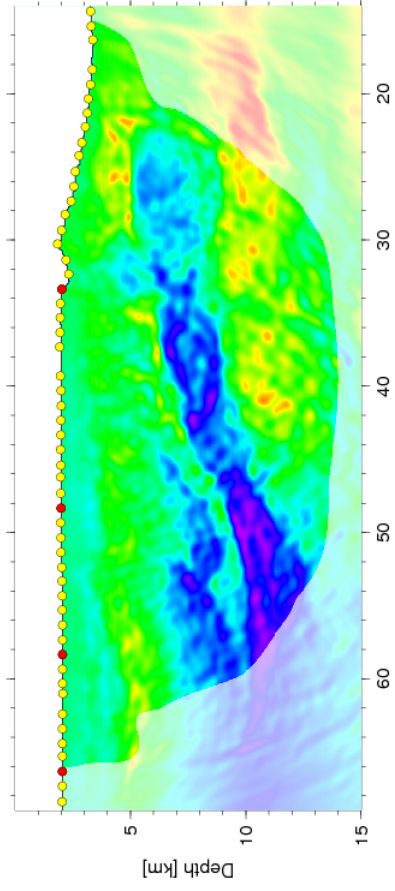


Land

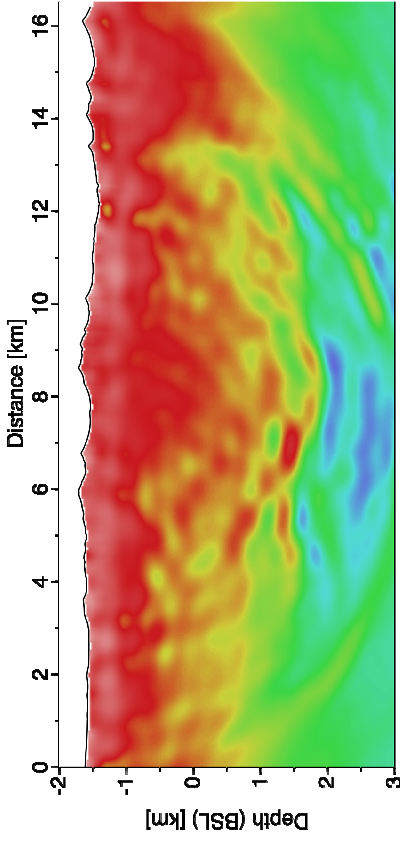


# Models and waveform fit

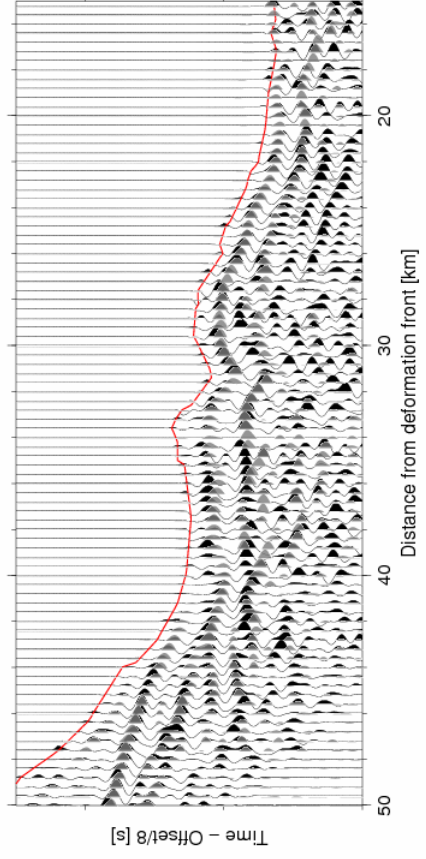
Marine



Land



Distance from deformation front [km]



Offset [km]

Reduced Time  $(t - x/5.4)$  [s]

Shot Position .686 km

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