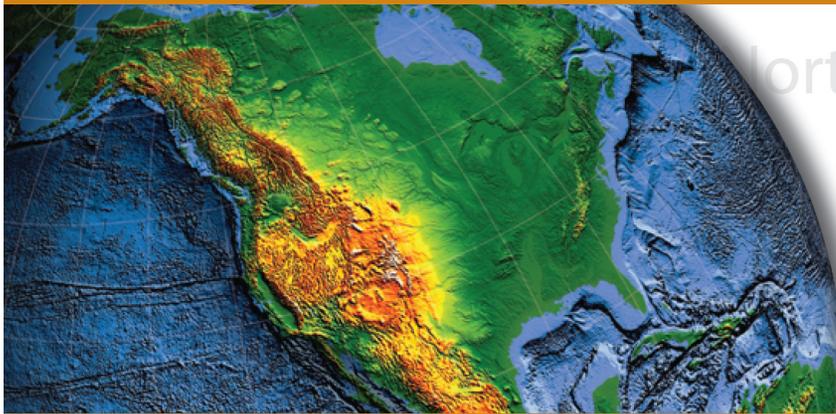
A map of North America showing noise tomography results. The map uses a color scale from blue (low velocity) to red (high velocity) to represent seismic wave velocity variations. White contour lines are overlaid on the map, indicating specific velocity values such as 1500, 3000, 4500, 6000, and 7500. The map also shows state boundaries and major geographical features like the Rocky Mountains and the Great Plains. The text is centered over the map.

Noise tomography of North America using USArray

Göran Ekström

1. A few words about EarthScope and USArray
2. Surface-wave studies of the crust and mantle
3. Tomography using noise and Aki's method
4. Remarkable images of US crust (and basins)!



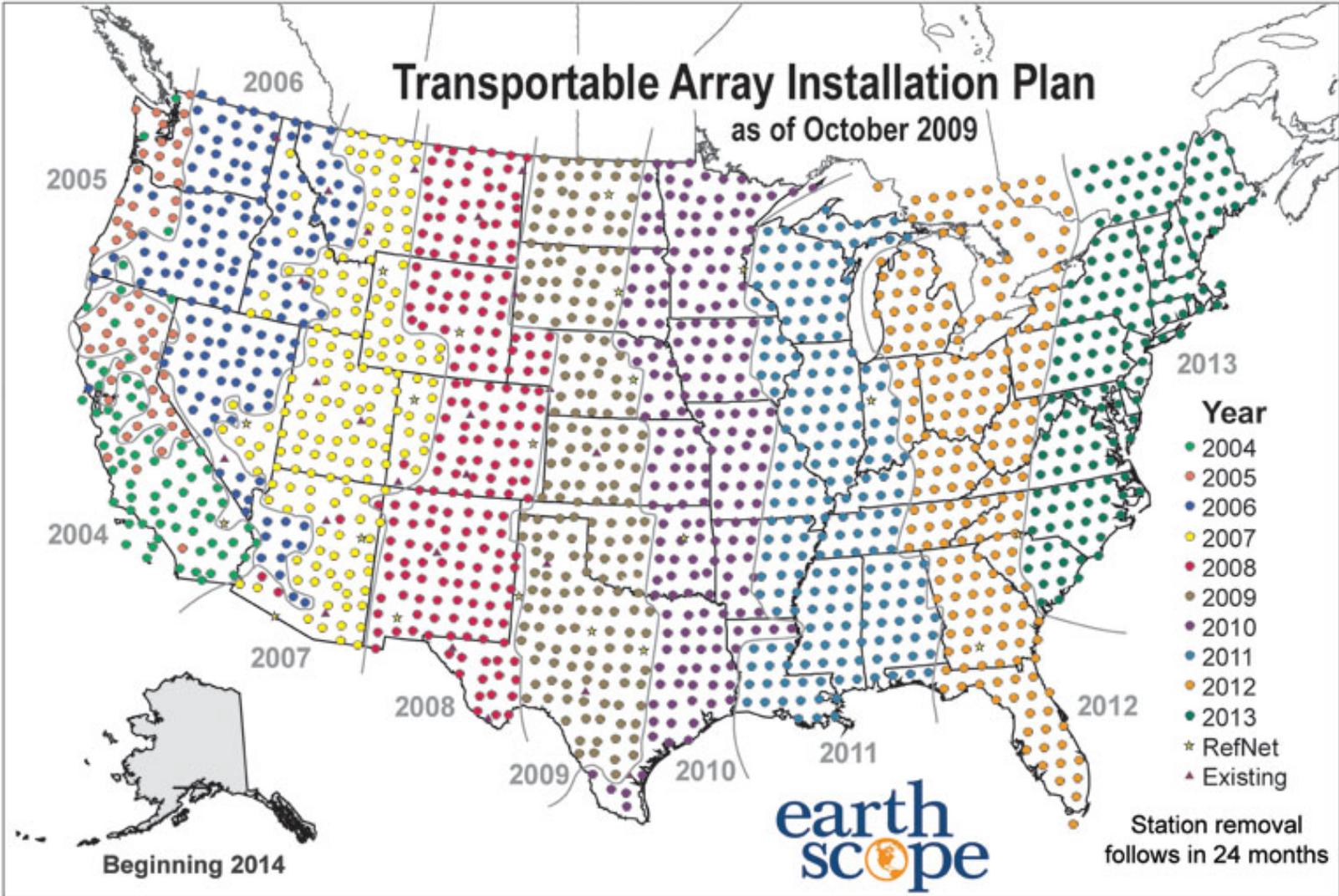
North American Continent

Unlocking the Secrets of the North American

Secrets of the North American

Unlocking the Secrets of the North American Continent

An EarthScope Science Plan for 2010–2020



2-year deployments - 2000 sites
uniform instrumentation
systematic quality control



The speed of surface waves (Love and Rayleigh) depend on the shallow structure of the Earth

BULLETIN OF THE GEOLOGICAL SOCIETY OF AMERICA
VOL. 70, PP. 229-244, 11 FIGS. MARCH 1959

DETERMINATION OF CRUSTAL STRUCTURE FROM PHASE VELOCITY OF RAYLEIGH WAVES PART III: THE UNITED STATES

BY MAURICE EWING AND FRANK PRESS

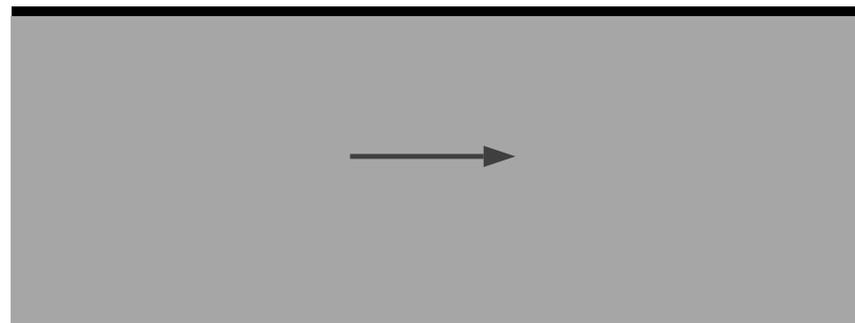
ABSTRACT

Variations in phase velocity of Rayleigh waves from the Samoa earthquake of April 14, 1957 are reported for the United States. These variations are correlated with topography and Bouguer gravity anomaly on a continental scale, demonstrating regional isostatic compensation. The correlation of phase-velocity variations with crustal-thickness changes is justified, and permits specification of the mechanism of compensation as the regional Airy system.

Regional average crustal thicknesses are: Peninsular Ranges and Southwestern Desert, 40 km; Basin and Range Province, 48 km; Rocky Mountains, 47 km; Interior Plains, 35-41 km; Appalachian Mountains, 40 km.

