



Towards global scale full-waveform inversion

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May 2012, Slovakia**

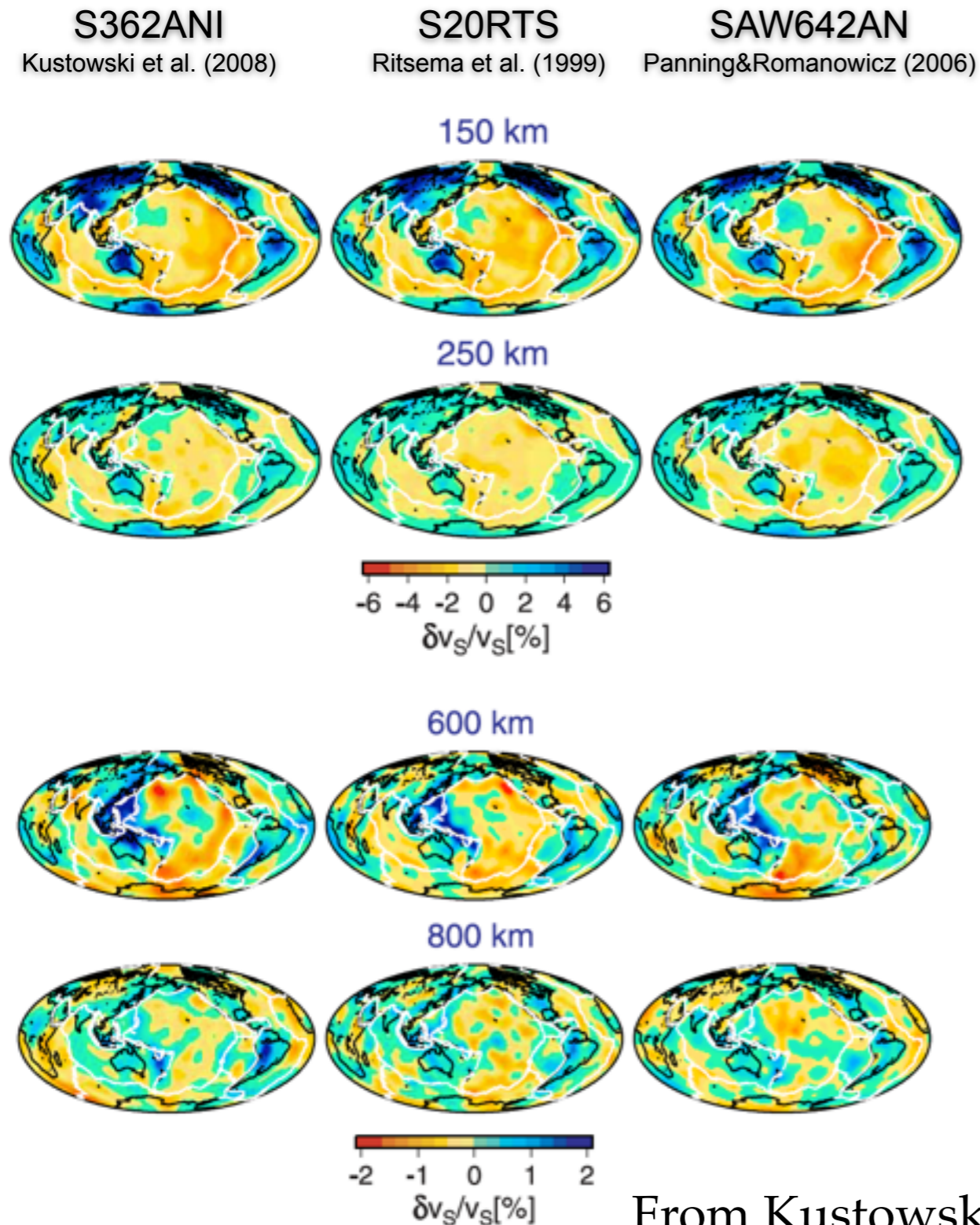


What we mean by full-waveform tomography

- Forward simulations in 3D models
- Fréchet kernels in 3D background models
- Use of complete seismograms at three components
- Use of both phase and amplitudes

Towards global adjoint tomography

Global tomography



From Kustowski et al. (2008)

- mostly based on ray theory, recently finite-frequency effects are also taken into account
- 1D background models
- Combination of different data sets
- Crustal corrections

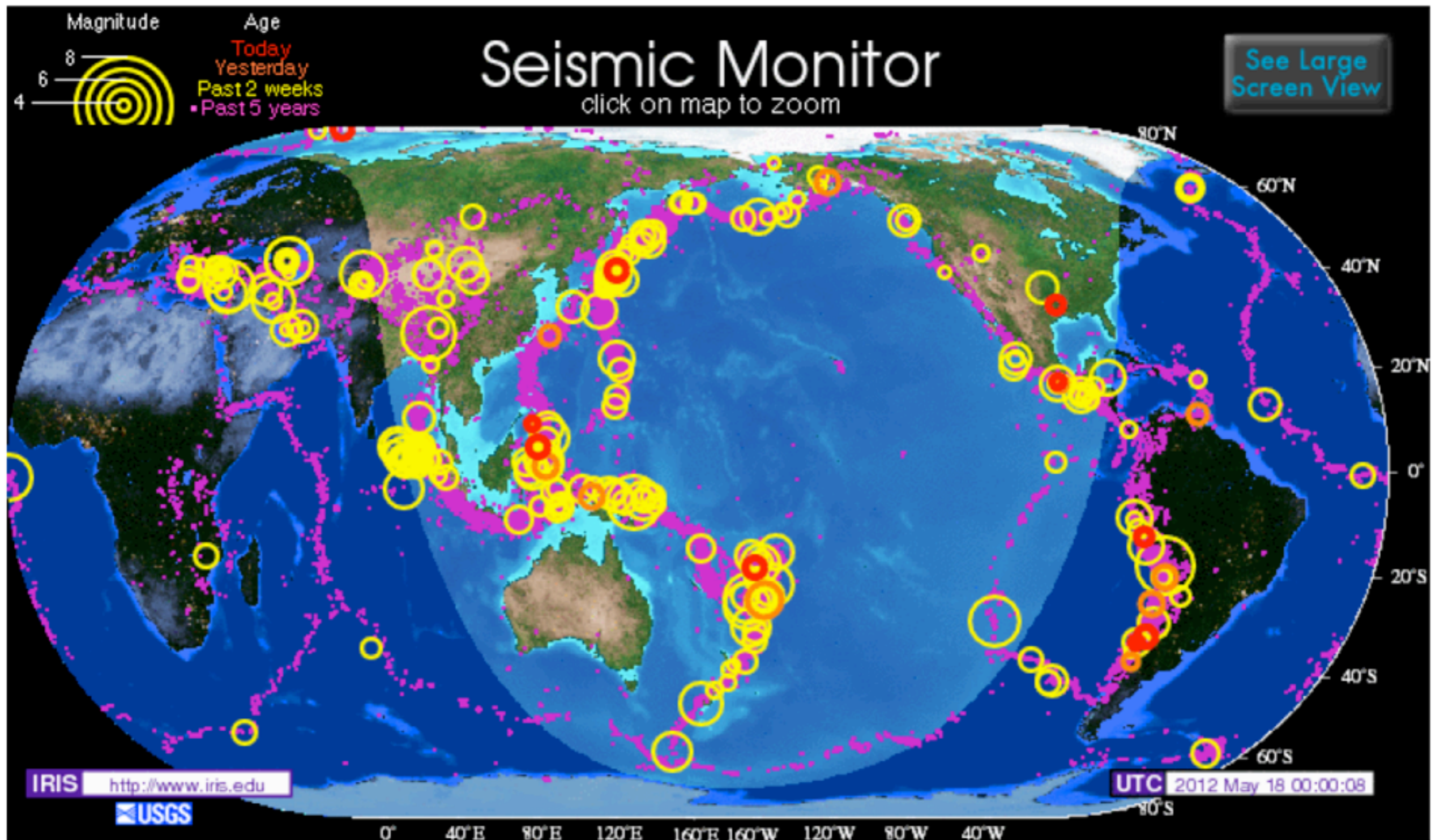
Challenges in global tomography

- Theoretical limitations
 - Finite-frequency effects have become important

Challenges in global tomography

- Theoretical limitations
 - Finite-frequency effects have become important
- Data coverage
 - Uneven distribution of earthquakes and stations on the globe

