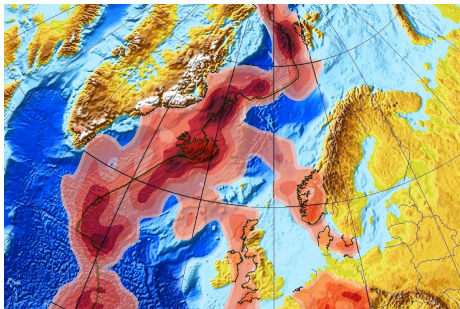


A new full-waveform model of the North Atlantic region

Florian Rickers, Andreas Fichtner, Jeannot Trampert

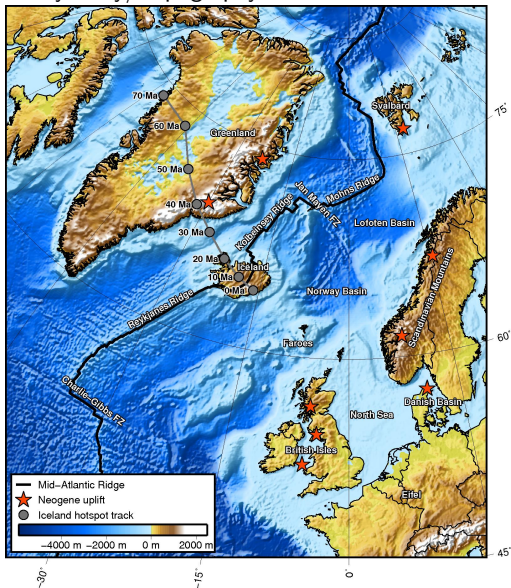


- 4th QUEST Workshop in Benodet (France), May 2013 -

- 1 The North Atlantic region
- 2 Adjoint tomography and instantaneous phase misfit
- 3 Dataset and Inversion
- 4 Features of the model
- 5 Conclusions

The North Atlantic region

Bathymetry/Topography

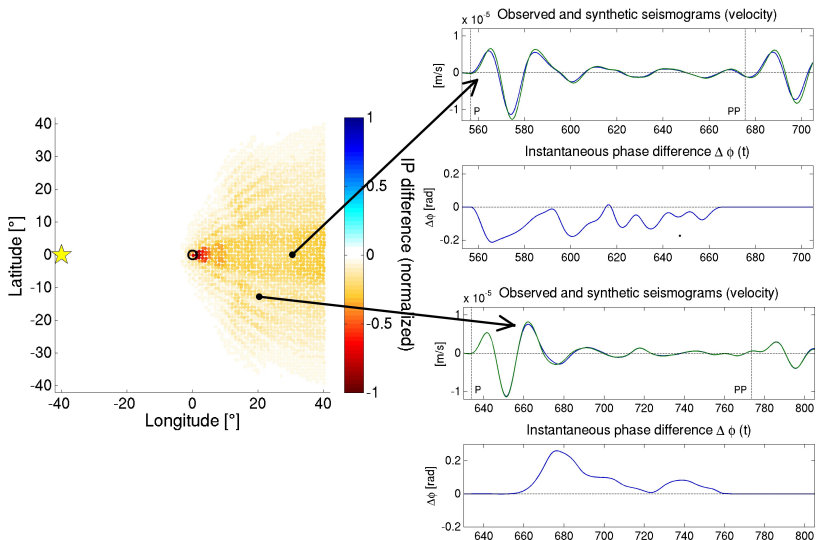


- ▶ Iceland/Jan Mayen: Deep mantle plume or shallower source?
- ▶ Are Iceland and Jan Mayen separate hotspots / plumes?
- ▶ What is the lateral extent of the 'plume head' beneath the lithosphere?
- ▶ Is the post-breakup uplift (< 30 Ma) at the continental margins related to the plume?