Sensitivity of surface wave velocities to elastic structure at depth

200 seconds



300 km

20 seconds



30 km

Noise cross correlations - two approaches:

Diffuse and equipartitioned 3-D wavefield:

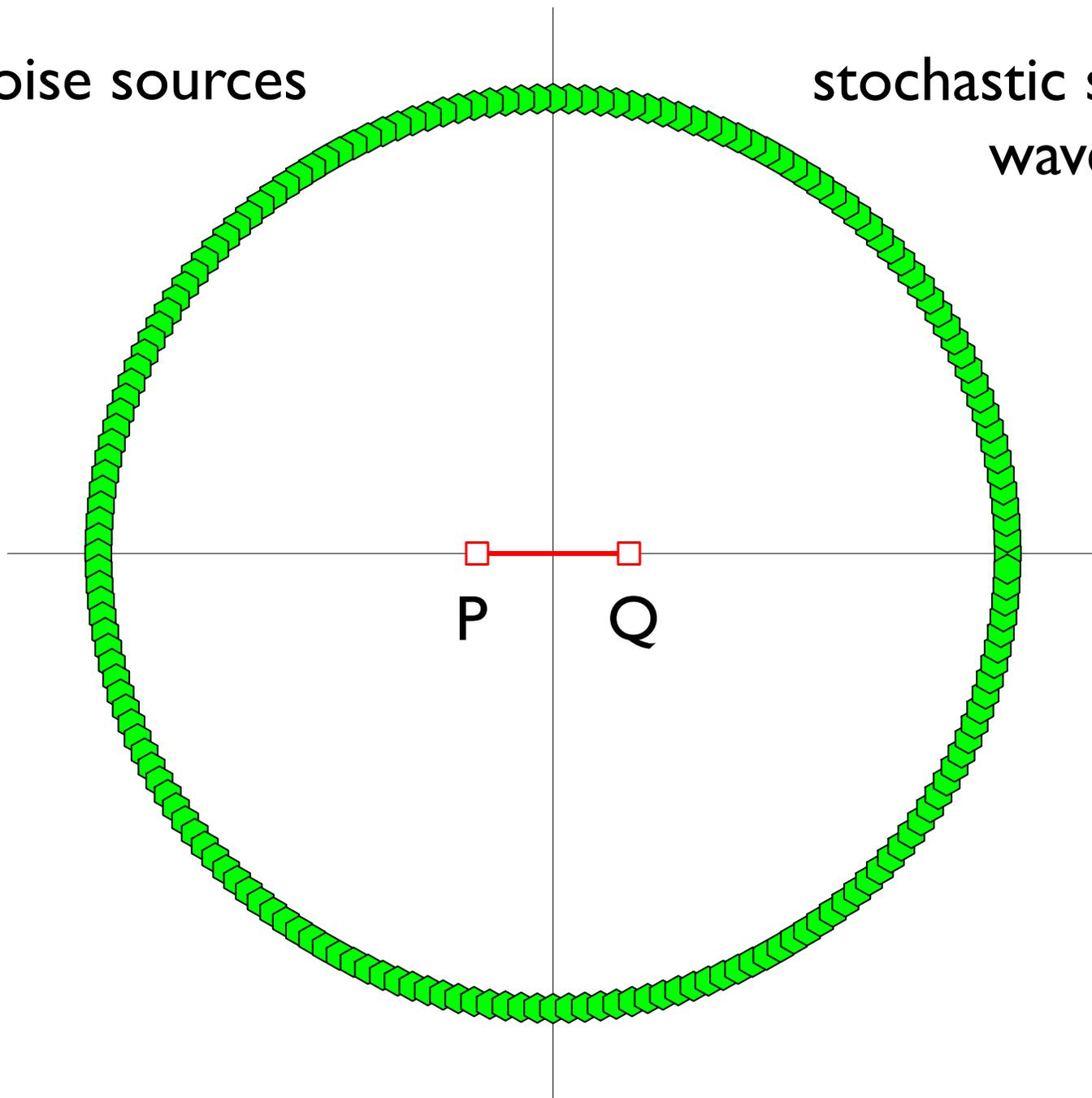
- leads to Green function (Campillo, ...)

Lots of traveling surface waves:

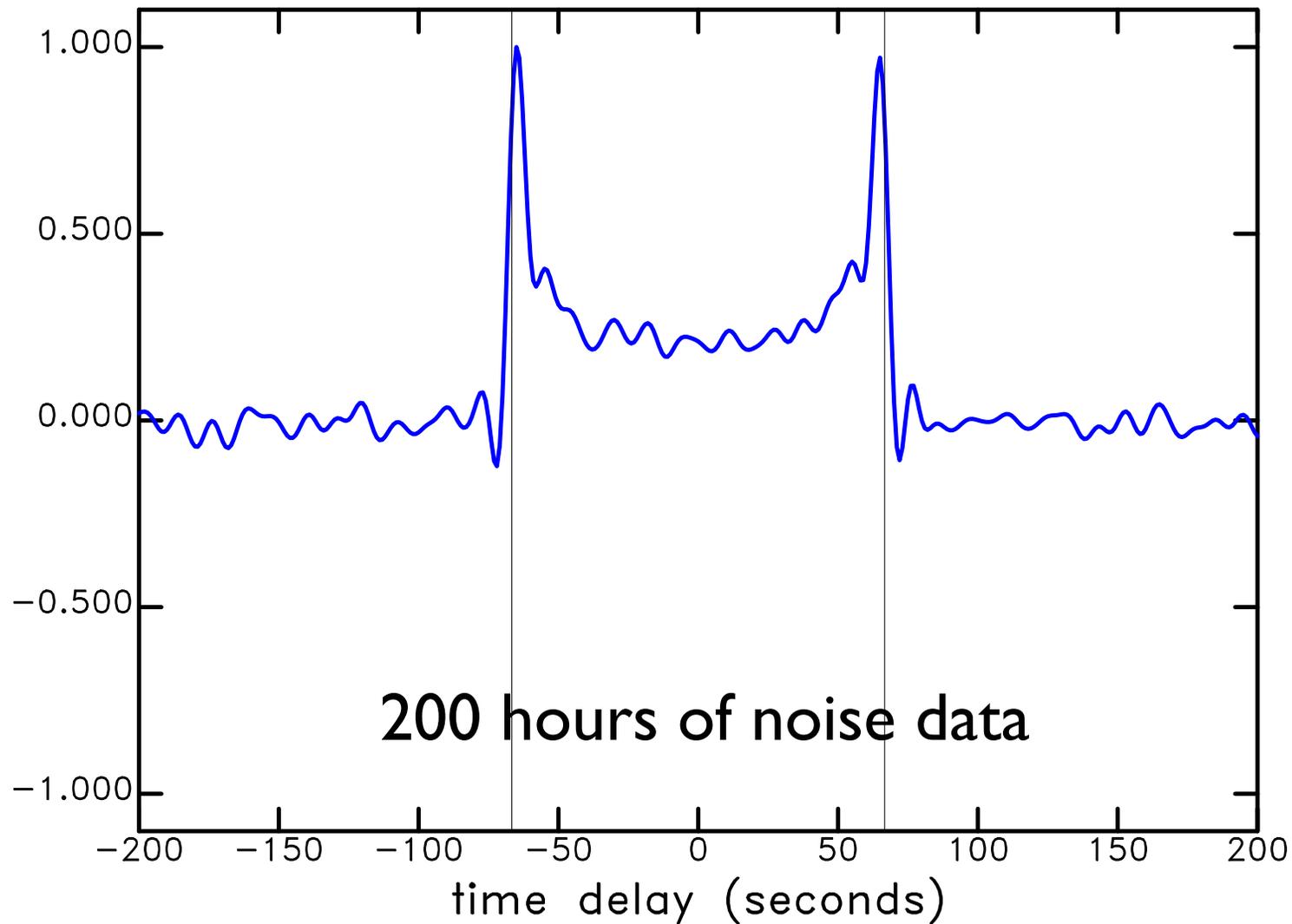
- leads to Bessel function (Aki, 1957)