World seismicity



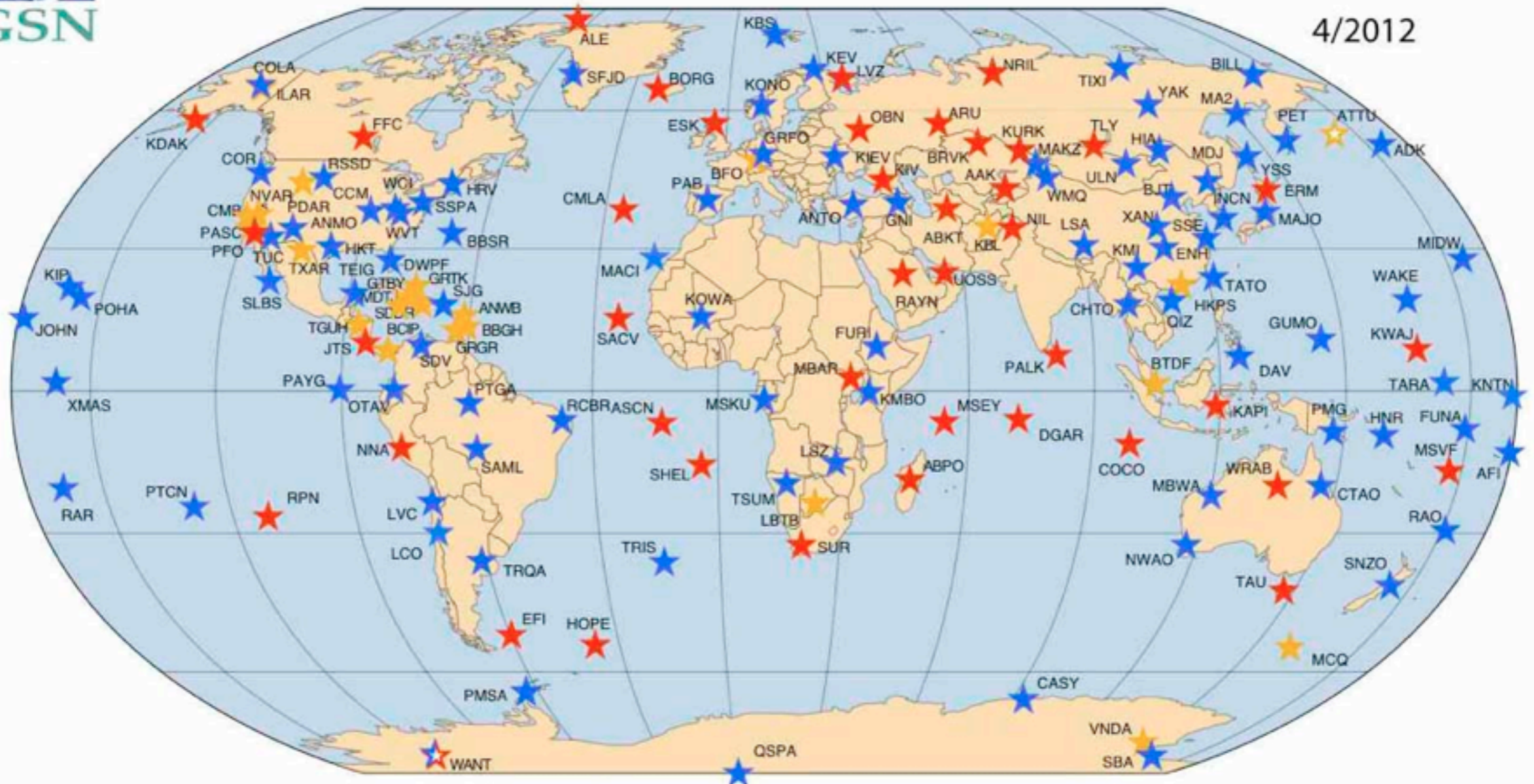
<http://www.iris.edu/dms/seismon.htm>

Seismic stations



GLOBAL SEISMOGRAPHIC NETWORK

4/2012



★ IRIS / IDA Stations

★ IRIS / USGS Stations

★ Affiliate Stations

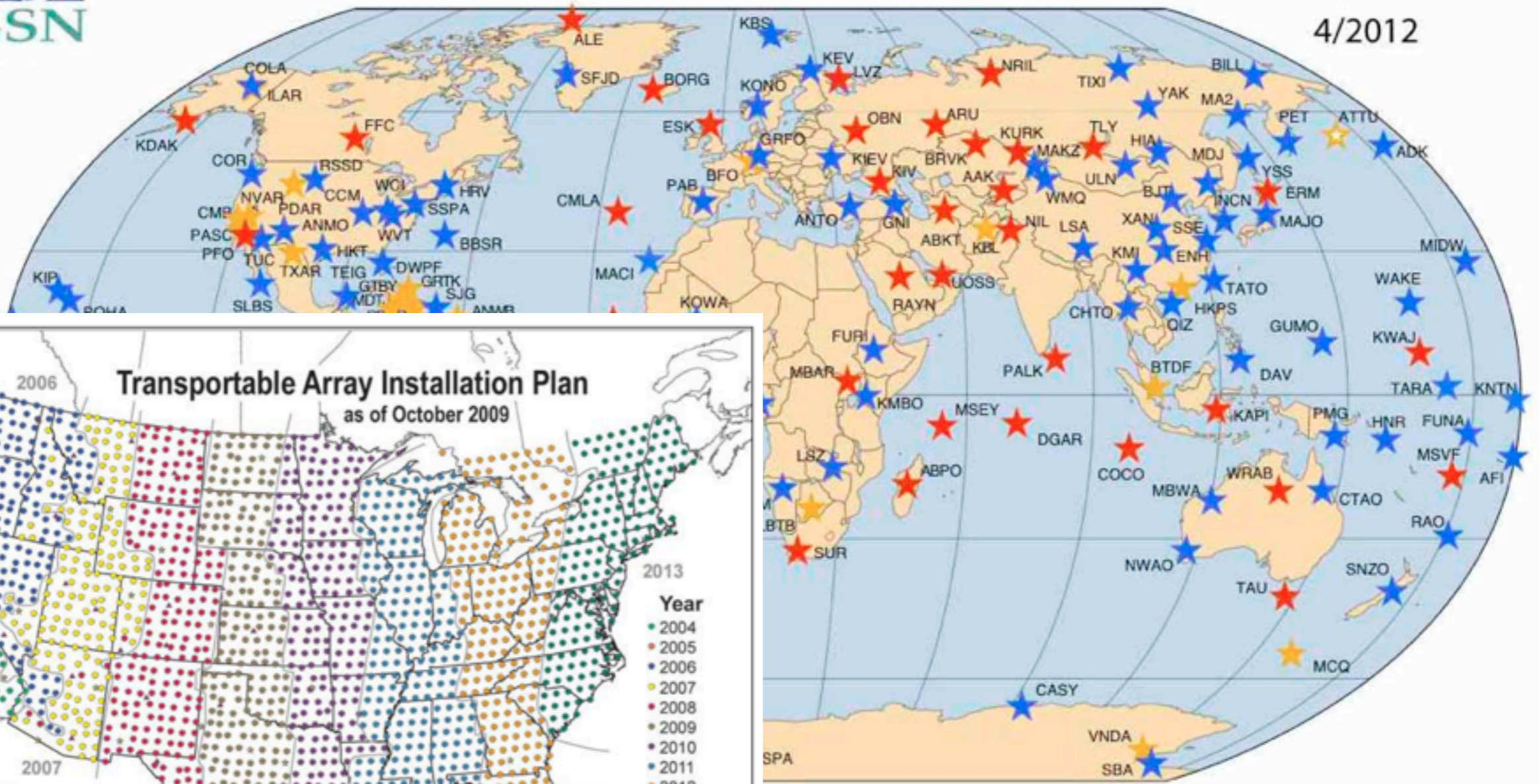
★ Planned Stations

Seismic stations



GLOBAL SEISMOGRAPHIC NETWORK

4/2012

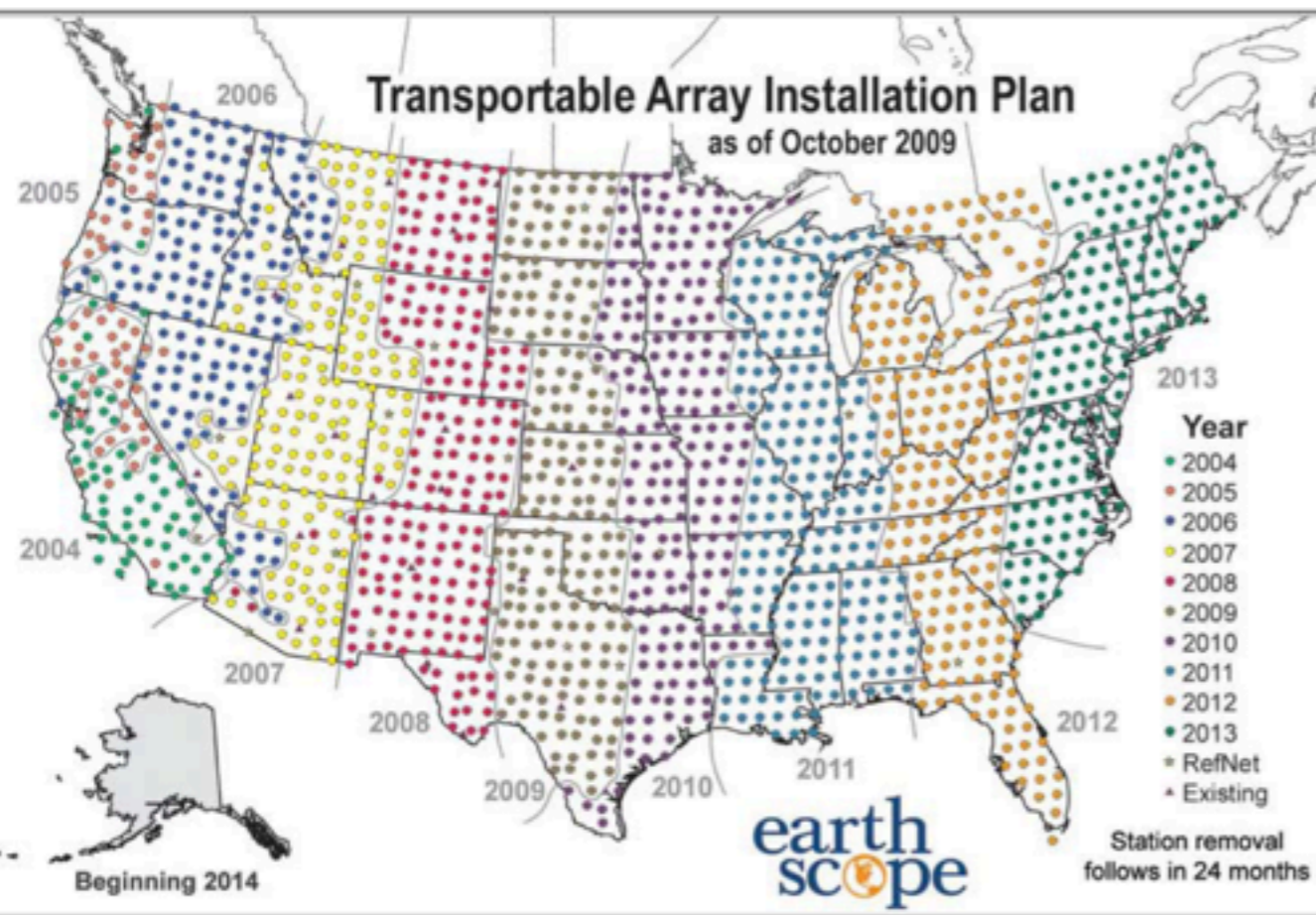


JSGS Stations



Affiliate Stations

Transportable Array Installation Plan as of October 2009



Seismic stations

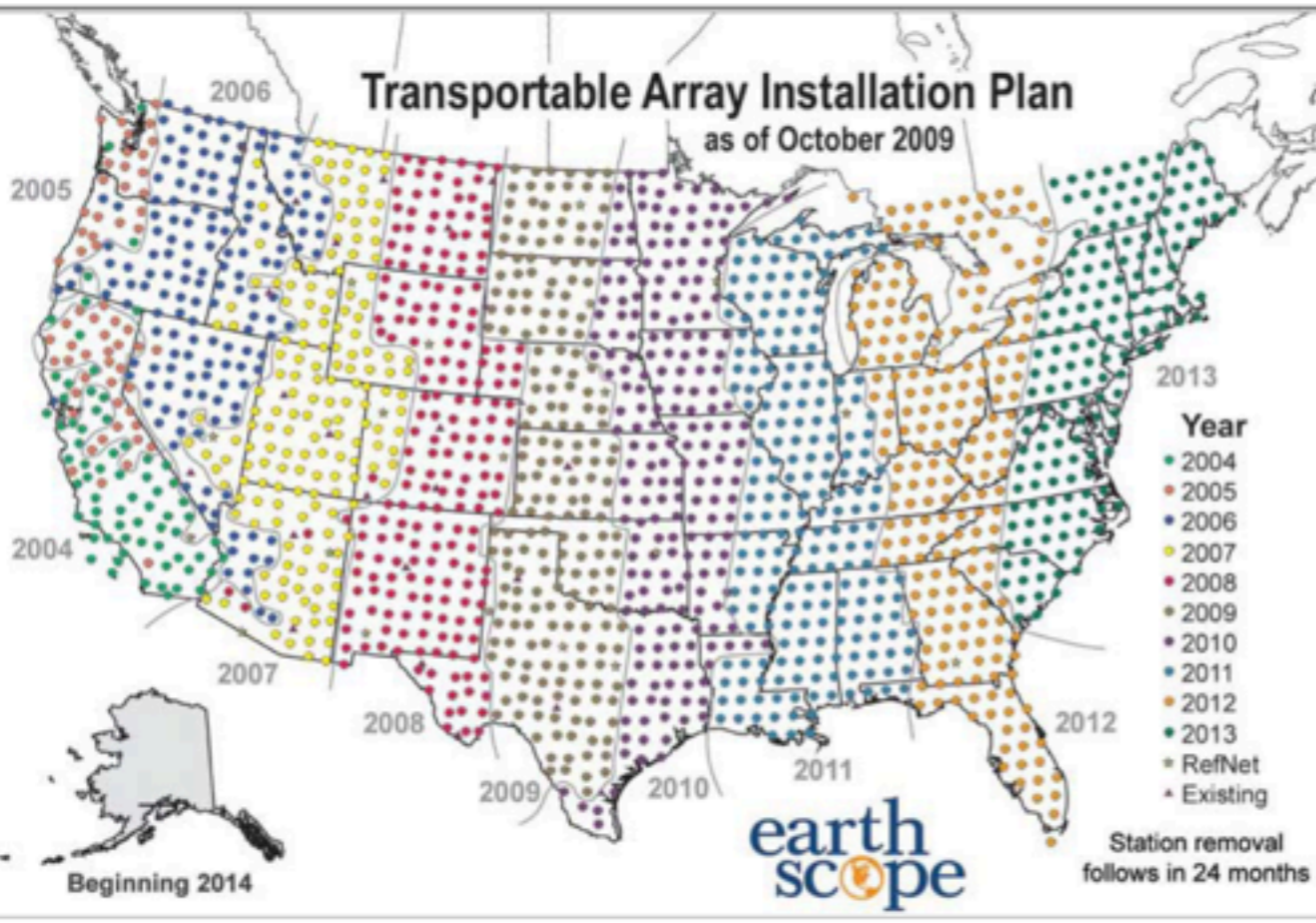


GLOBAL SEISMOGRAPHIC NETWORK

4/2012



Transportable Array Installation Plan as of October 2009



Simons et al. 2009



Challenges in global tomography

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- Data coverage
 - Uneven distribution of earthquakes and stations on the globe
 - Usable data is subjected to the forward theory

Challenges in global tomography

- Theoretical limitations
 - Finite-frequency effects have become important
- Data coverage
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 - Usable data is subjected to the forward theory
- Crustal effects
 - Can be highly nonlinear, thus “crustal corrections” are questionable

3D wave simulations - Adjoint tomography

- Full nonlinearity of wave propagation
- Dramatic increase in usable data, resulting better data coverage
- 3D background models help reduce nonlinearity of problem
- Iterative update of models
- No crustal corrections!