1. Computation of the **forward wavefield** using the spectral element method (SEM)
→ Accurate seismograms in 3-D reference model
2. Evaluation of **misfit** between data and synthetics, computation of adjoint sources
3. Computation of the **adjoint wavefield** by time-reverse propagation of adjoint sources (again using SEM)
4. **Gradients** (sensitivity kernels) are constructed from interaction of forward and adjoint wavefields
5. Iterative model update using a **conjugate gradient scheme**

⇒ Fully 3-D non-linear inversion based on arbitrary misfit
Potential to exploit the full seismogram

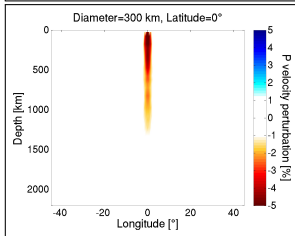
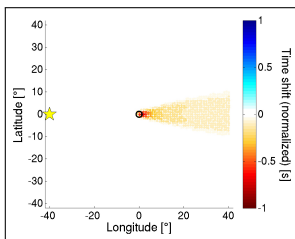
Instantaneous phase misfit



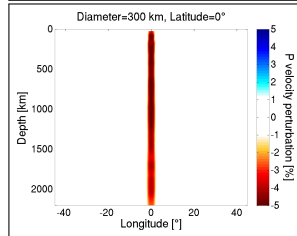
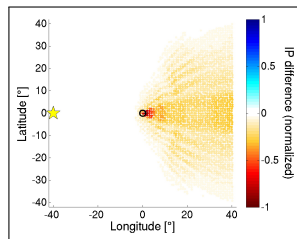
Time-continuous, amplitude-independent phase difference

⇒ Diffracted waves in the coda can be measured and included in the inversion

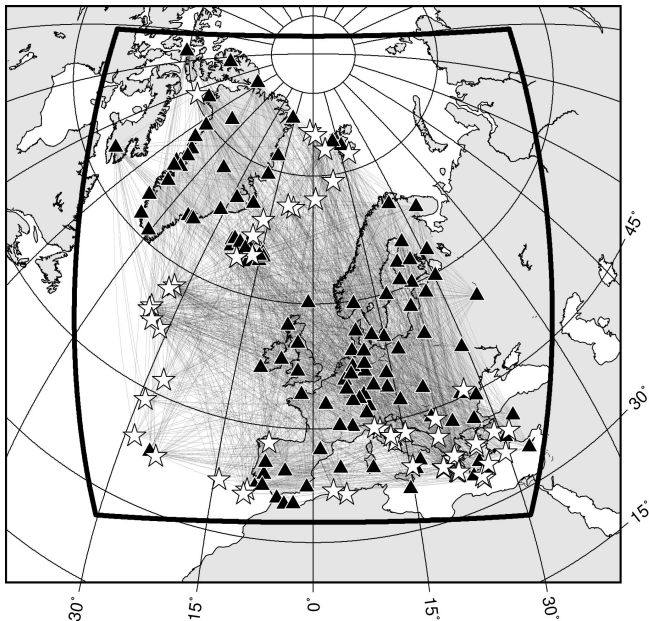
Travel times (CC)



Instantaneous phase

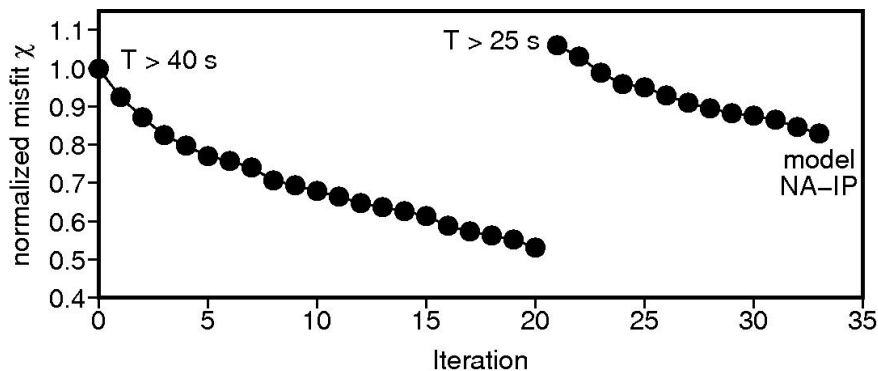


⇒ Plumes in the deep mantle can be imaged if diffracted waves are considered



- ▶ Depth: 1500 km
- ▶ 60 events
- ▶ ca. 4500 seismograms
- ▶ Full-waveform starting model (Fichtner et al., 2011)
- ▶ Inversion for v_{sh} and v_{sv} (but only v_{sh} presented)

Inversion: Reduction of the total misfit



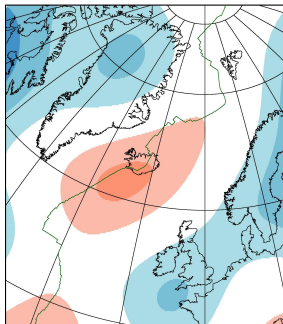
- ▶ \approx 2.5 million CPU hours

Validation: Comparison with S20RTS and EU-TF (at 80 km depth)