180 noise sources

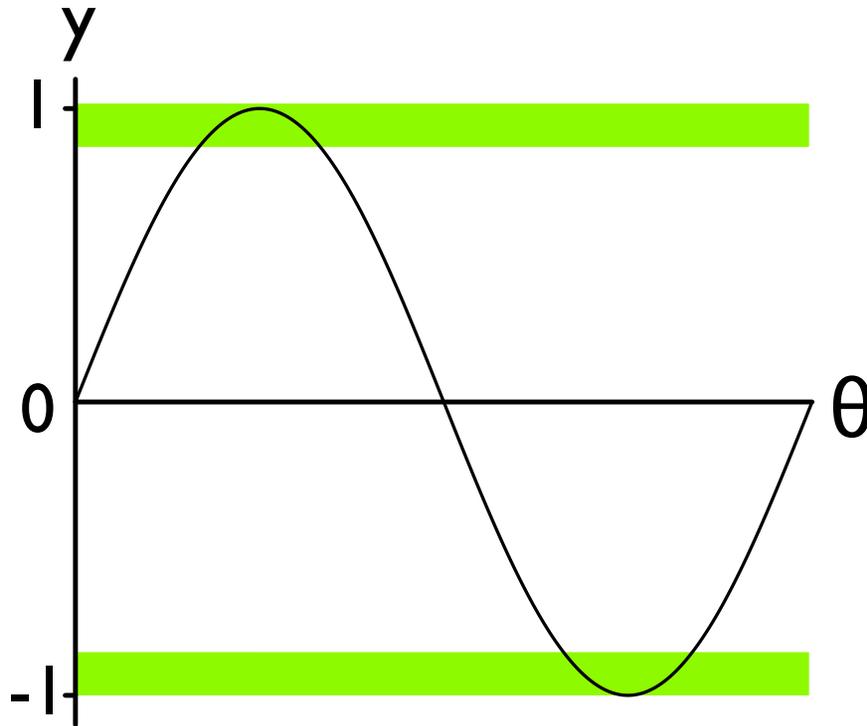
stochastic surface
waves



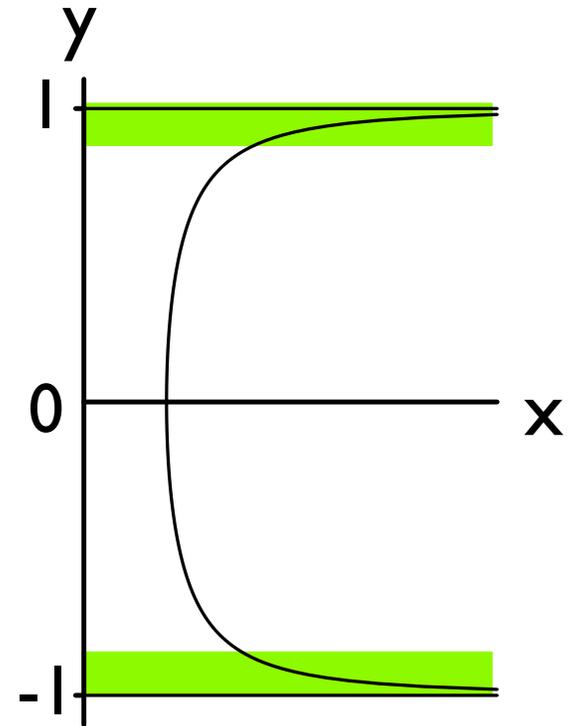
Cross-correlation function, P and Q



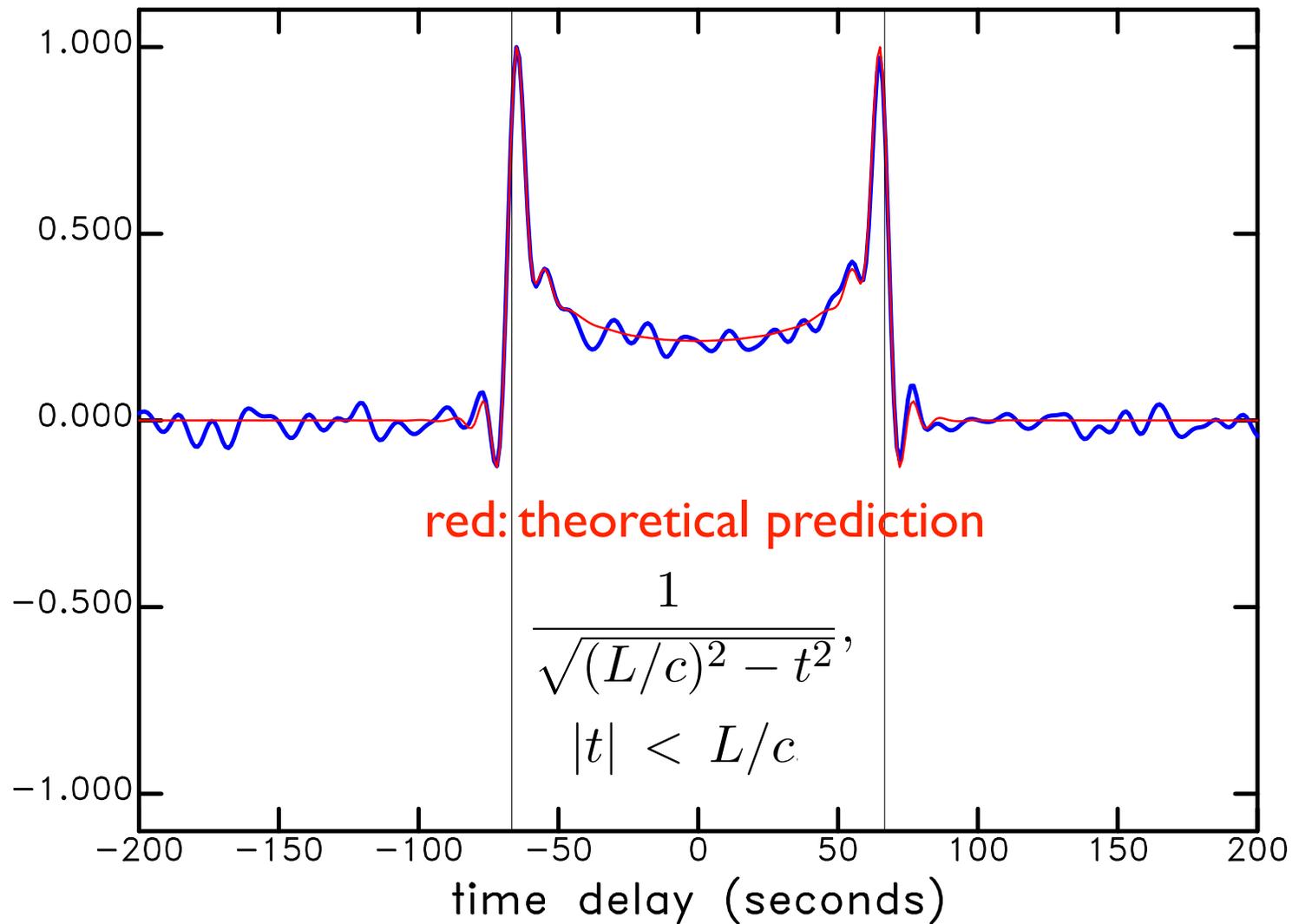
sine function
 $y = \sin(\theta)$



pdf of sine function
 $x = p(y)$



Cross-correlation function, P and Q



What about the Fourier transform?

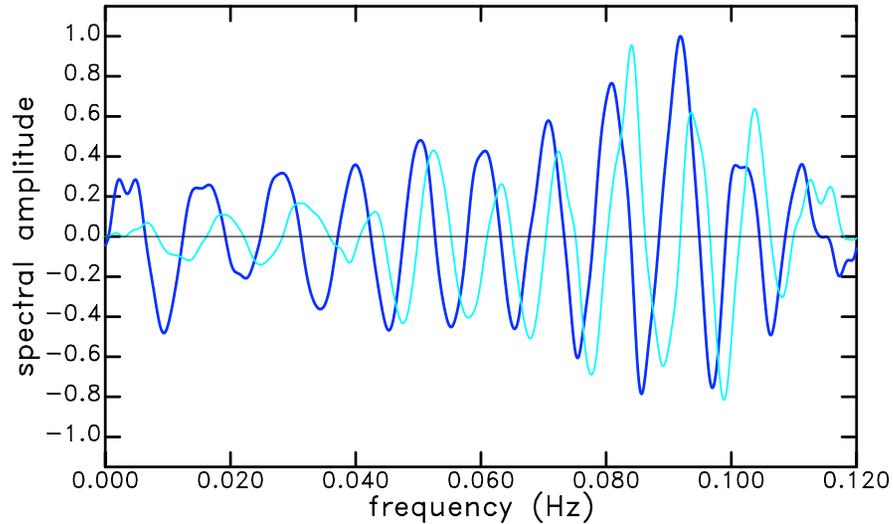
$$\frac{1}{\sqrt{(L/c)^2 - t^2}} \longrightarrow J_0\left(\frac{\omega L}{c}\right)$$

$$\bar{\rho}(r, \omega_0) = J_0 \left(\frac{\omega_0}{c(\omega_0)} r \right)$$

“This formula clearly indicates that if one measures $\bar{\rho}(r, \omega_0)$ for a certain r and for various ω_0 's, he can obtain the function $c(\omega_0)$, i.e., the dispersion curve of the wave for the corresponding range of frequency ω_0 ”.