Outline

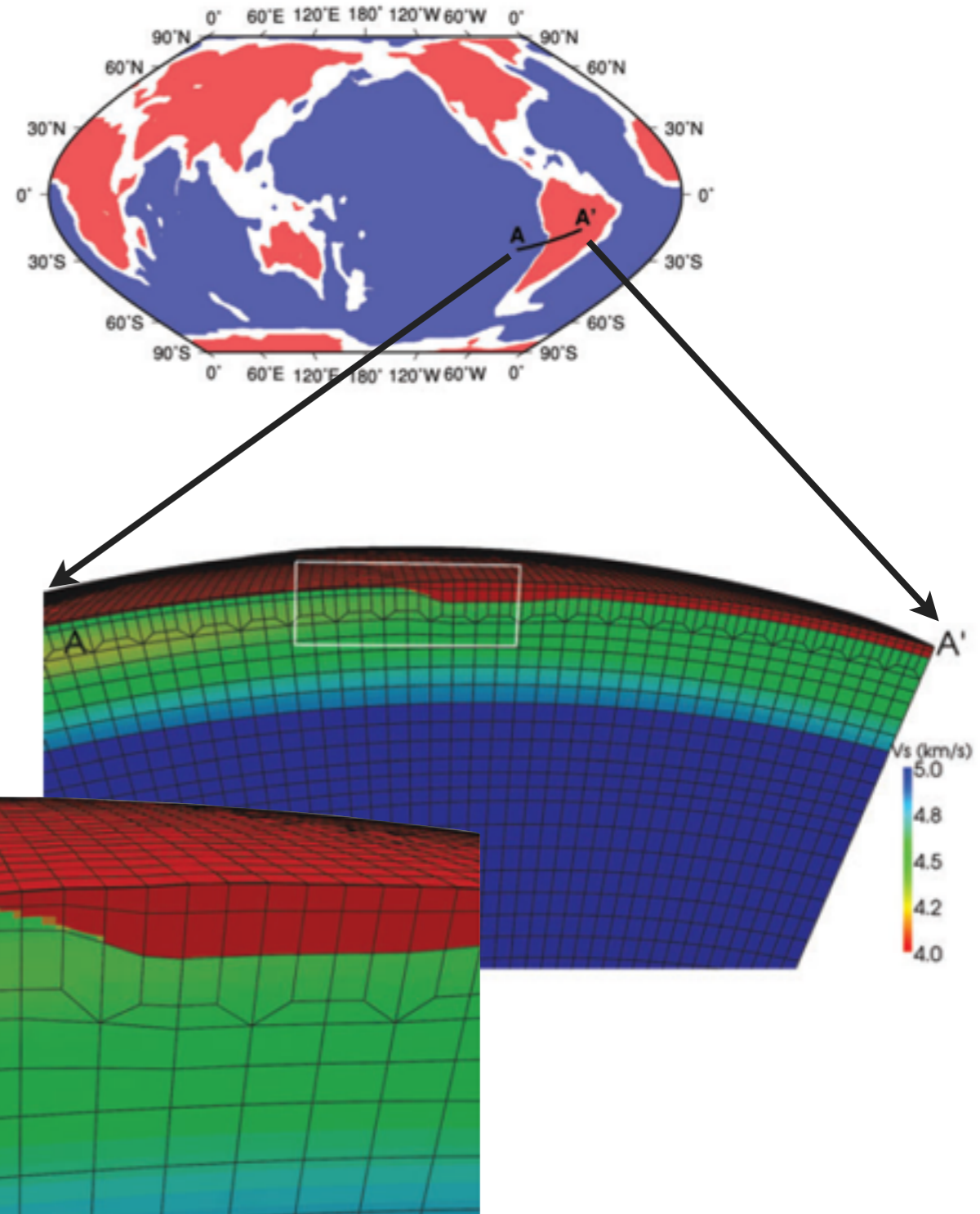
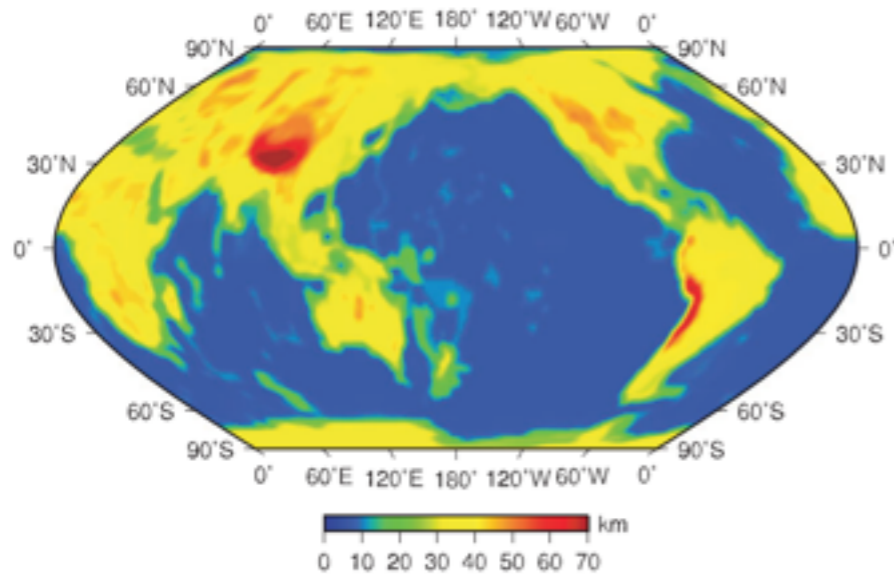
- Numerical simulations
- Source inversions
- Adjoint tomography
 - 1st iteration results!

Numerical simulations

- SPECFEM3D_GLOBE (Komatitsch & Tromp 2002)
- 3D Reference model: S362ANI (Kustowski et al. 2008)
+ Crust2.0 (Bassin et al. 2000)
- Topography / bathymetry / attenuation / ellipticity /
rotation / gravity
- Length of seismograms = 100 m
- $T_{\min} = \sim 27$ s

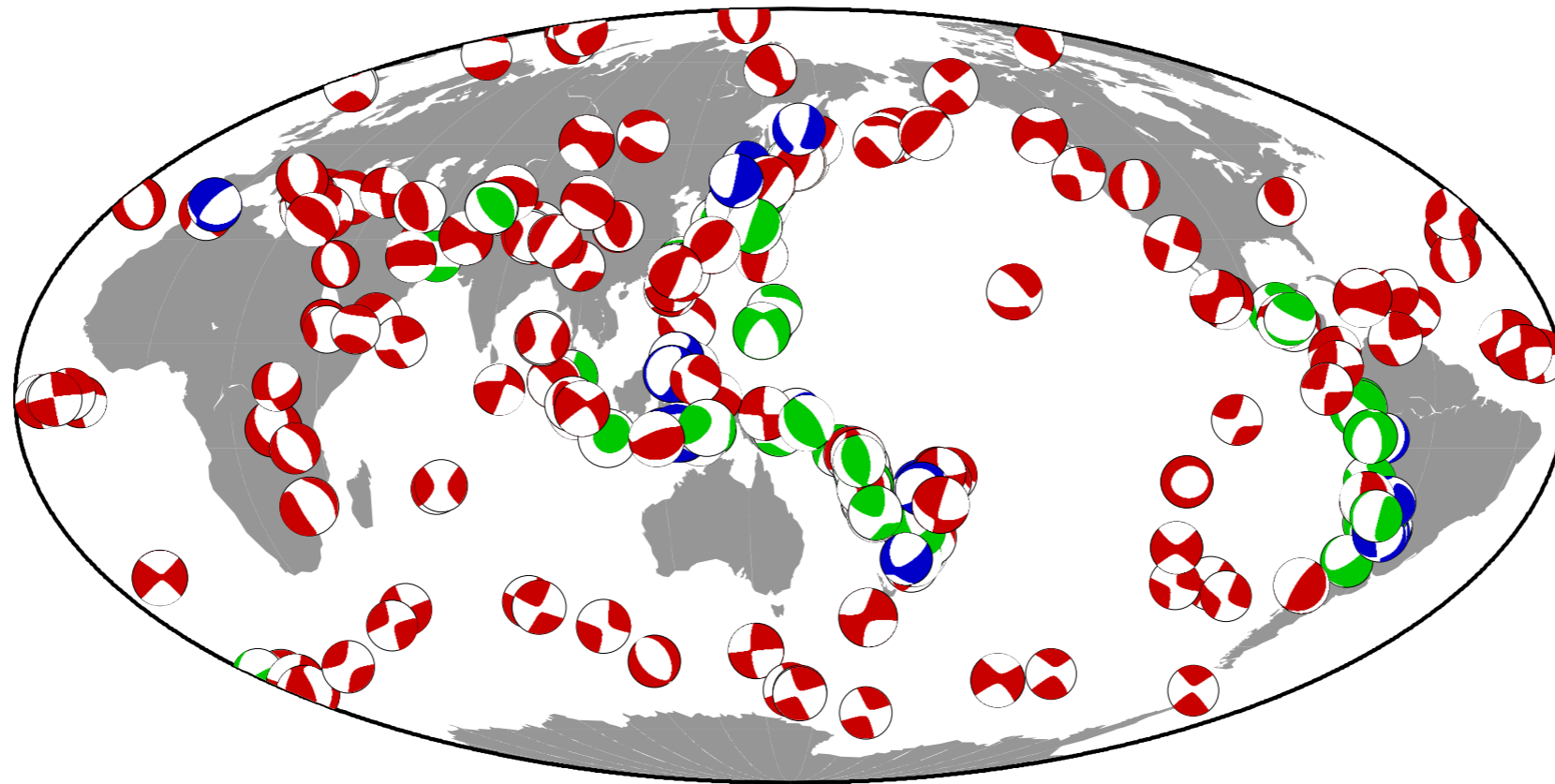
Implementation of crust in simulations

Smoothed Moho depths
from Crust2.0



Moho is honored
if crustal thickness is
 ≤ 15 km and ≥ 35 km.