S20RTS

(Ritsema et al., 1999)

Depth 80 km – $\delta \ln \beta_{sh}$ w.r.t. 4.406 km/s

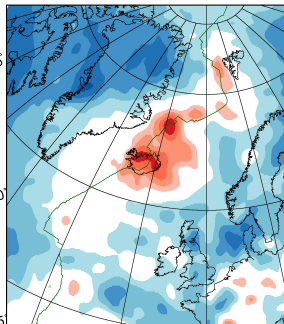


⇒

EU-TF

(Fichtner et al., 2011)

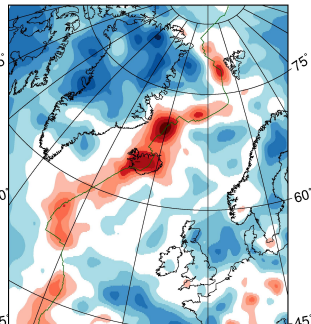
Depth 80 km – $\delta \ln \beta_{sh}$ w.r.t. 4.406 km/s



⇒

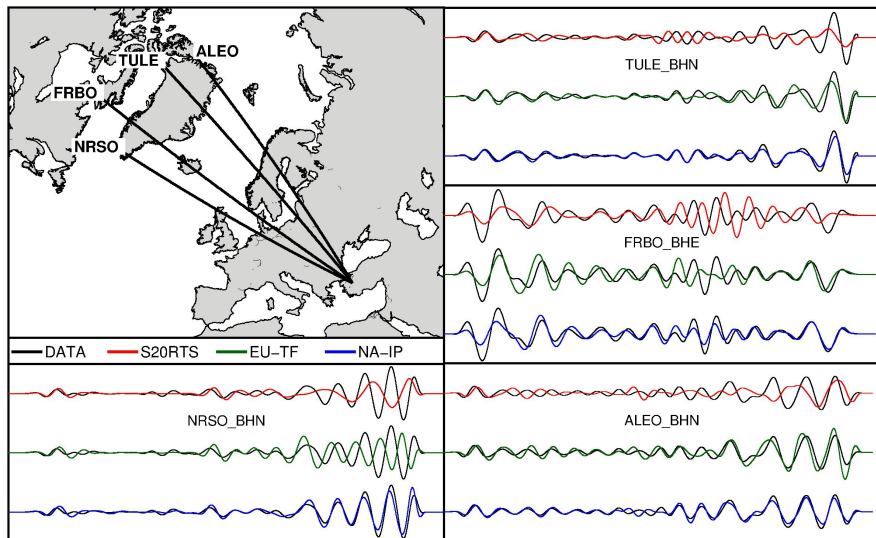
NA-IP

Depth 80 km – $\delta \ln \beta_{sh}$ w.r.t. 4.406 km/s



► correlation with tectonic and geologic features

Validation: Waveform fits compared to S20RTS and EU-TF

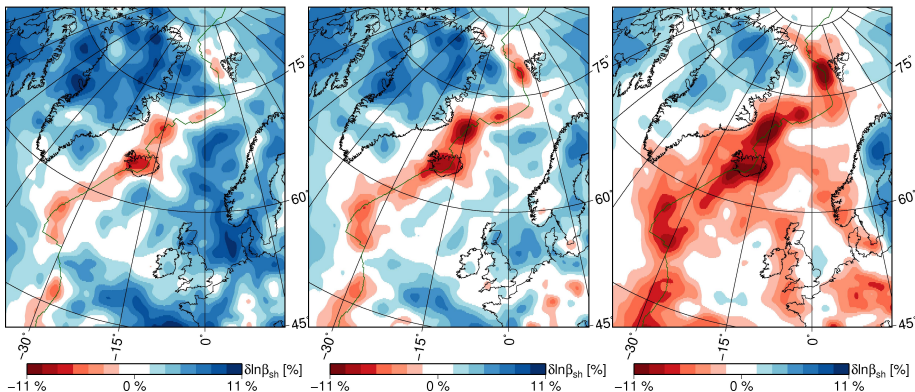


Model NA-IP: Horizontal slices 60, 80, 120 km

Depth 60 km – $\delta \ln \beta_{sh}$ w.r.t. 4.413 km/s

Depth 80 km – $\delta \ln \beta_{sh}$ w.r.t. 4.406 km/s