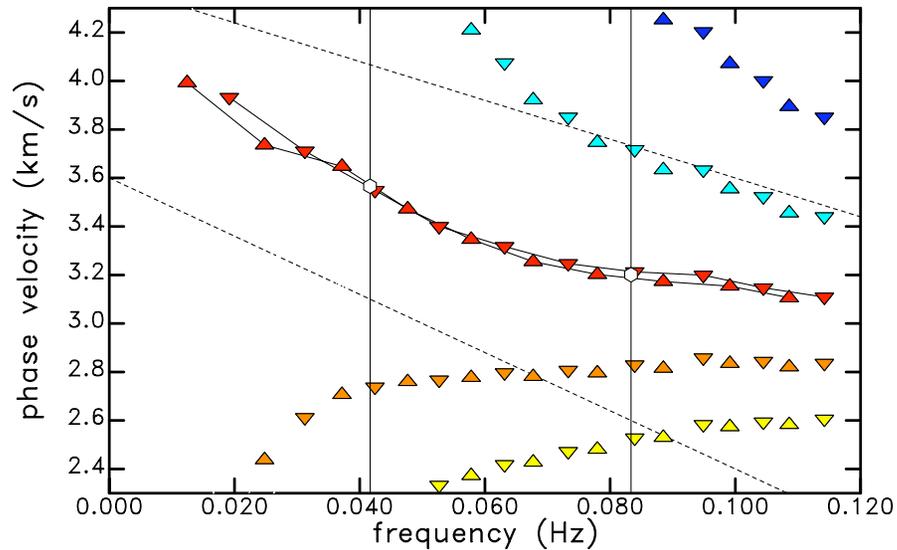
Aki, 1957

Matching zero crossings for dispersion



D07A-B04A
282 km

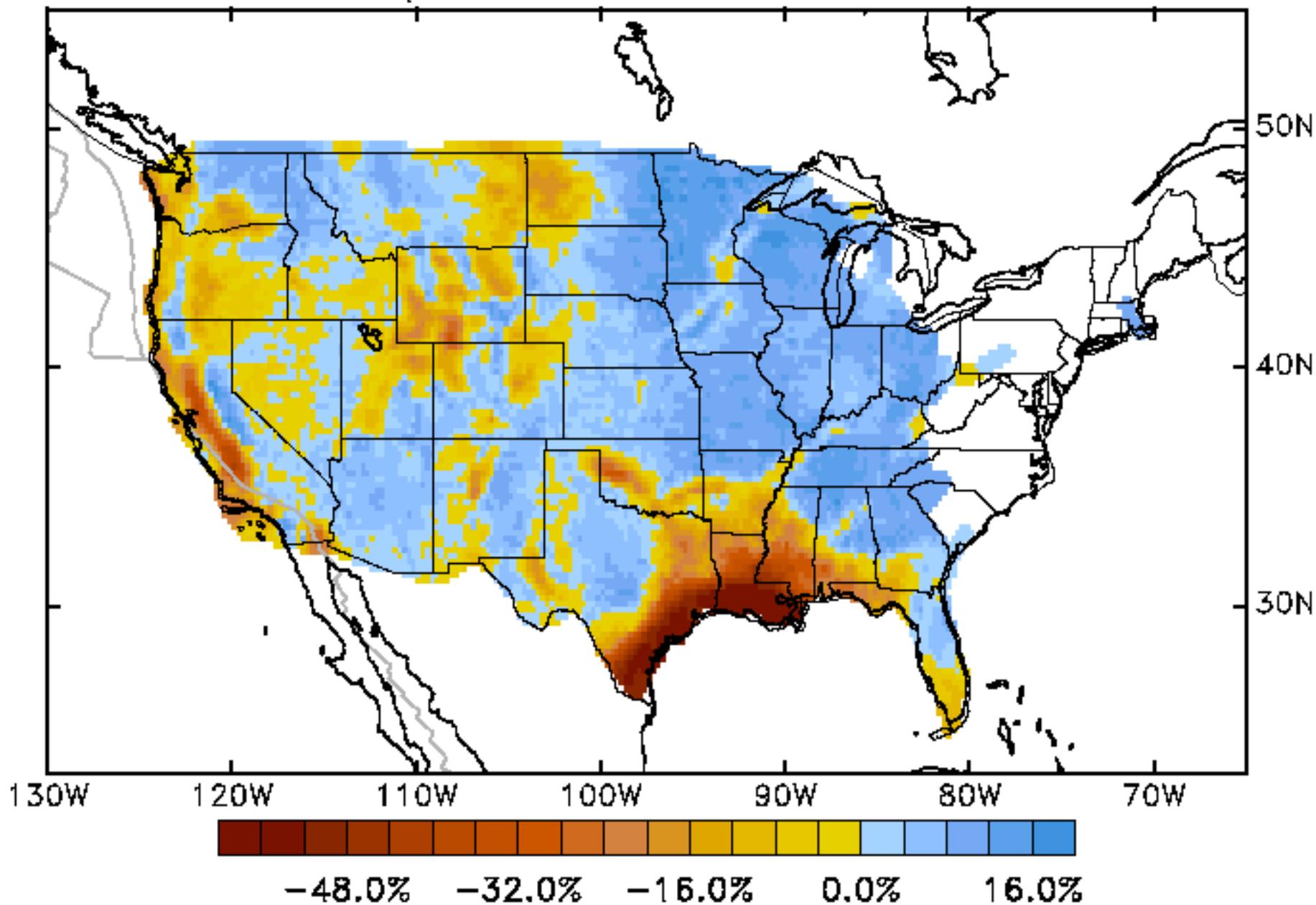
$$c(\omega_n) = \frac{\omega_n r}{Z_n}$$



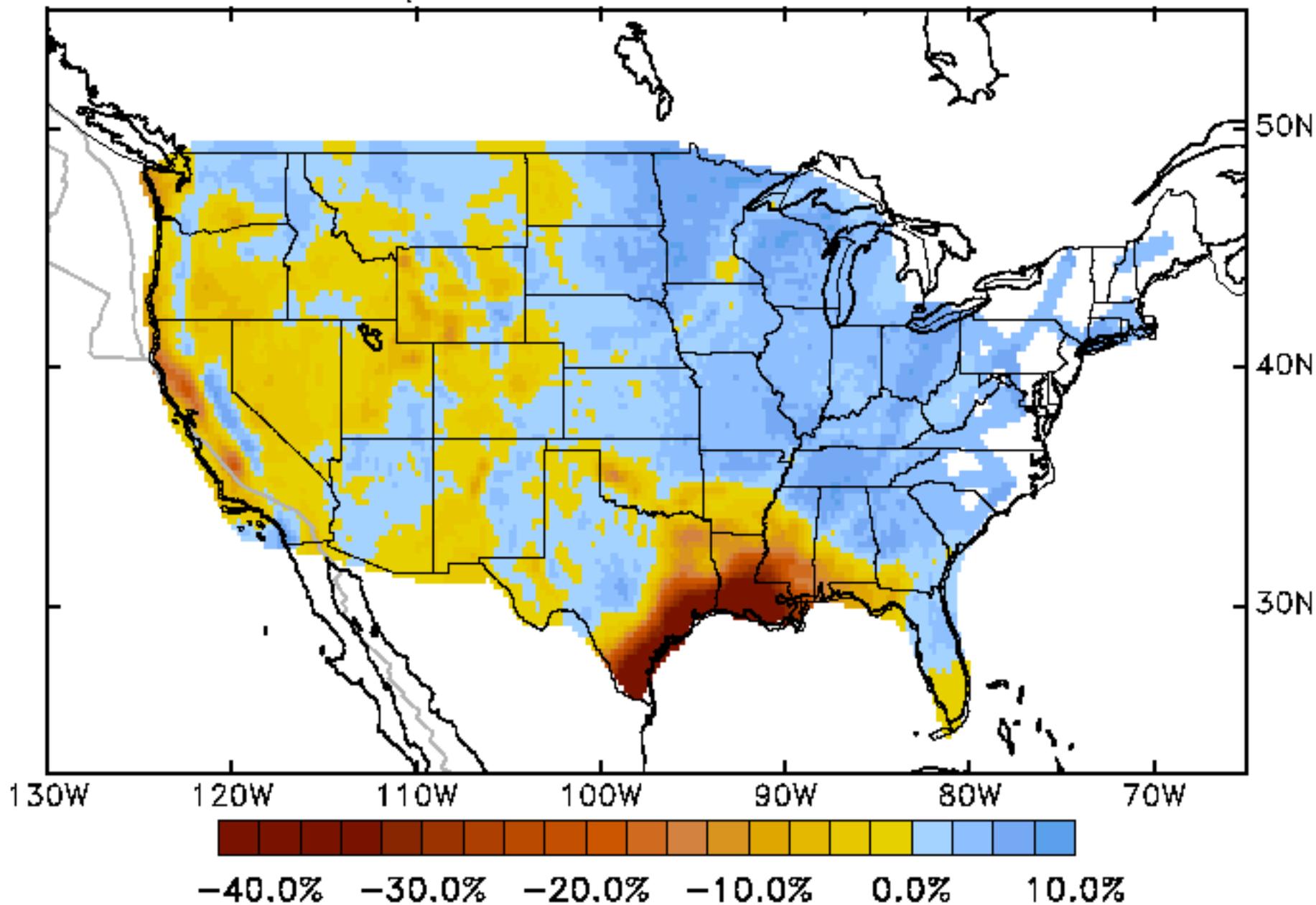
Recipe for tomographic success:

1. Correlate continuous recorded signals at all pairs of USArray stations in 4-h windows (note - this is a big calculation)
 2. Stack all correlation functions for each pair
 3. Determine zero crossings of stacked cross-correlation spectra
 4. Determine phase velocities using Aki's formula
 5. Invert phase-velocity observations to determine phase-velocity maps
- (no one-bit, no whitening, no nuthin')

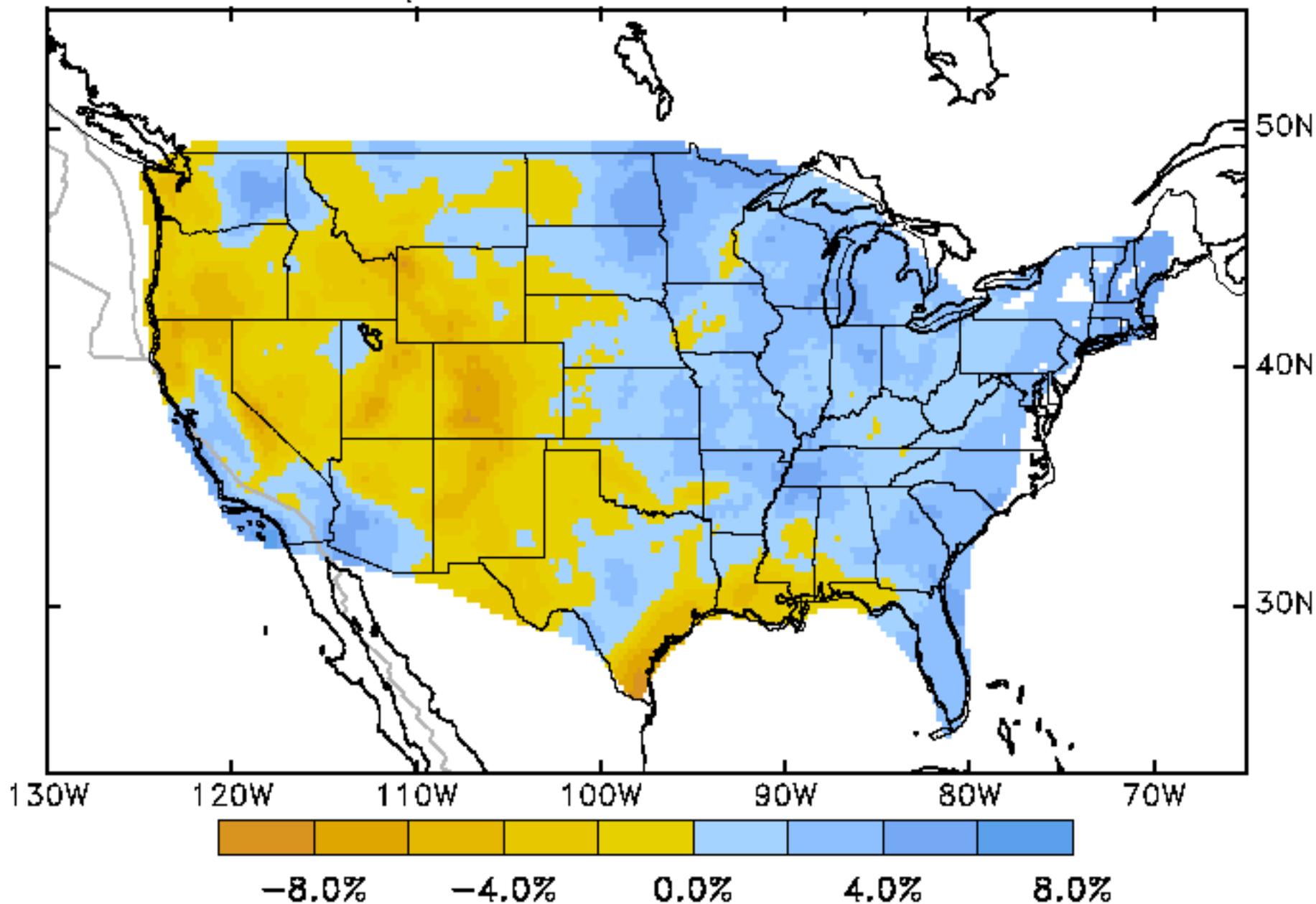
R005.1212 † bo.pix



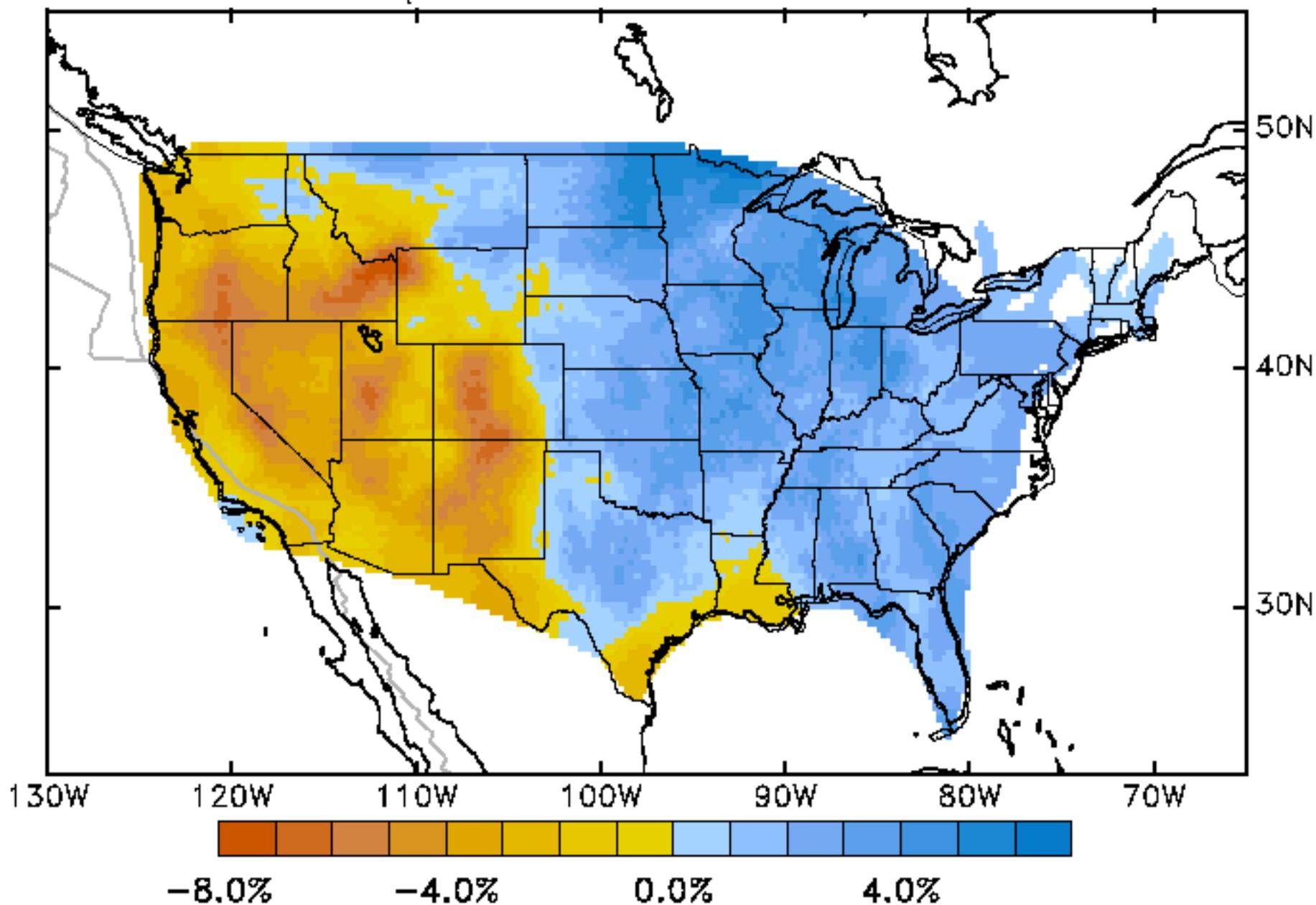
R010.1212 † bo.pix



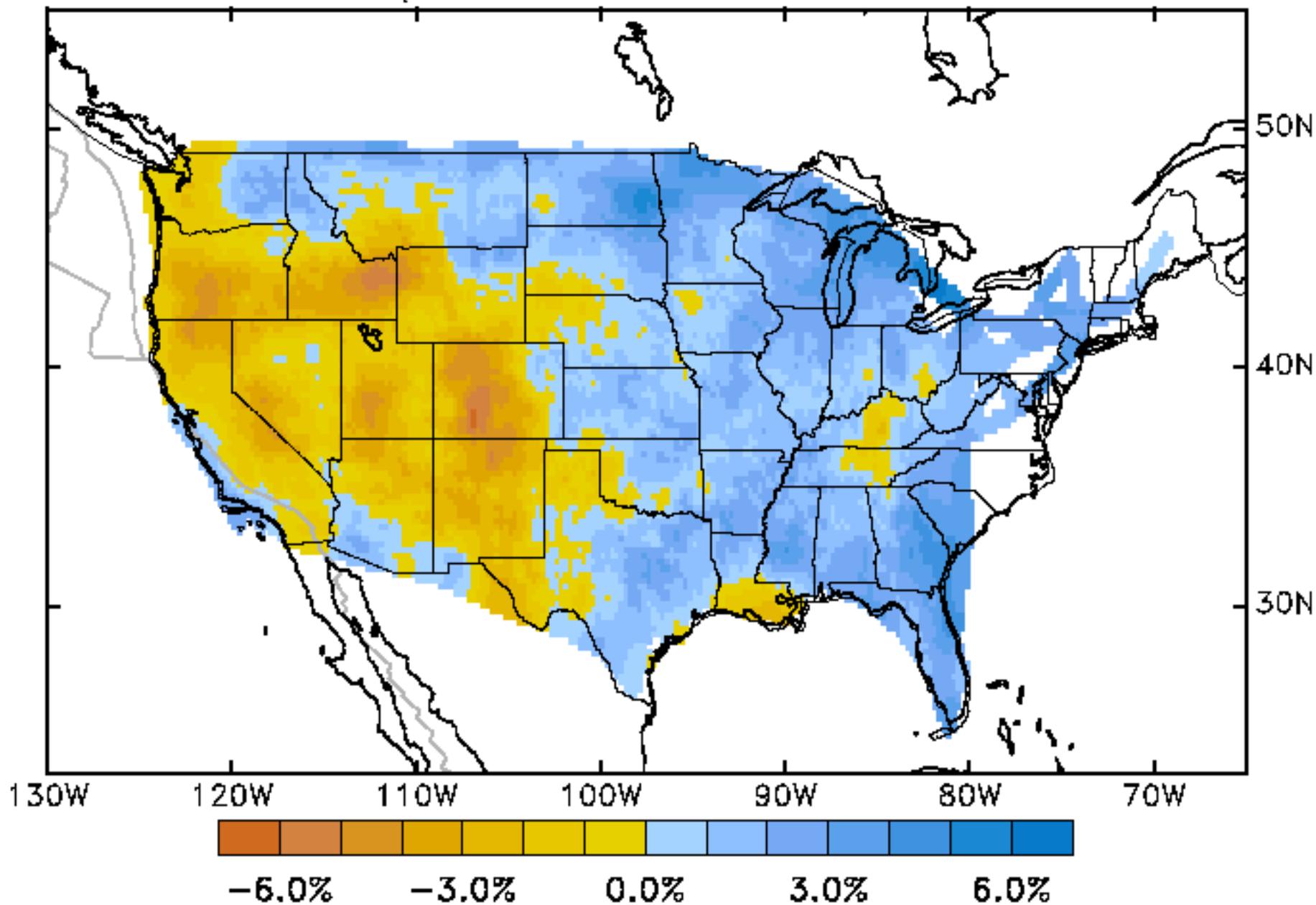
R020.1212 † bo.pix