255 global CMT earthquakes



shallow



intermediate



deep

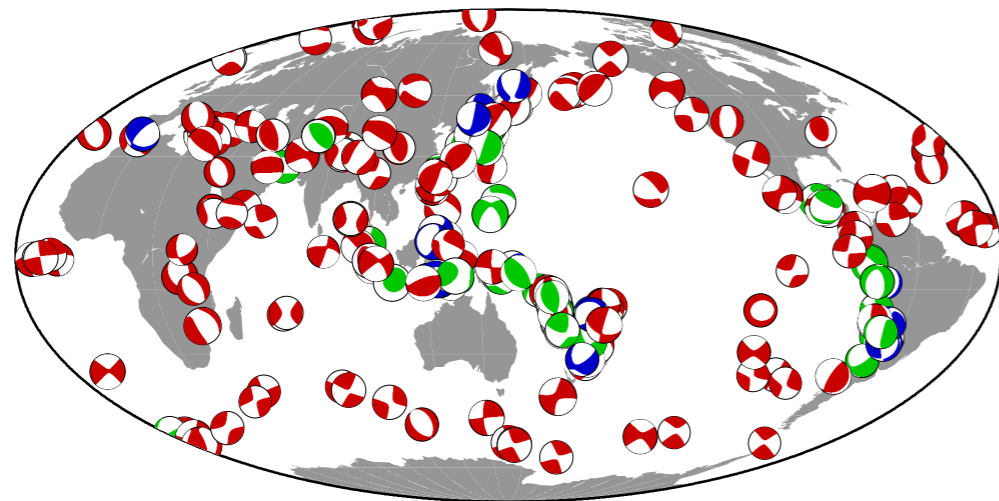
$5.8 \leq M_w \leq 7.0$

shallow: $d \leq 50$ km

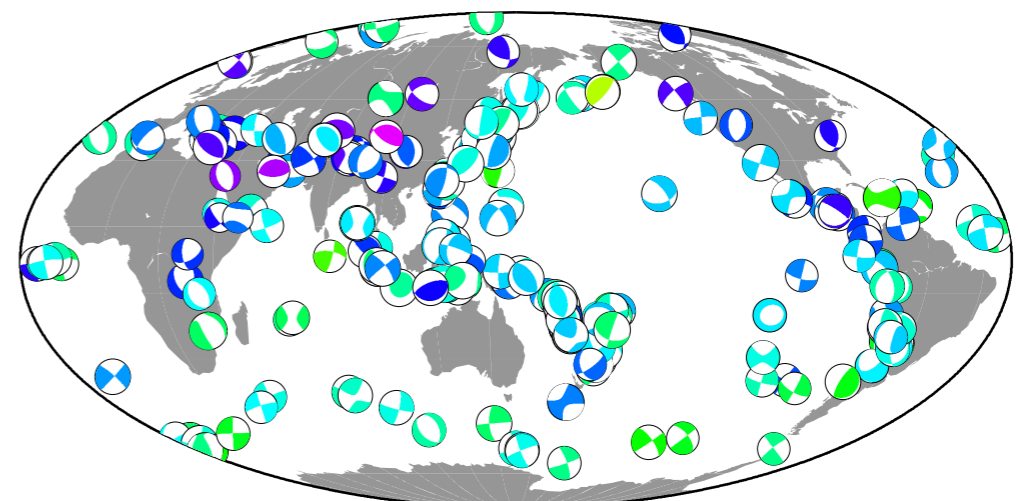
intermediate: 50 km $< d \leq 300$ km

deep: $d > 300$ km

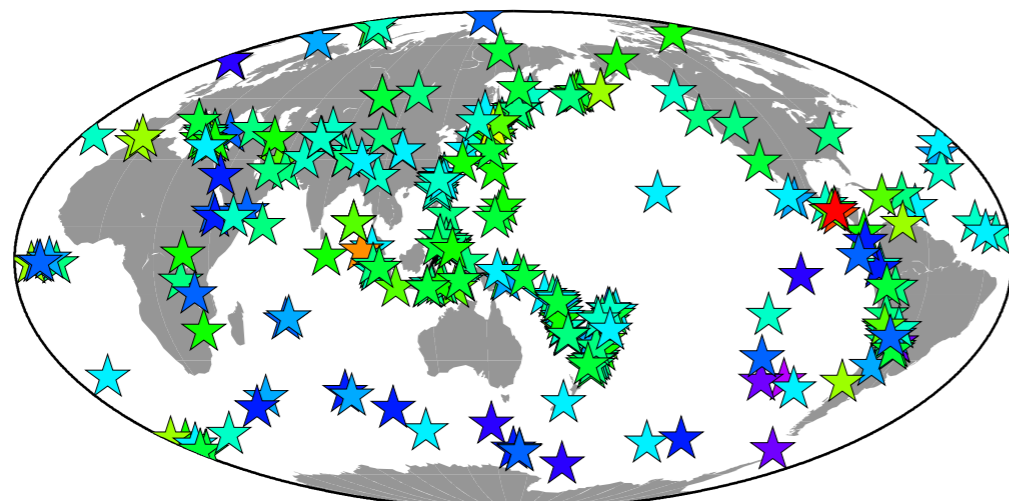
Source inversions - summary



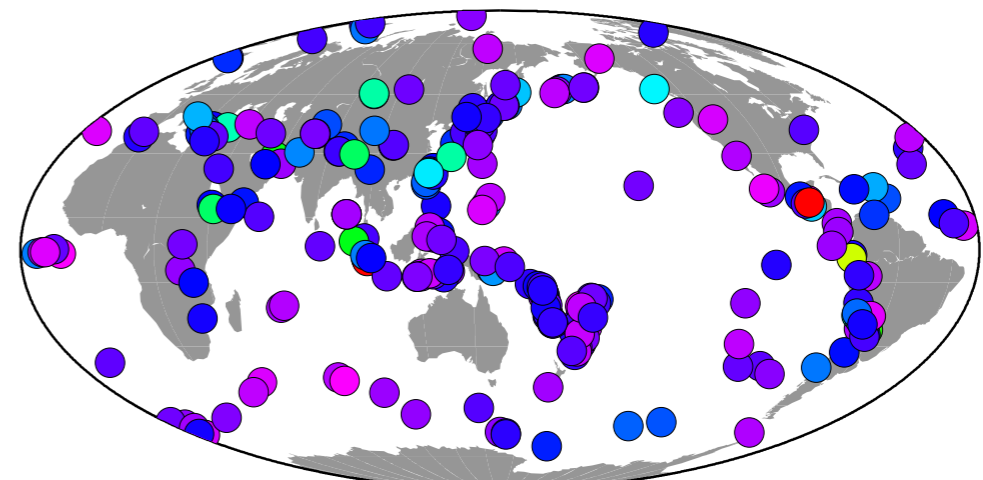
shallow intermediate deep



$\delta \ln M_0$



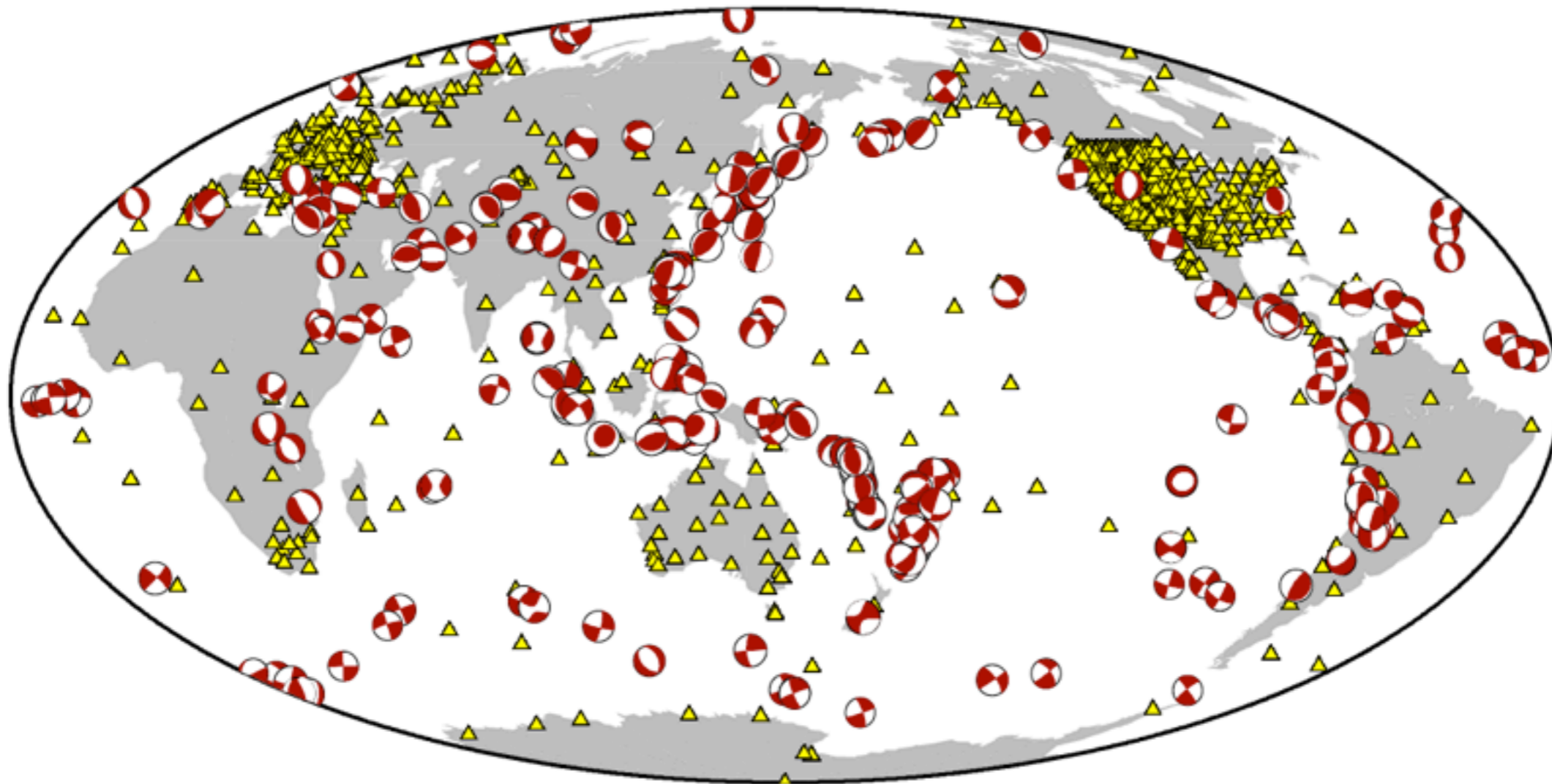
$\delta \text{depth-km}$



$\delta \text{loc-km}$

Adjoint tomography

Earthquake-station distribution



253 global CMT events ($5.8 \leq M_w \leq 7.0$)

Data from IRIS & ORFEUS

Inversion strategies

Multitaper travelttime measurements

$$\chi_c = \frac{1}{N_c} \sum_{s=1}^S \sum_{i=1}^{N_c^s} \int w_i(\omega) \left[\frac{\Delta\tau_i(\omega)}{\sigma_i(\omega)} \right]^2 d\omega$$

χ_c : misfit per category

N_c : number of picks
per category

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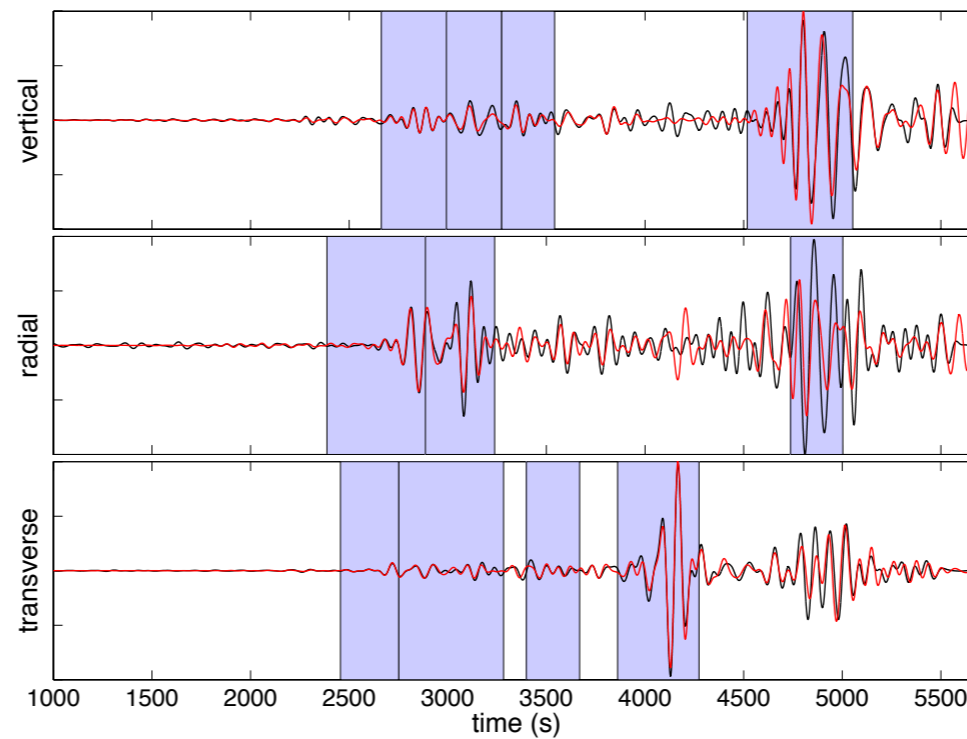
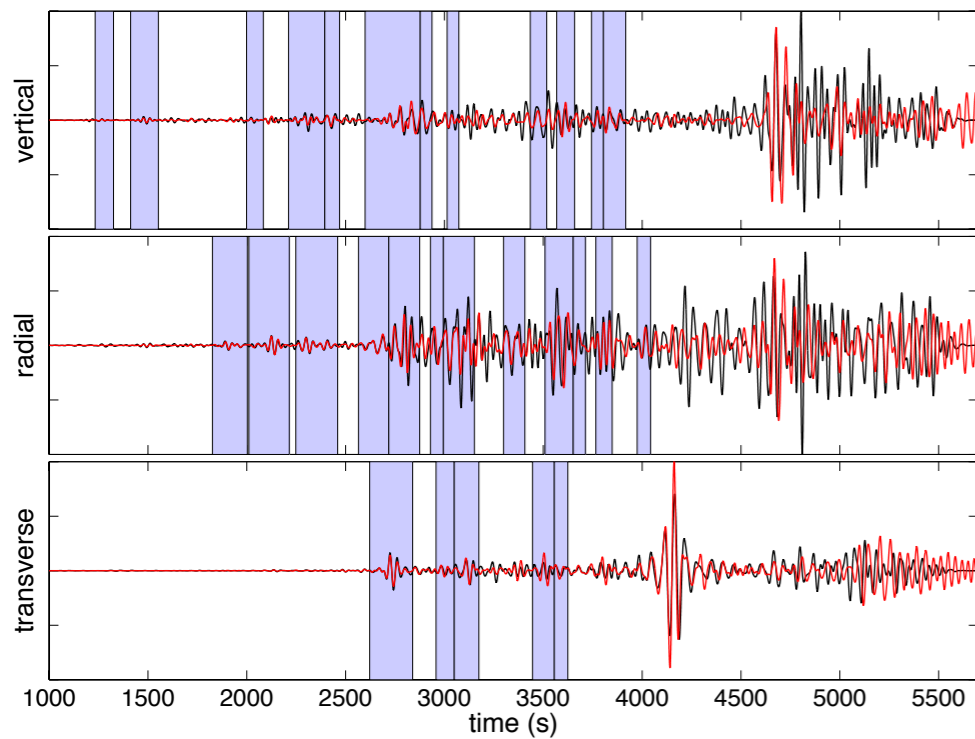
$$\chi^{total} = \chi^{27-60s} + \chi^{60-120s}$$

- 1) P-SV on vertical
- 2) P-SV on radial
- 3) SH on transverse

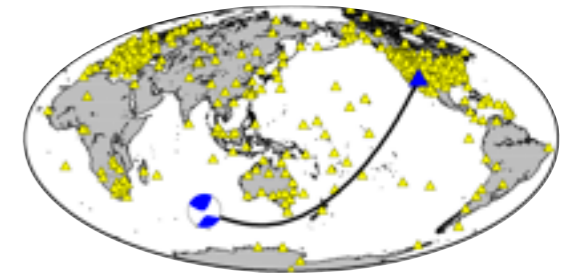
- 4) P-SV-Rayleigh on vertical
- 5) P-SV-Rayleigh on radial
- 6) SH-Love on transverse

Data selection

2008, May 31, Mid-Indian Ridge event
 $M_w=6.4$, depth=6.5 km



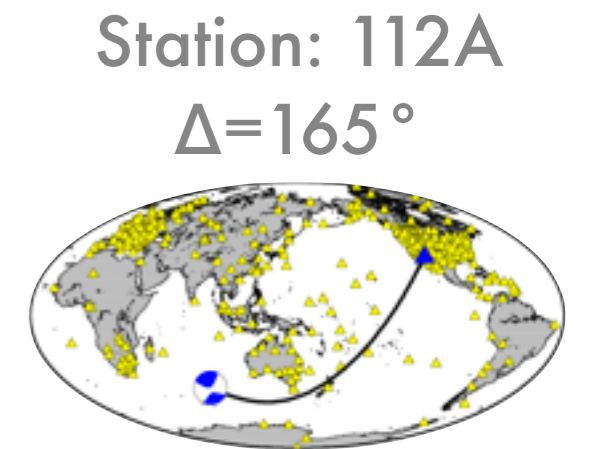
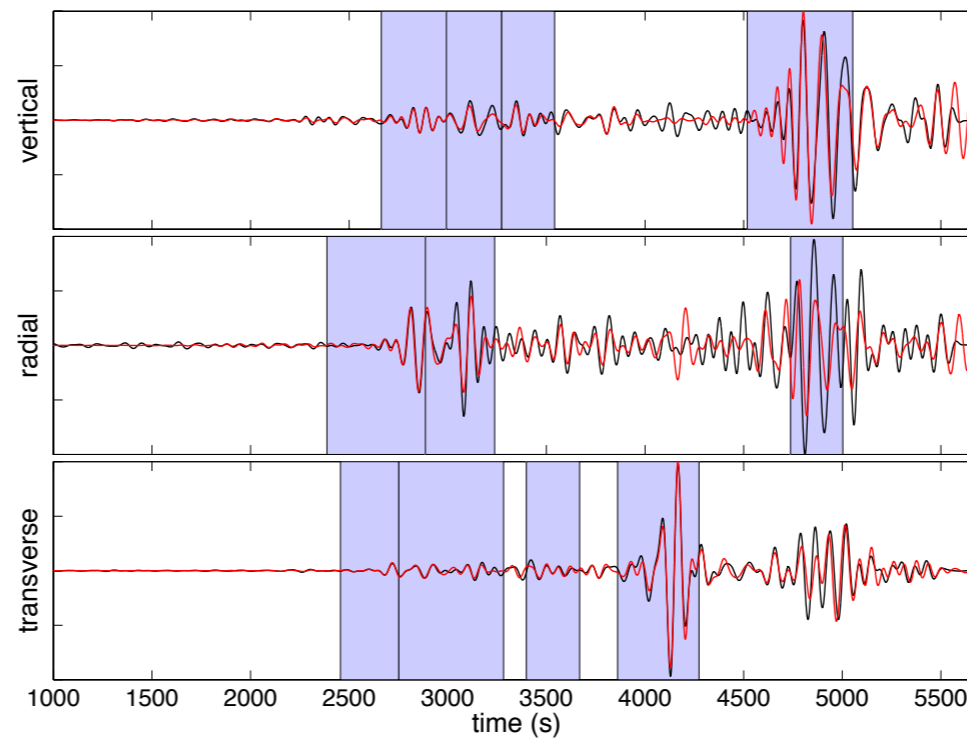
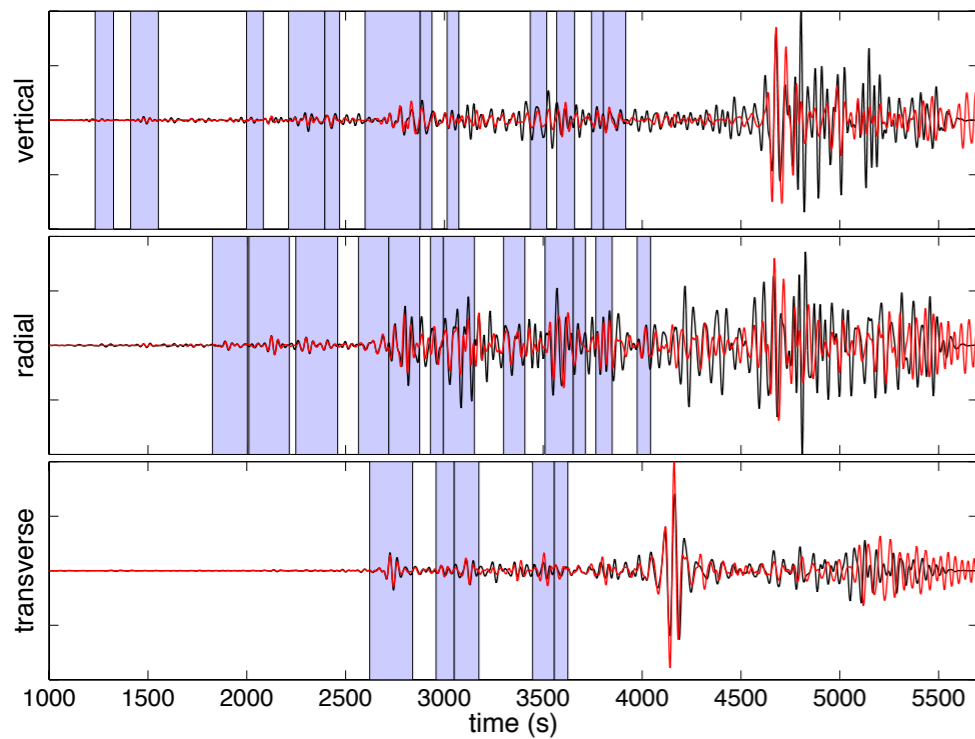
Station: 112A
 $\Delta=165^\circ$