Depth 120 km – $\delta \ln \beta_{sh}$ w.r.t. 4.391 km/s



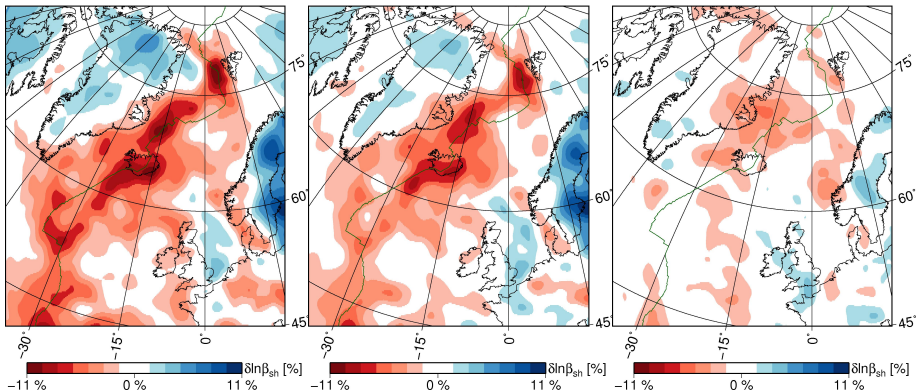
- ▶ separate hotspots beneath Iceland and Jan Mayen
- ▶ deep cratons Greenland + Baltic shield
- ▶ low-velocity layer beneath oceanic lithosphere
- ▶ 'fingers' extend beneath continental lithosphere

Model NA-IP: Horizontal slices 160, 200, 300 km

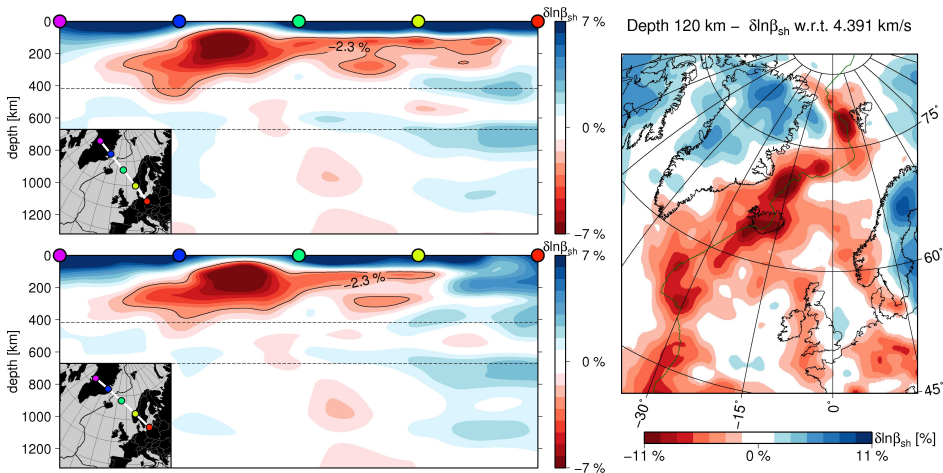
Depth 160 km – $\delta \ln \beta_{sh}$ w.r.t. 4.376 km/s

Depth 200 km – $\delta \ln \beta_{sh}$ w.r.t. 4.409 km/s

Depth 300 km – $\delta \ln \beta_{sh}$ w.r.t. 4.609 km/s



Model NA-IP: Cross-sections along Scandinavian low-velocity 'finger'



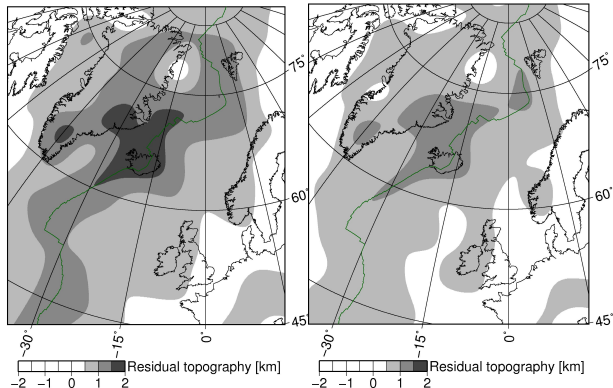
Dynamic topography estimated from gravity