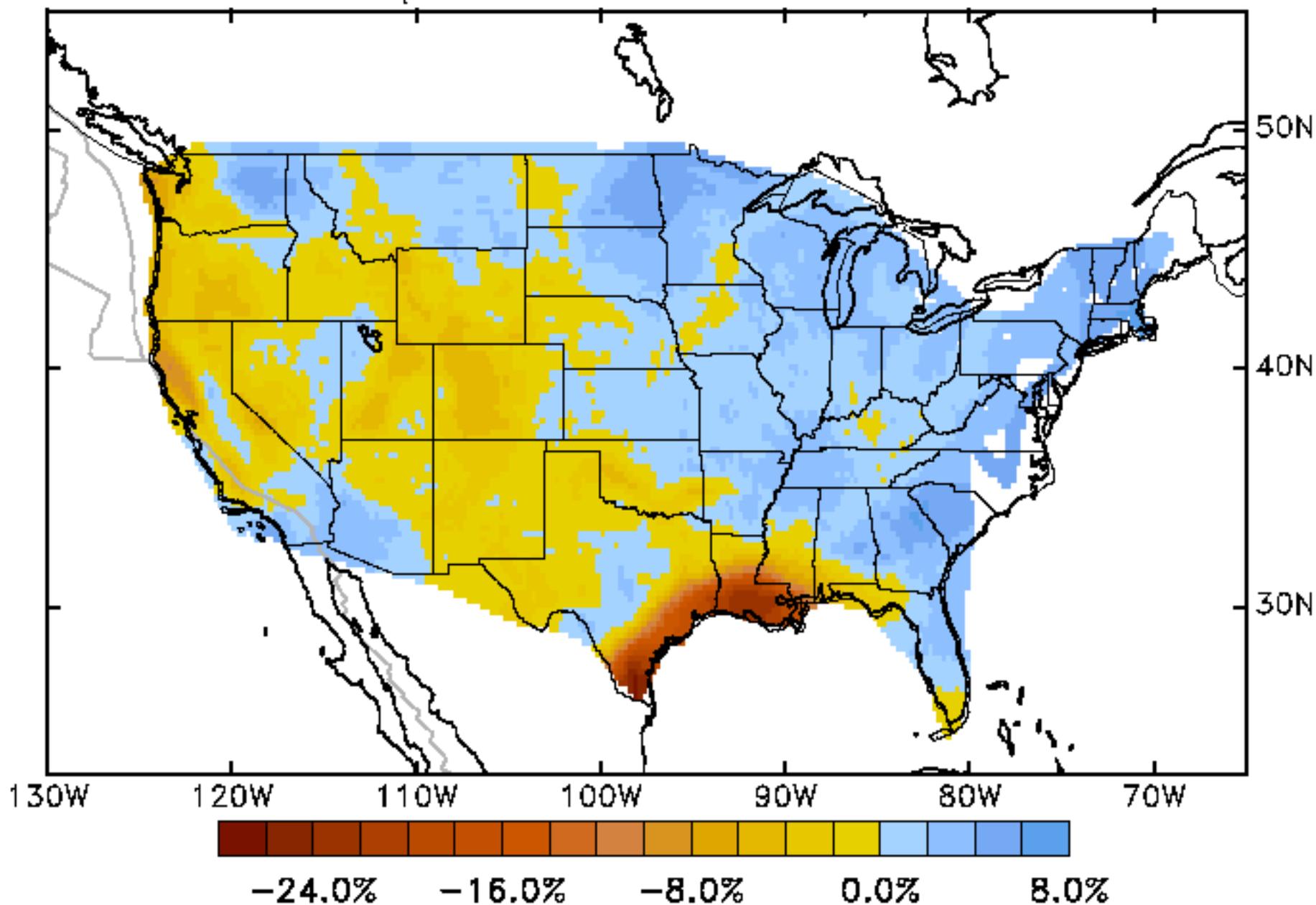
R040.1212 † bo.pix



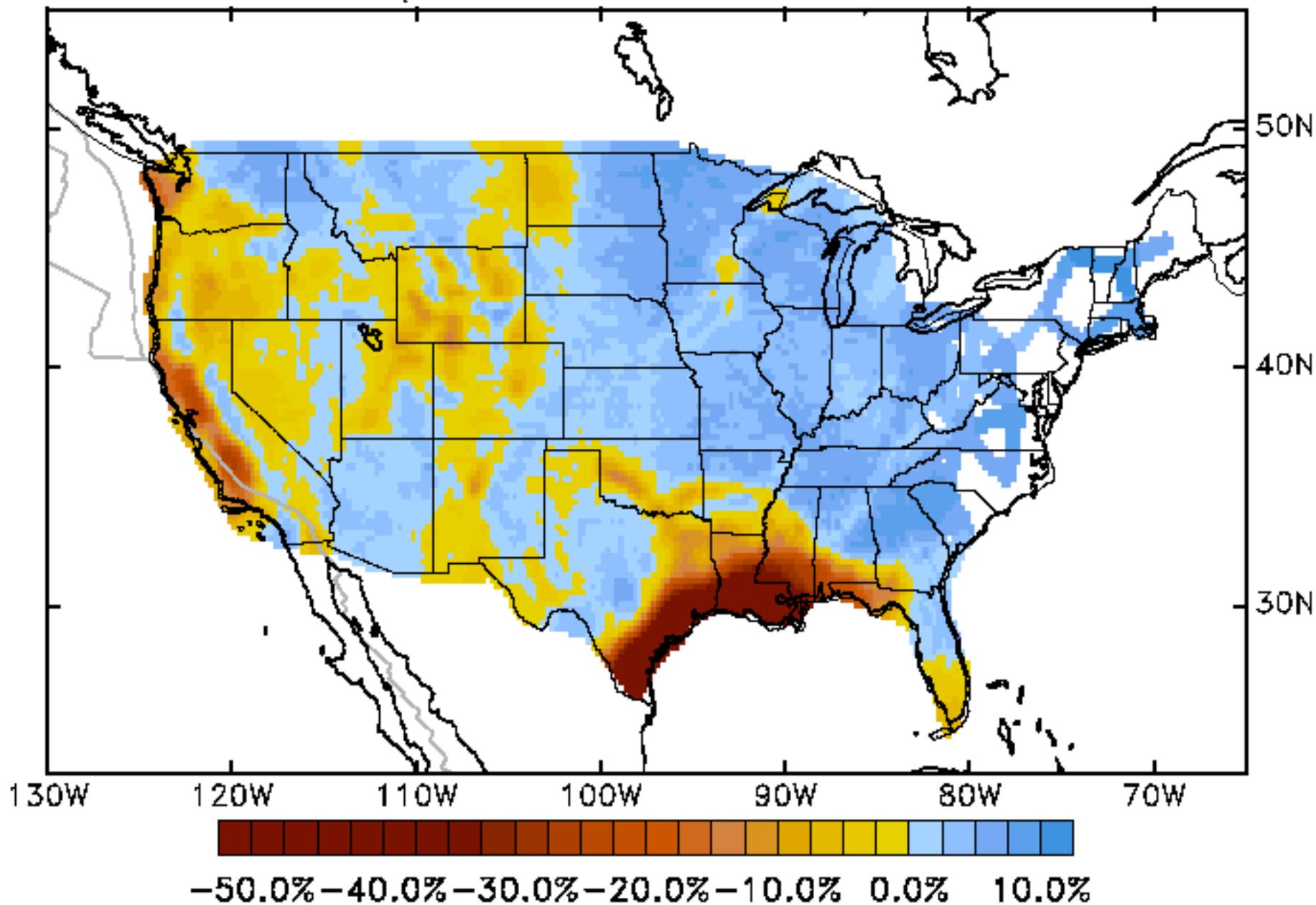
L040.1212 † bo.pix



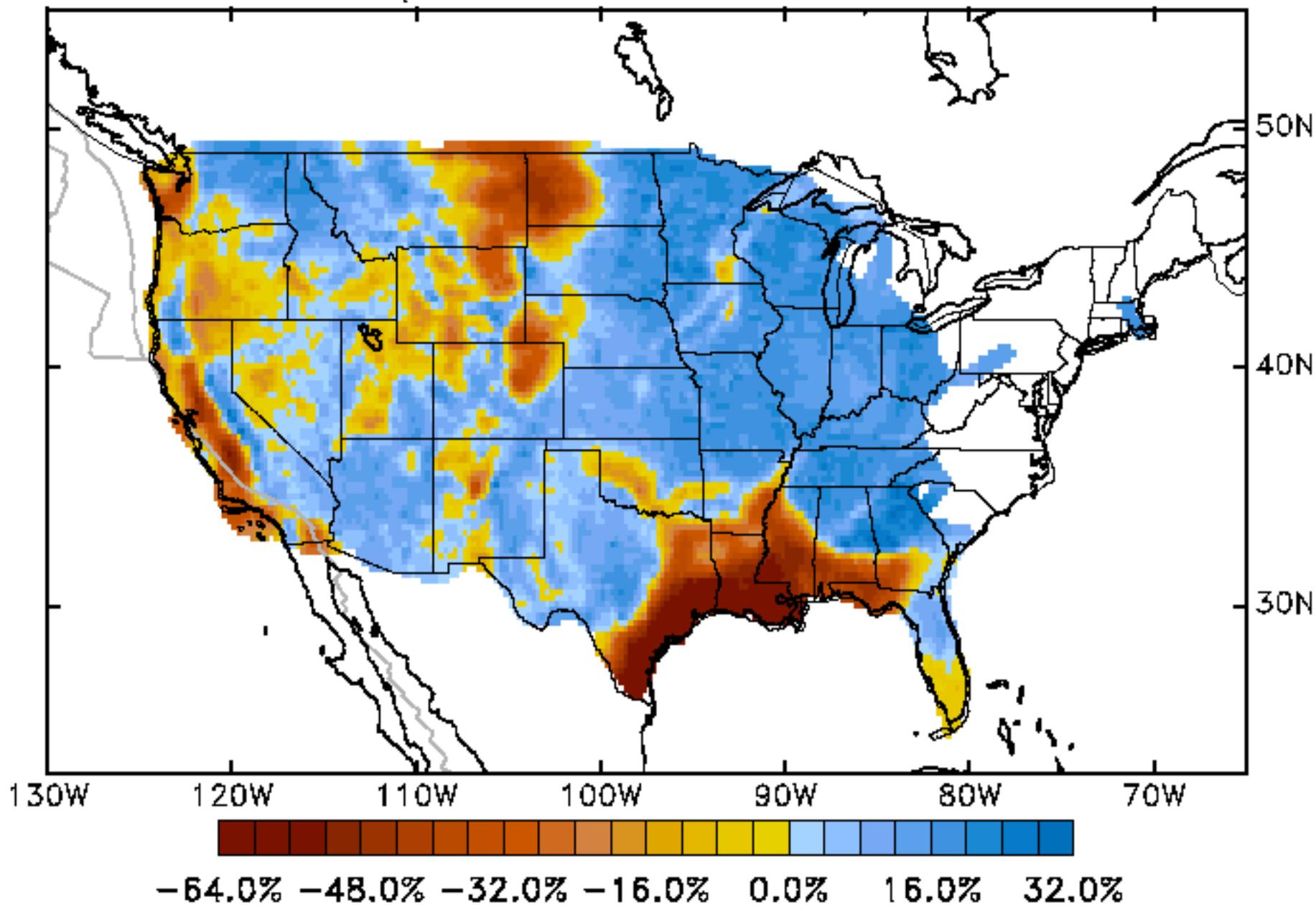
L020.1212 † bo.pix



L010.1212 † bo.pix



L005.1212 † bo.pix

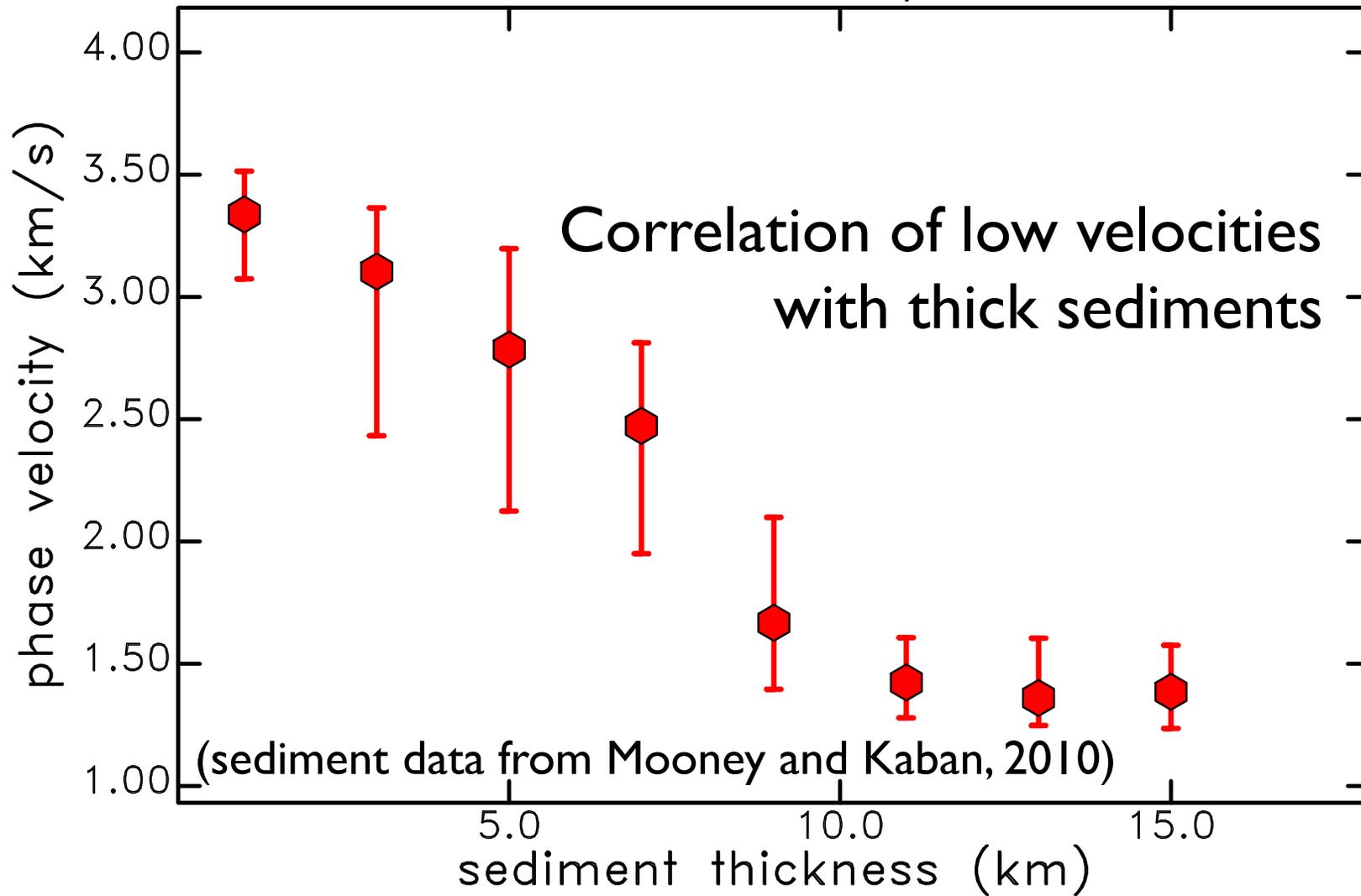


What are we looking at?