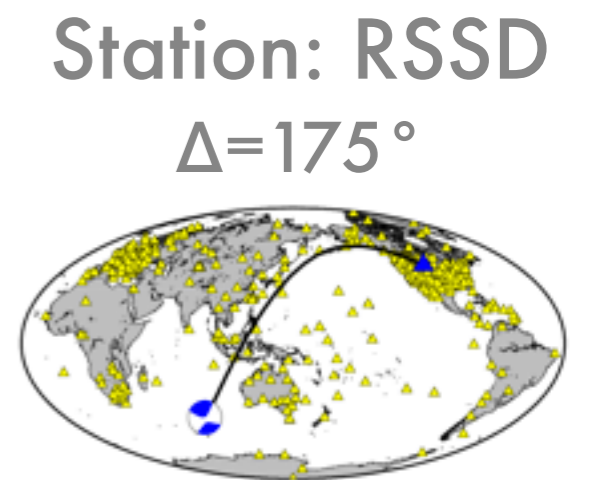
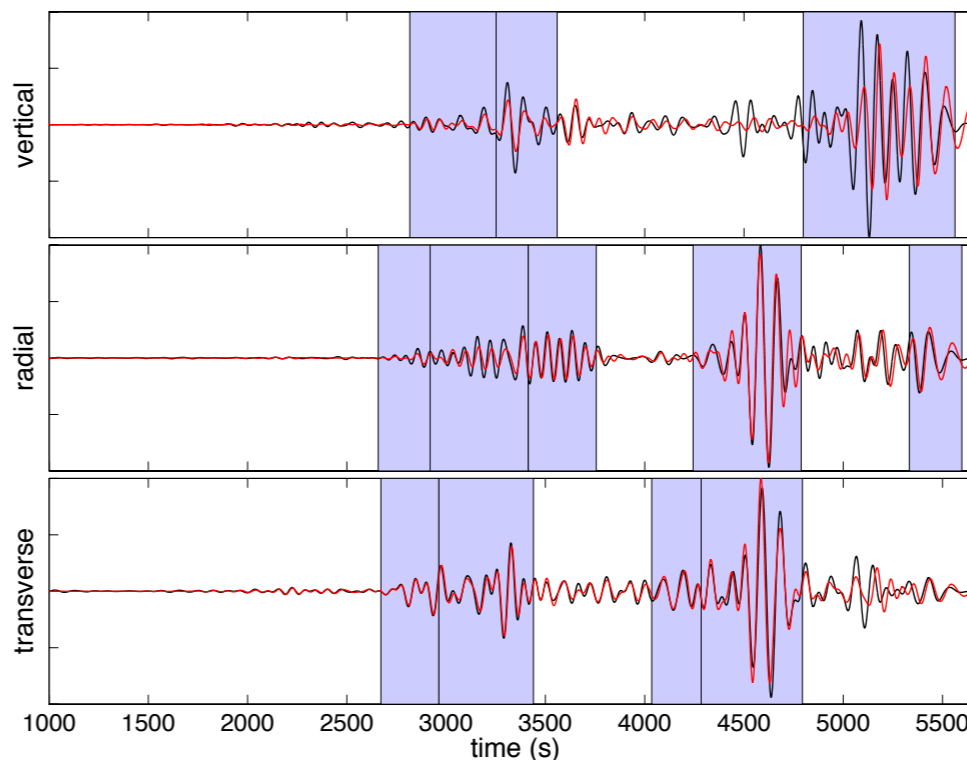
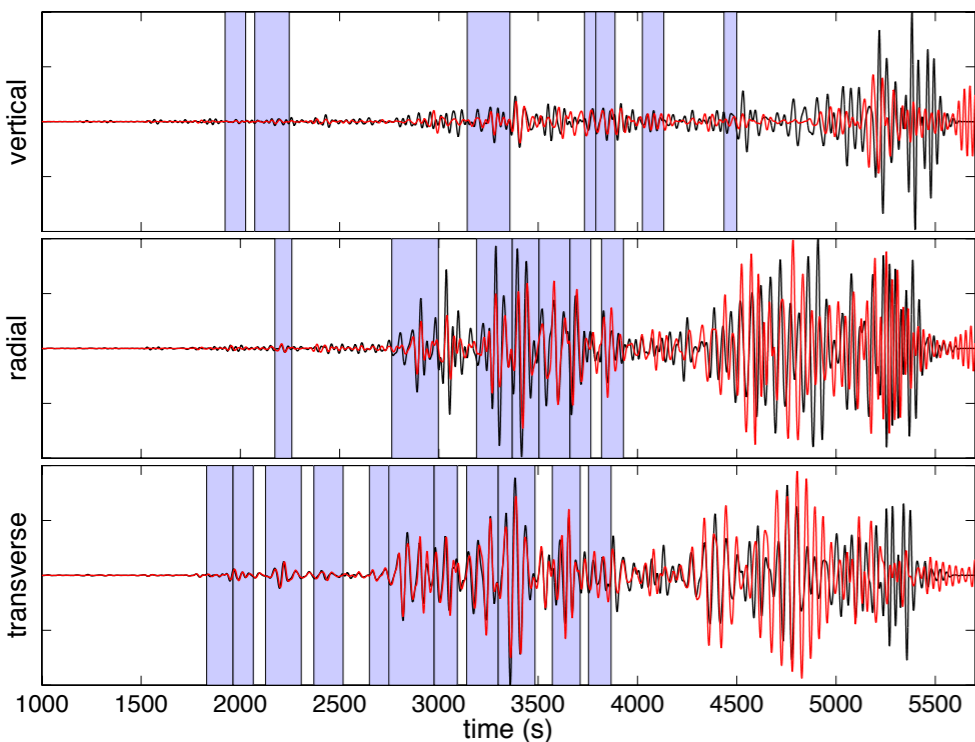
window selection:
FLEXWIN
(Maggi et al. 2009)

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window selection:
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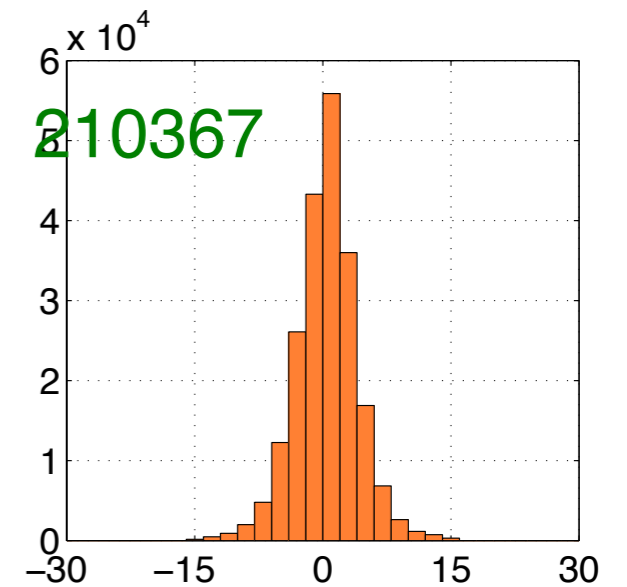
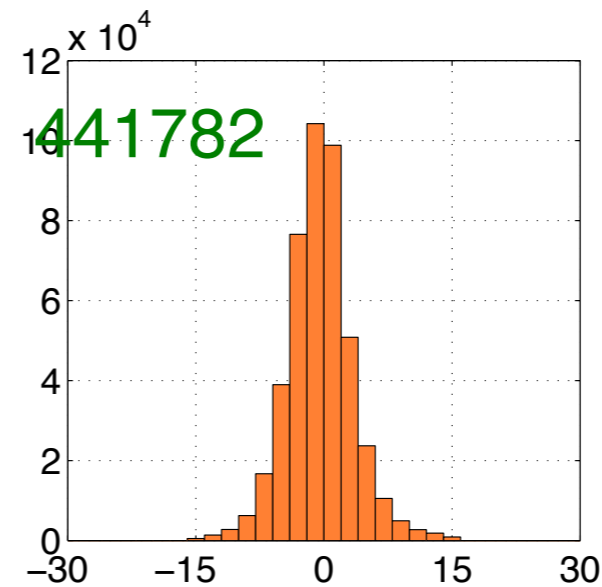
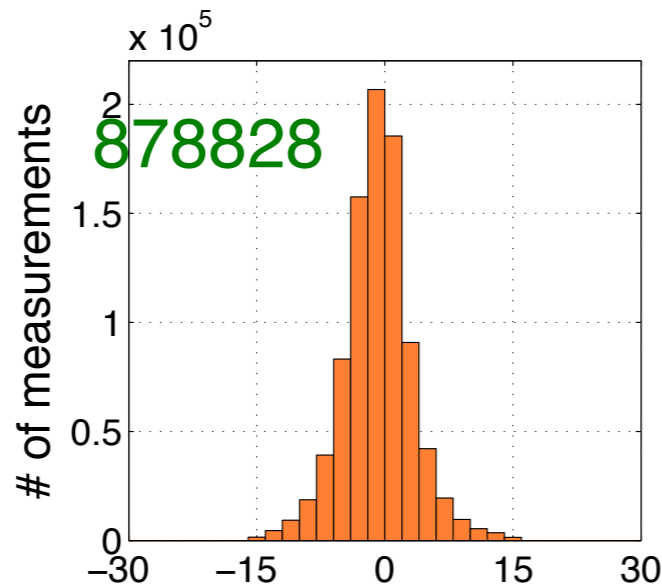