(oceanic)

$Z = 35 \text{ mGal km}^{-1}$ (for sub-aqueous regions)

(continental)

$Z = 50 \text{ mGal km}^{-1}$ (for sub-aerial regions)



(after Jones et al., 2002)

Model NA-IP: Comparison with dynamic topography estimates

Dynamic topography estimated from gravity

S-velocity

(oceanic)

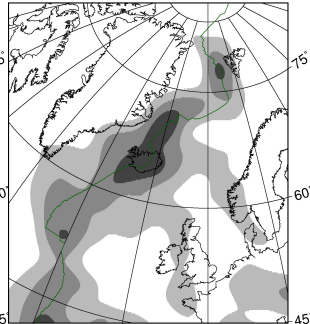
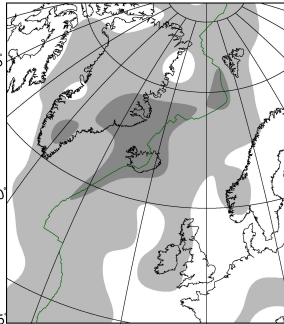
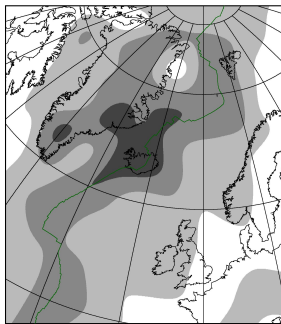
(continental)

(average 100-200 km)

$Z = 35 \text{ mGal km}^{-1}$ (for sub-aqueous regions)

$Z = 50 \text{ mGal km}^{-1}$ (for sub-aerial regions)

Model NA-IP, average 100-200 km



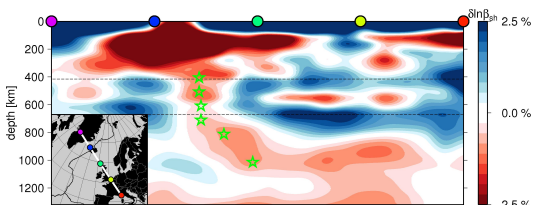
-2 -1 0 1 2 Residual topography [km]

-2 -1 0 1 2 Residual topography [km]

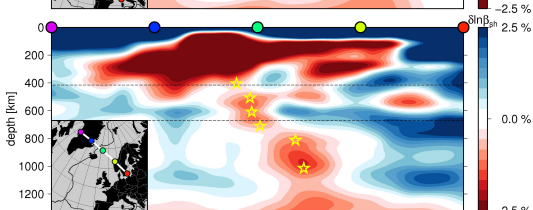
-9% 0% 9% $\delta \ln \beta_{sh}$ (averaged) [%]

(after Jones et al., 2002)

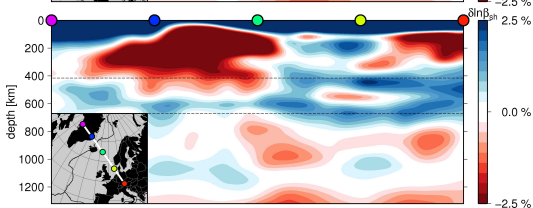
Model NA-IP: Cross-sections of plume conduits



Iceland



Jan Mayen / Kolbeinsey Ridge



In-between Iceland and Jan Mayen

- ▶ We successfully improved the tomographic resolution of the North Atlantic region using adjoint tomography with the instantaneous phase misfit.
- ▶ Iceland and Jan Mayen are separate hotspots, caused by deep plumes that are clearly separated in the upper mantle.
- ▶ A widespread low-velocity layer is found beneath much of the oceanic lithosphere of the North Atlantic.
- ▶ Fingers of the layer extend beneath the continental lithosphere of southern Scandinavia and Britain, supporting the uplift of these regions.
- ▶ Excellent agreement with independent, gravity-based estimates of dynamic topography.

- ▶ Rickers, F., Fichtner, A. and Trampert, J., 2013. **The Iceland - Jan Mayen plume system and its impact on mantle dynamics in the North Atlantic region: Evidence from full-waveform inversion.** Earth Planetary Science Letters 367, 39–51

- ▶ Rickers, F., Fichtner, A. and Trampert, J., 2012. **Imaging mantle plumes with instantaneous phase measurements of diffracted waves.** Geophysical Journal International 190, 650–664