Cross-correlation time-shifts

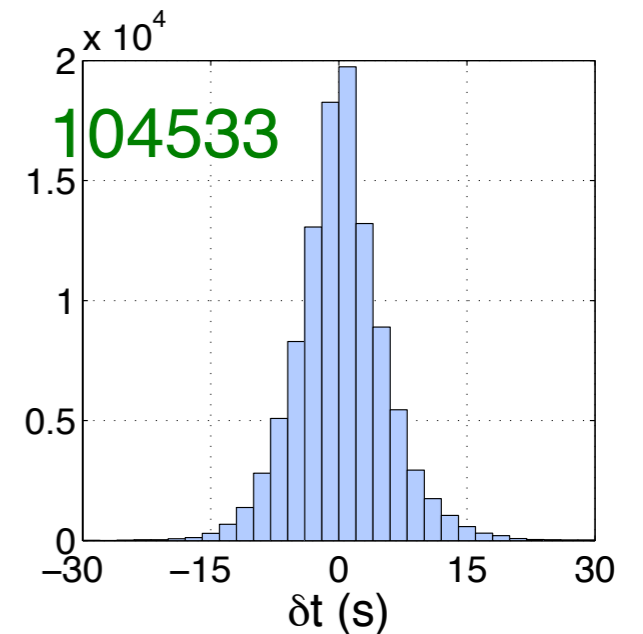
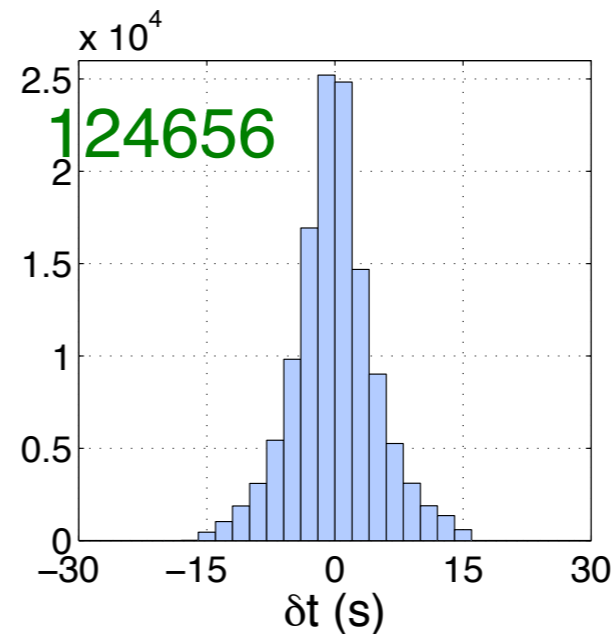
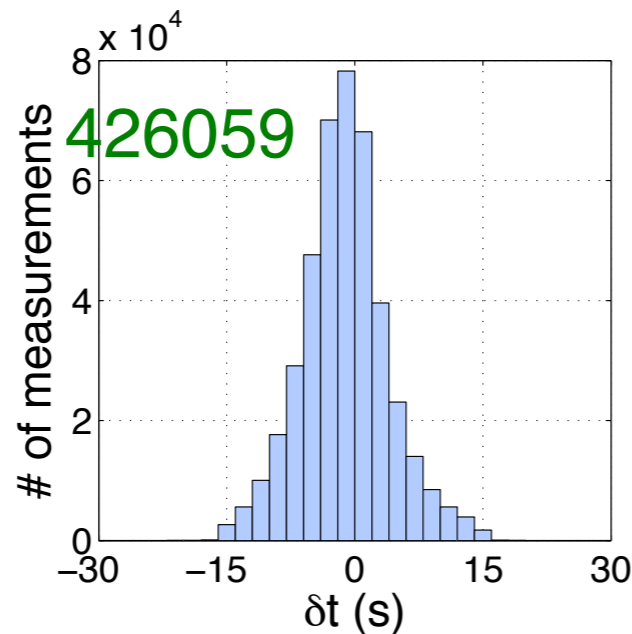
vertical

radial

transverse



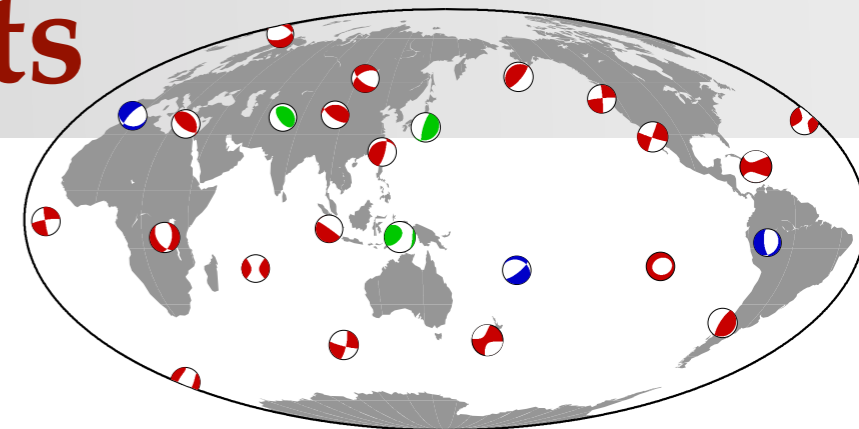
27 - 60 s



60 - 120 s

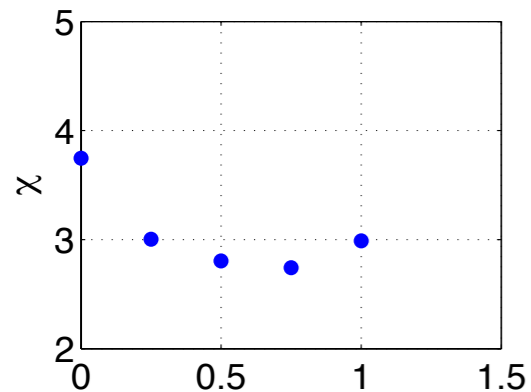
~2.2 million measurements

Line search with 24 test events

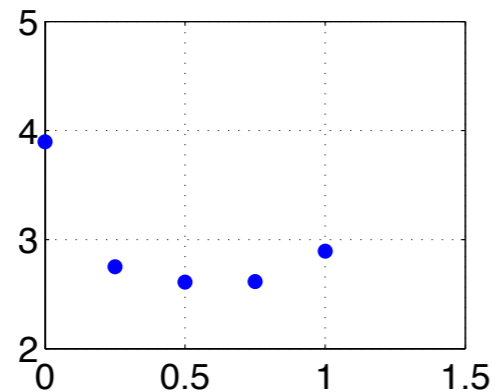


27 - 60 s

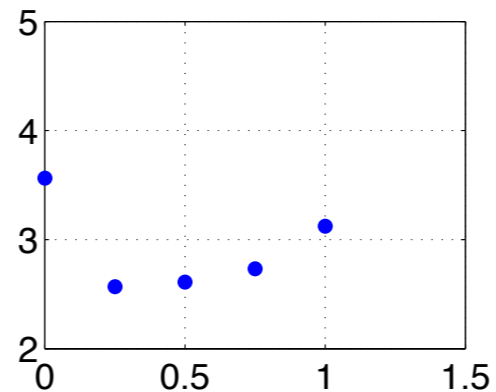
Vertical



Radial



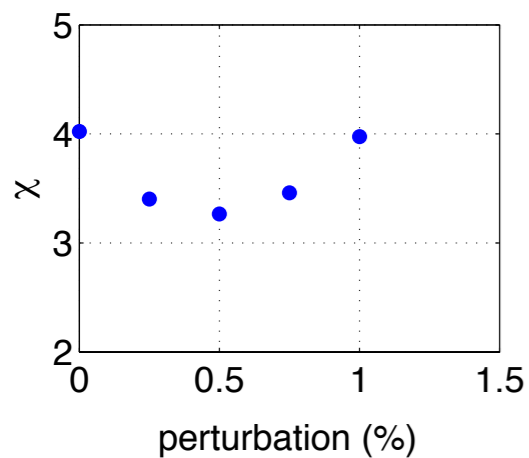
Transverse



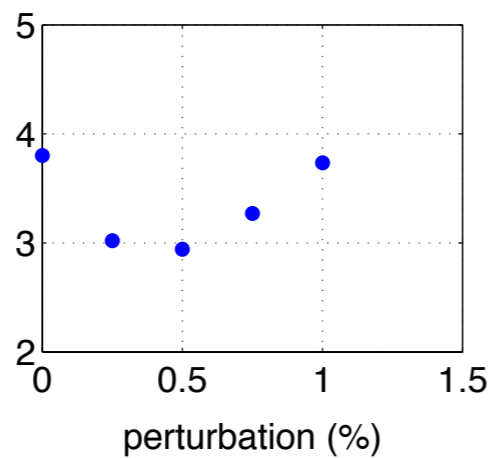
shallow intermediate deep

60 - 120 s

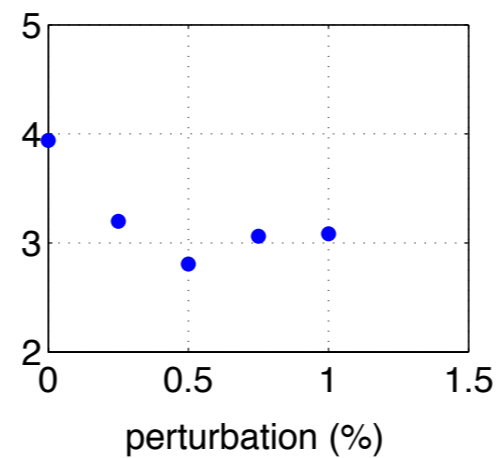
Vertical



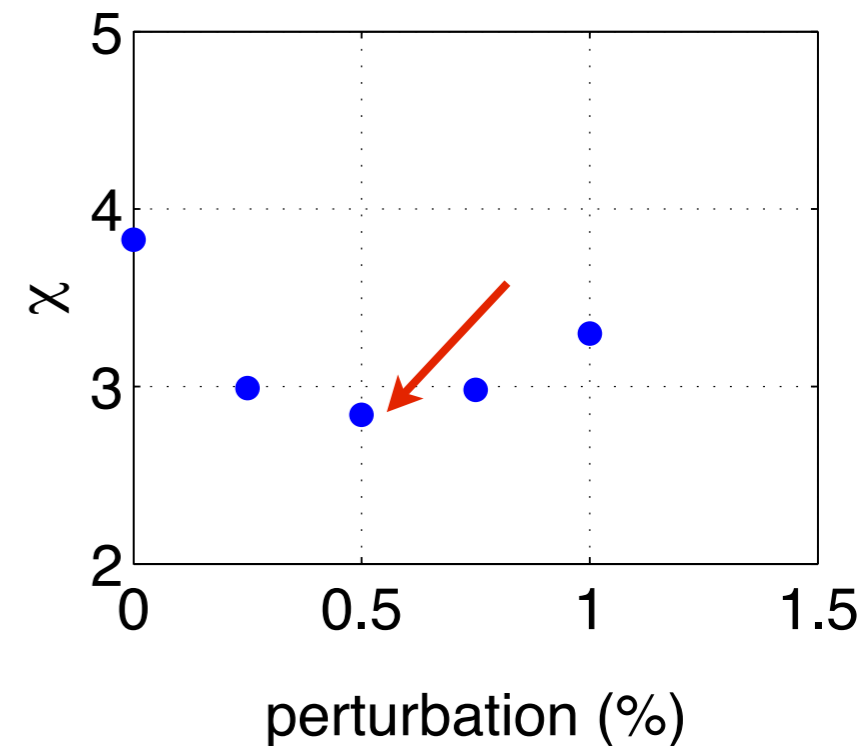
Radial



Transverse



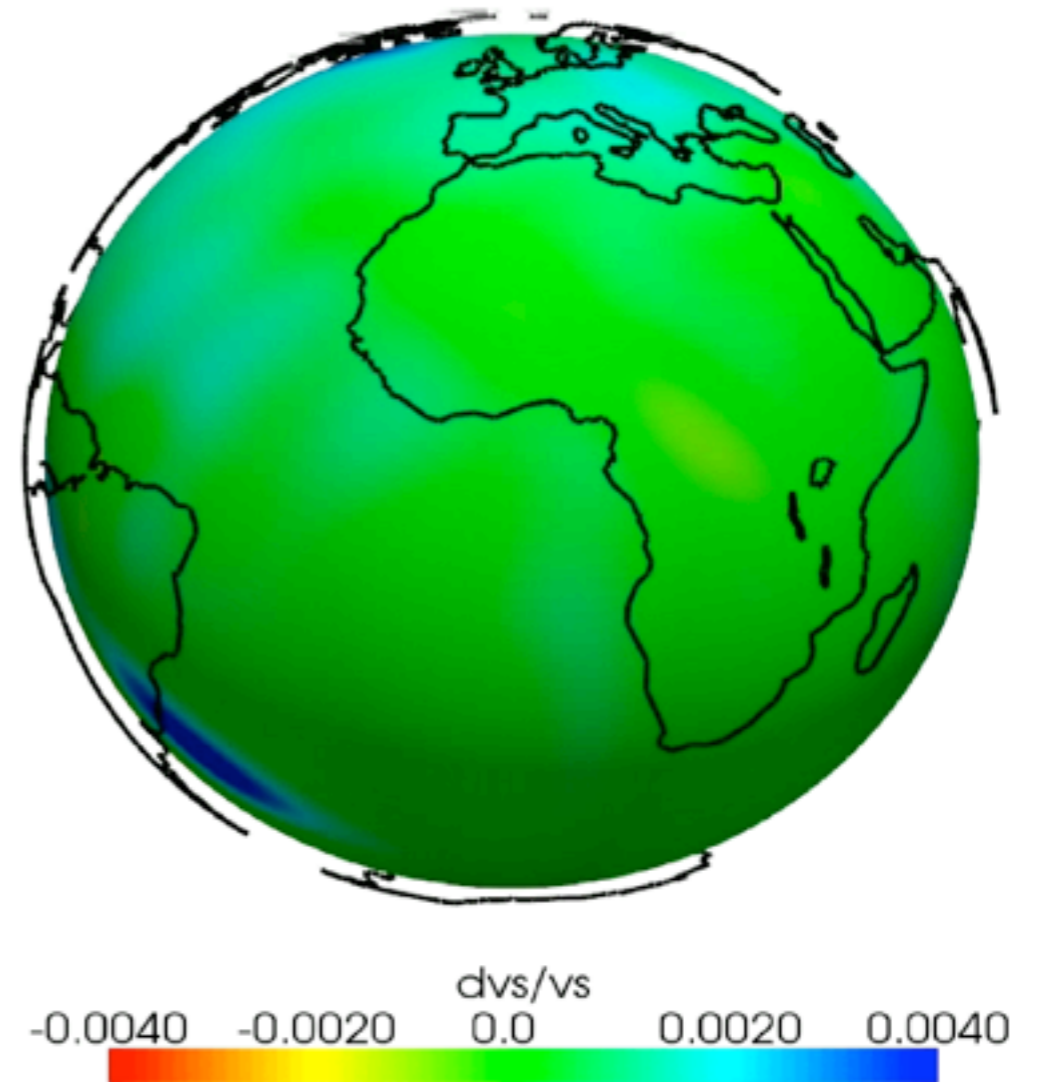
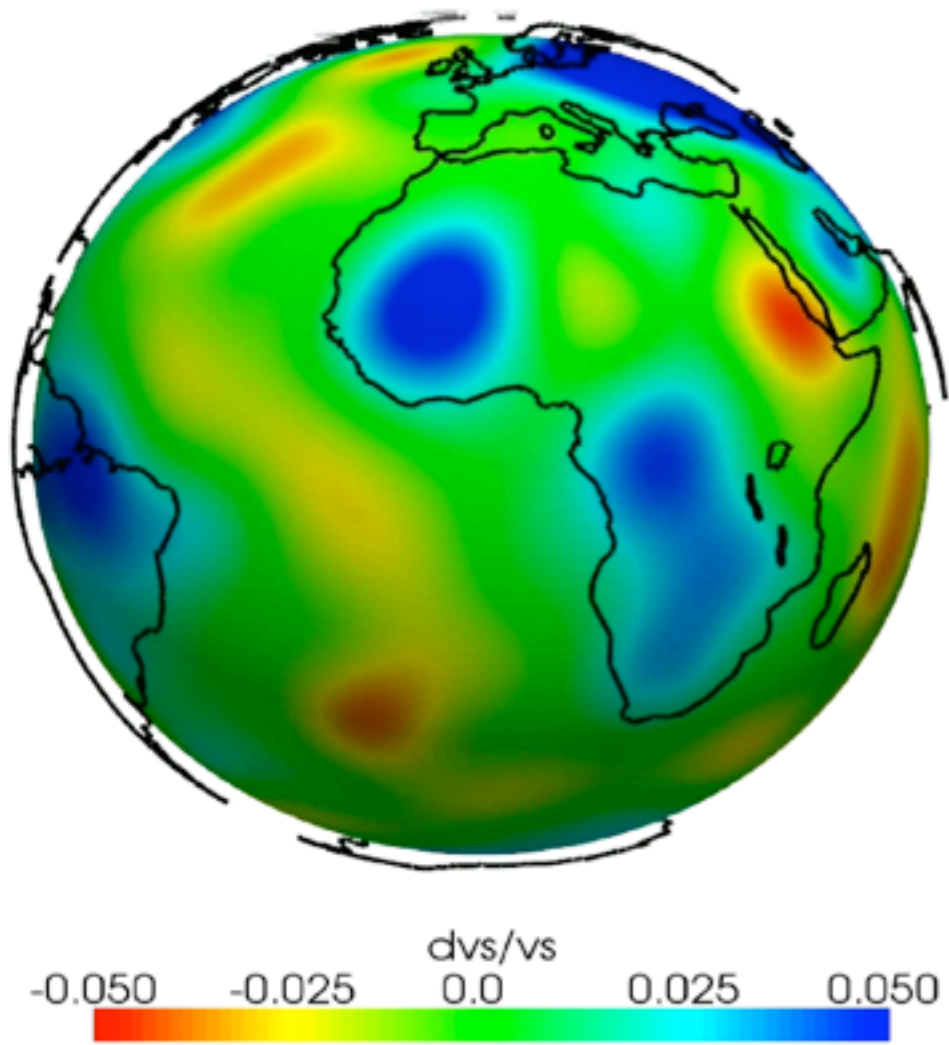
Total misfit



M00 - 1DREF

M01 - M00

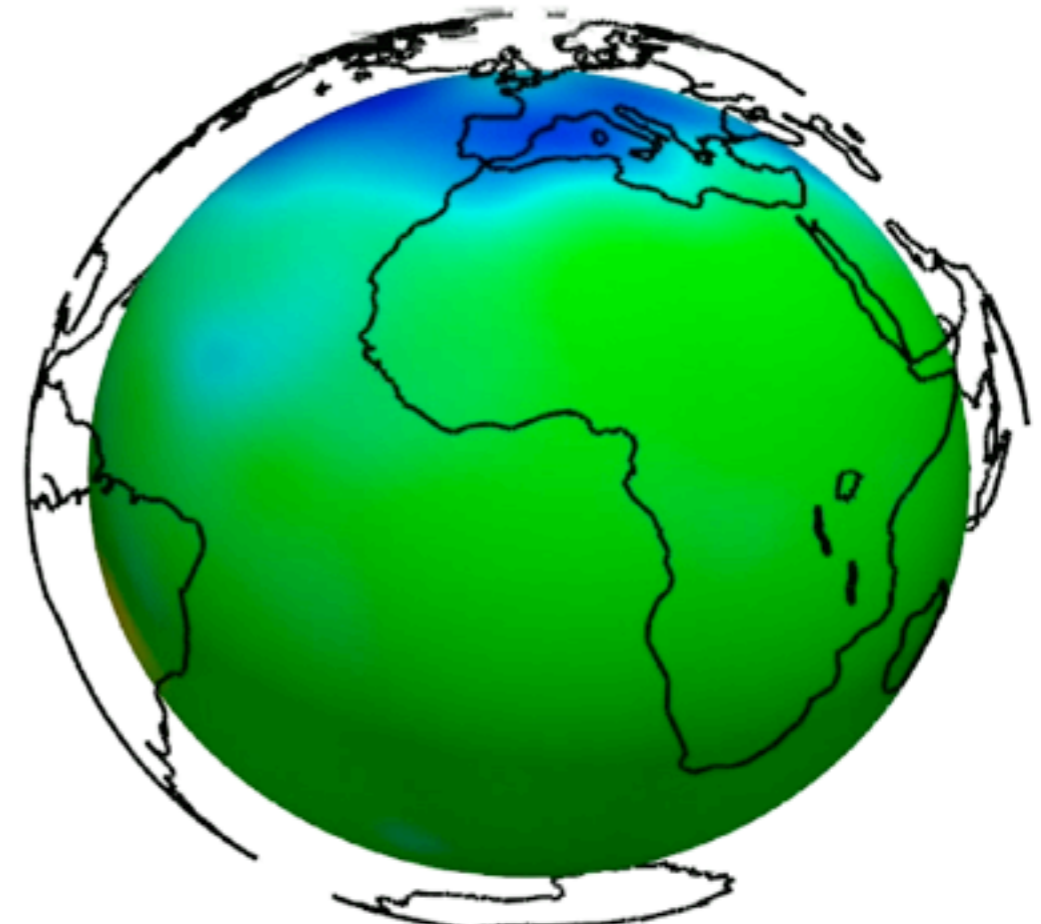
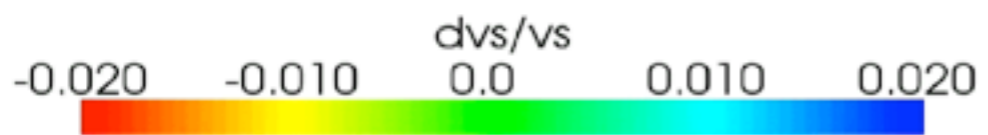
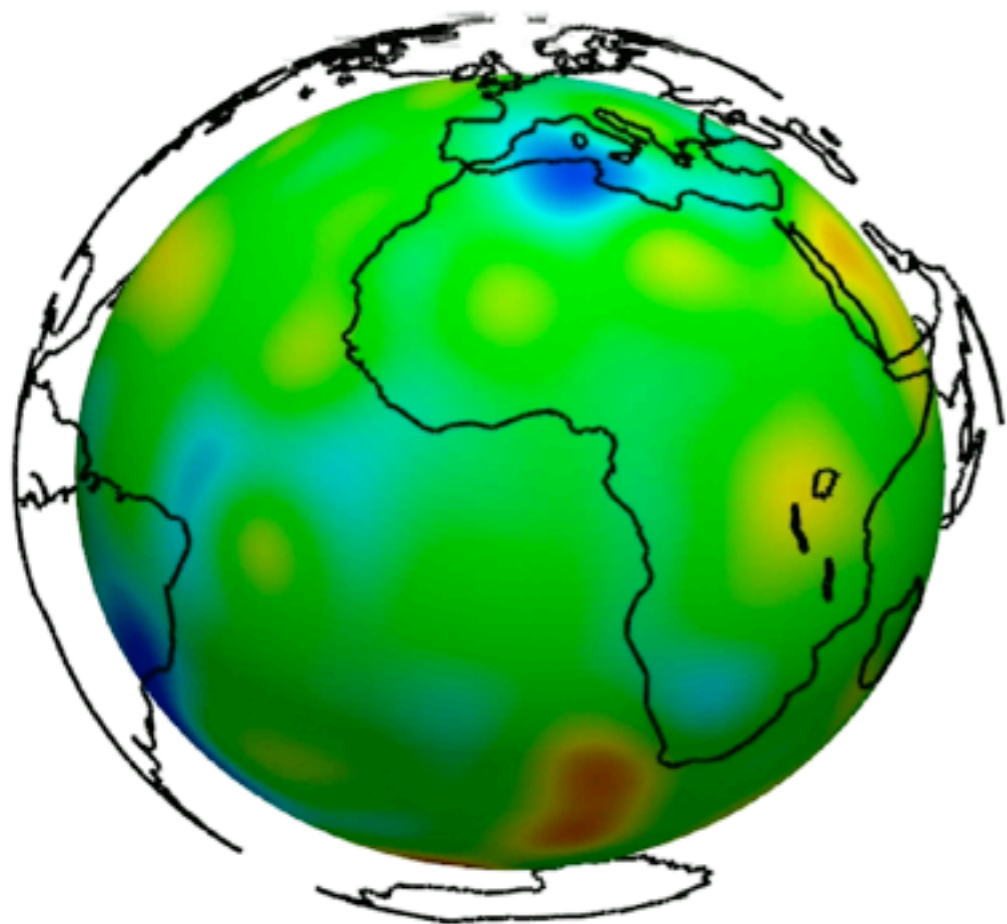
150 km



M00 - 1DREF

M01 - M00

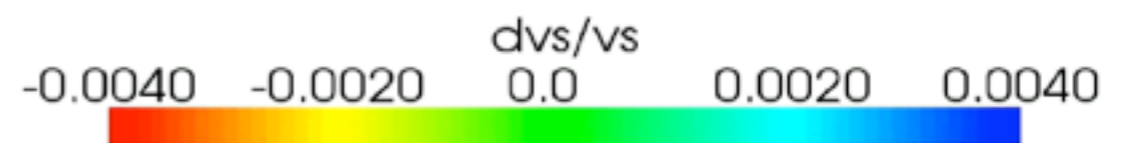
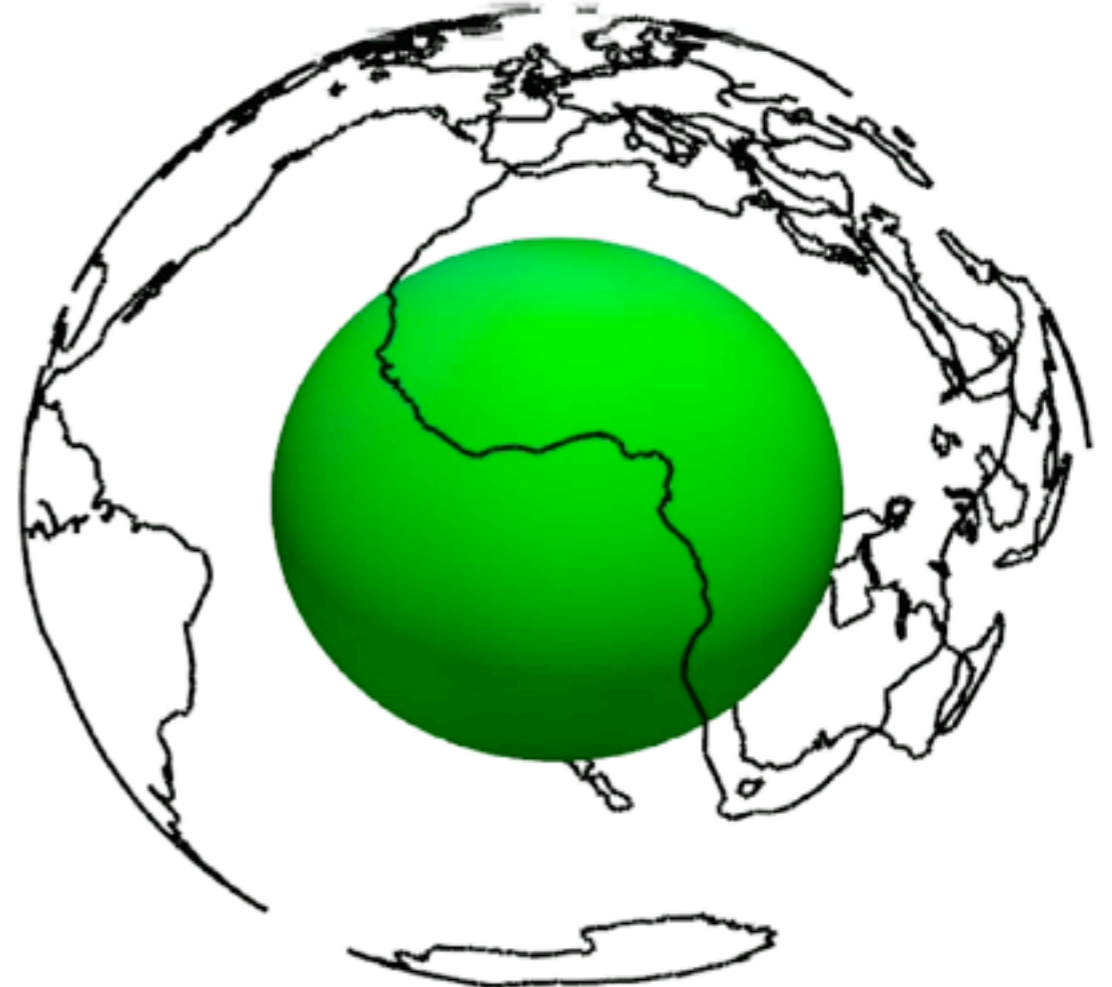
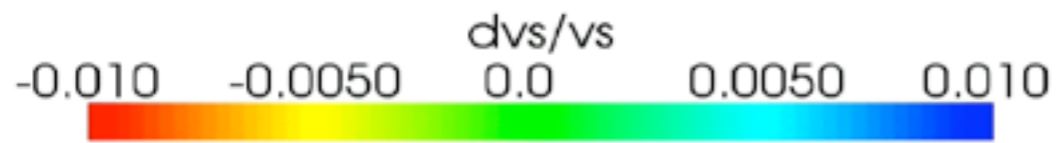
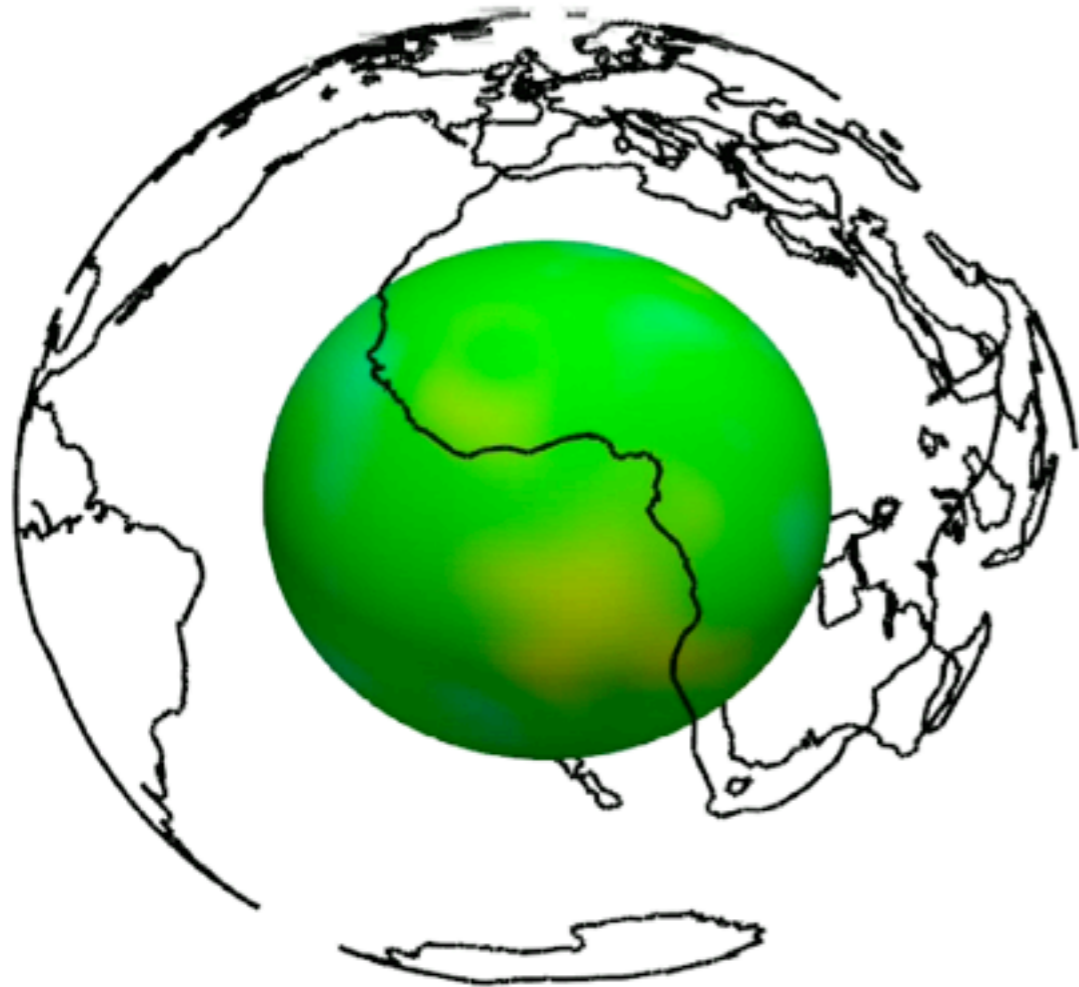
660 km



M00 - 1DREF

M01 - M00

CMB

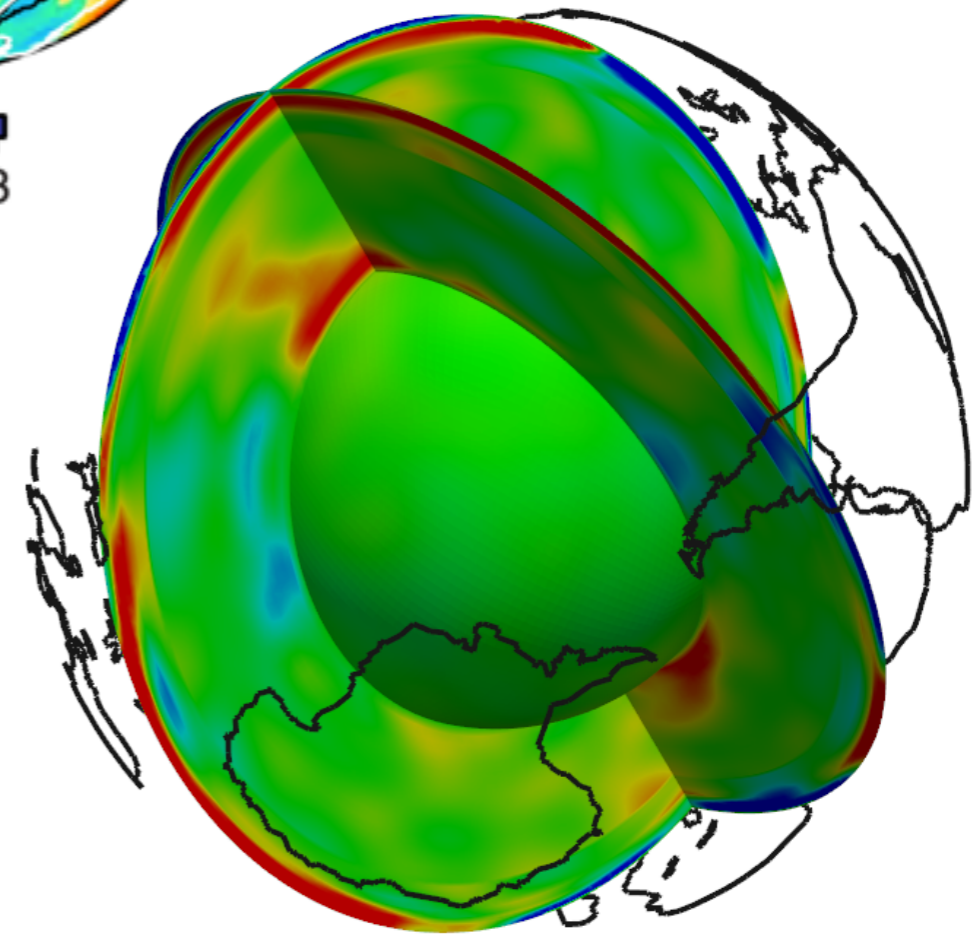
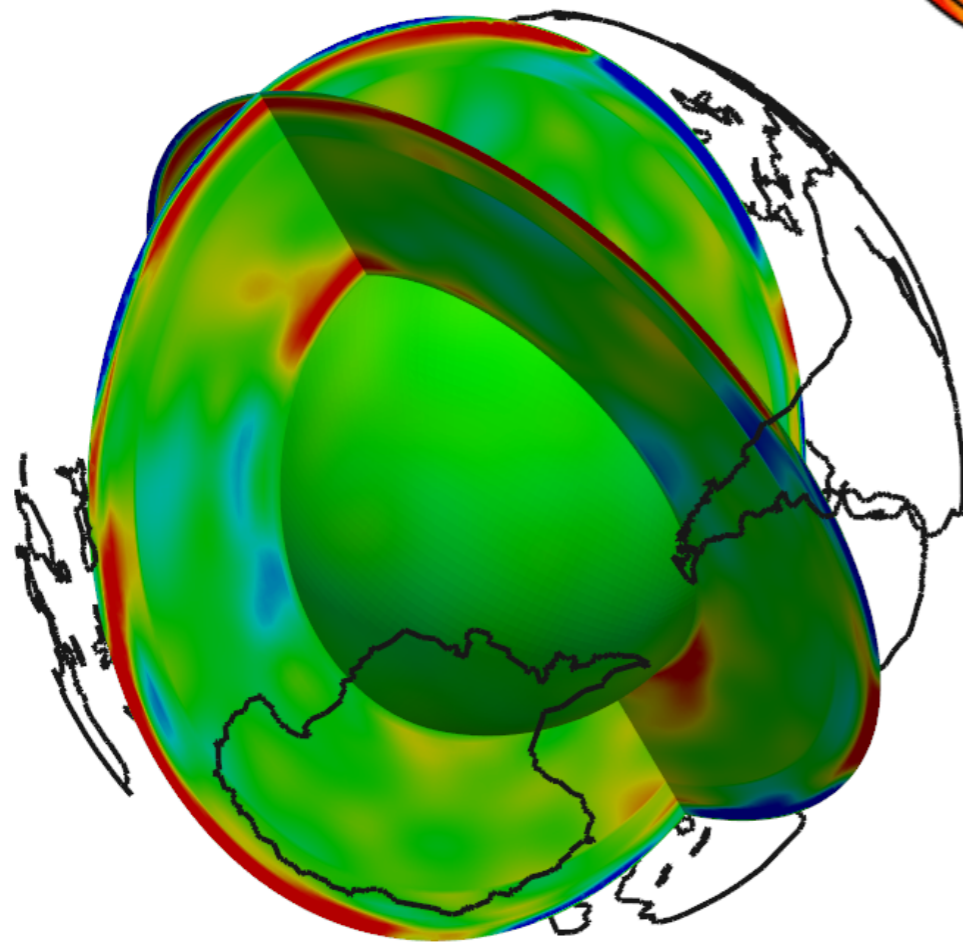
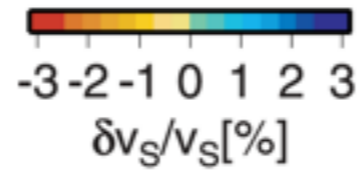
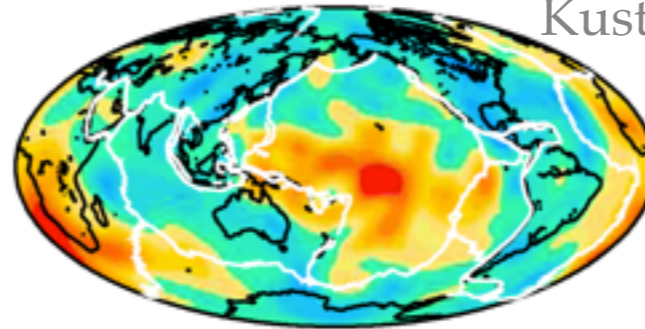


M00 - 1DREF

M01 - 1DREF

2800 km

Kustowski et al. 2008



Challenges

- Computational requirements

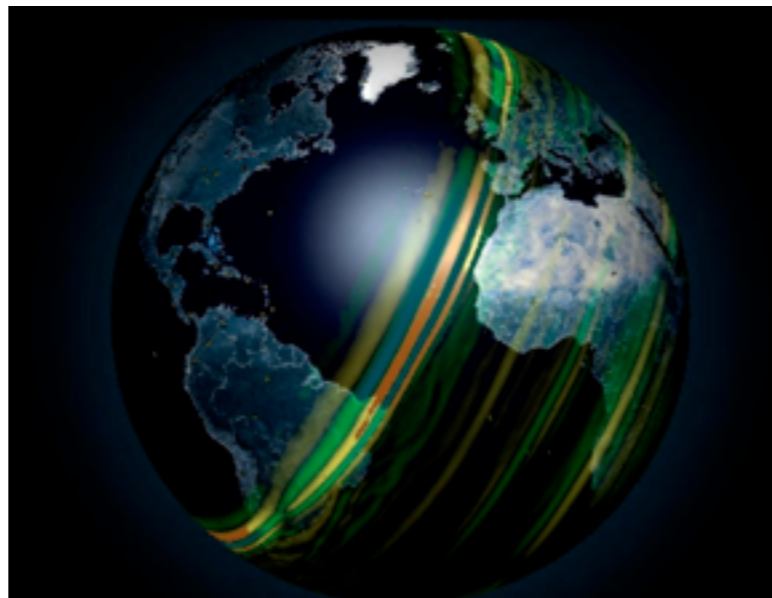
CPU hours	1 event	1 iteration (255 events)	20 iterations
forward + adjoint	3000	765,000	15,300,000

- Data processing - manual quality check
- Uneven distribution of source and receivers - Balance in gradient

Remedies

- More computational resources!
- Speeding up the forward / adjoint simulations: GPU computing
- Increasing data: using more earthquakes!

- First slide global wave propagation picture: April 12, 2012 Gulf of California Earthquake ($M_w = 7$, depth = 14 km) (global.shakemovie.princeton.edu).



- Master slide seismogram is from SPICE presentation template (www.spice-rtn.org).