Mainly elastic effects of the crust

Including strong signals of slow sediments

L005.1212 t.pix



Correlation of low velocities
with thick sediments

(sediment data from Mooney and Kaban, 2010)

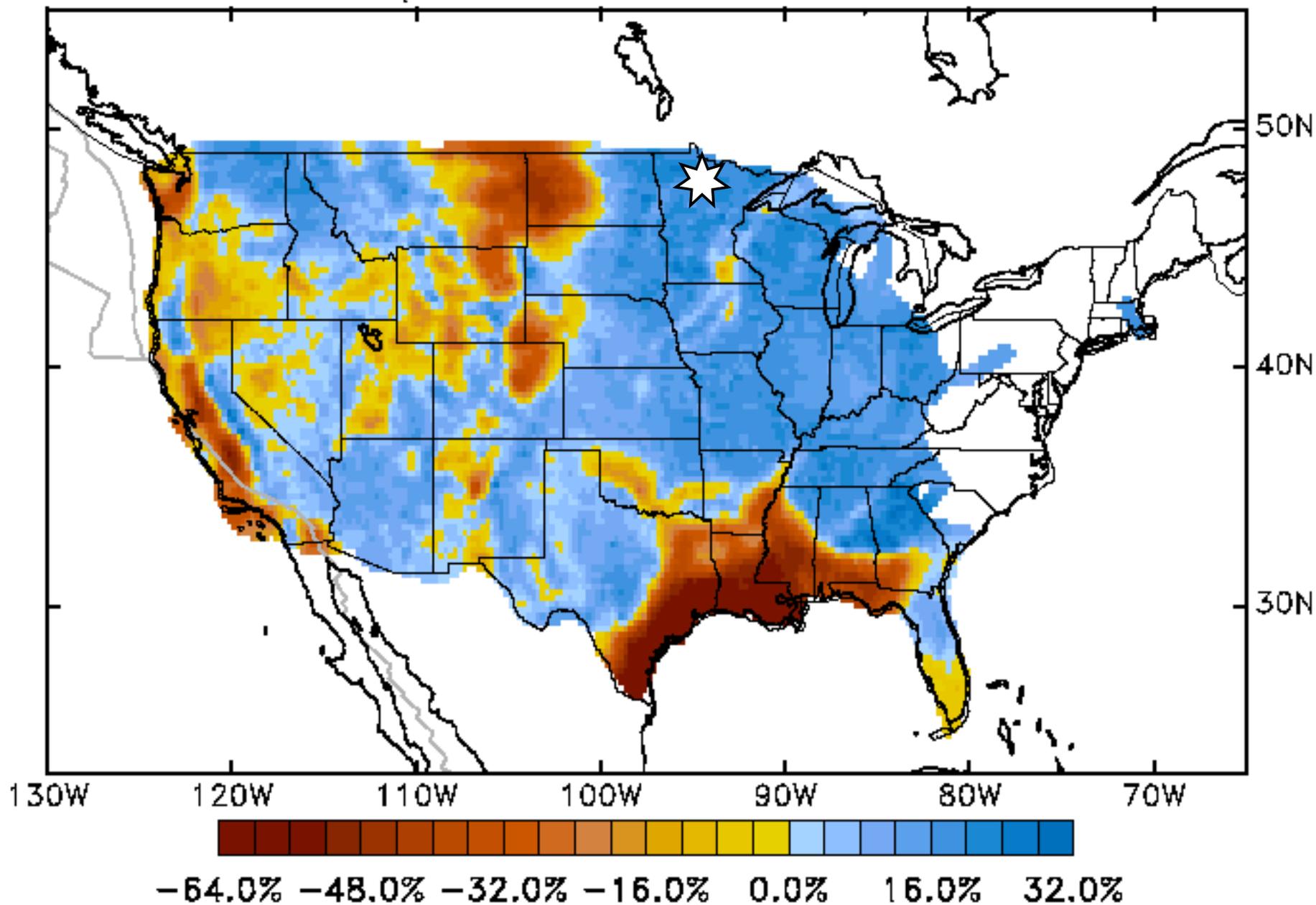
What are we looking at?

Mainly elastic effects of the crust

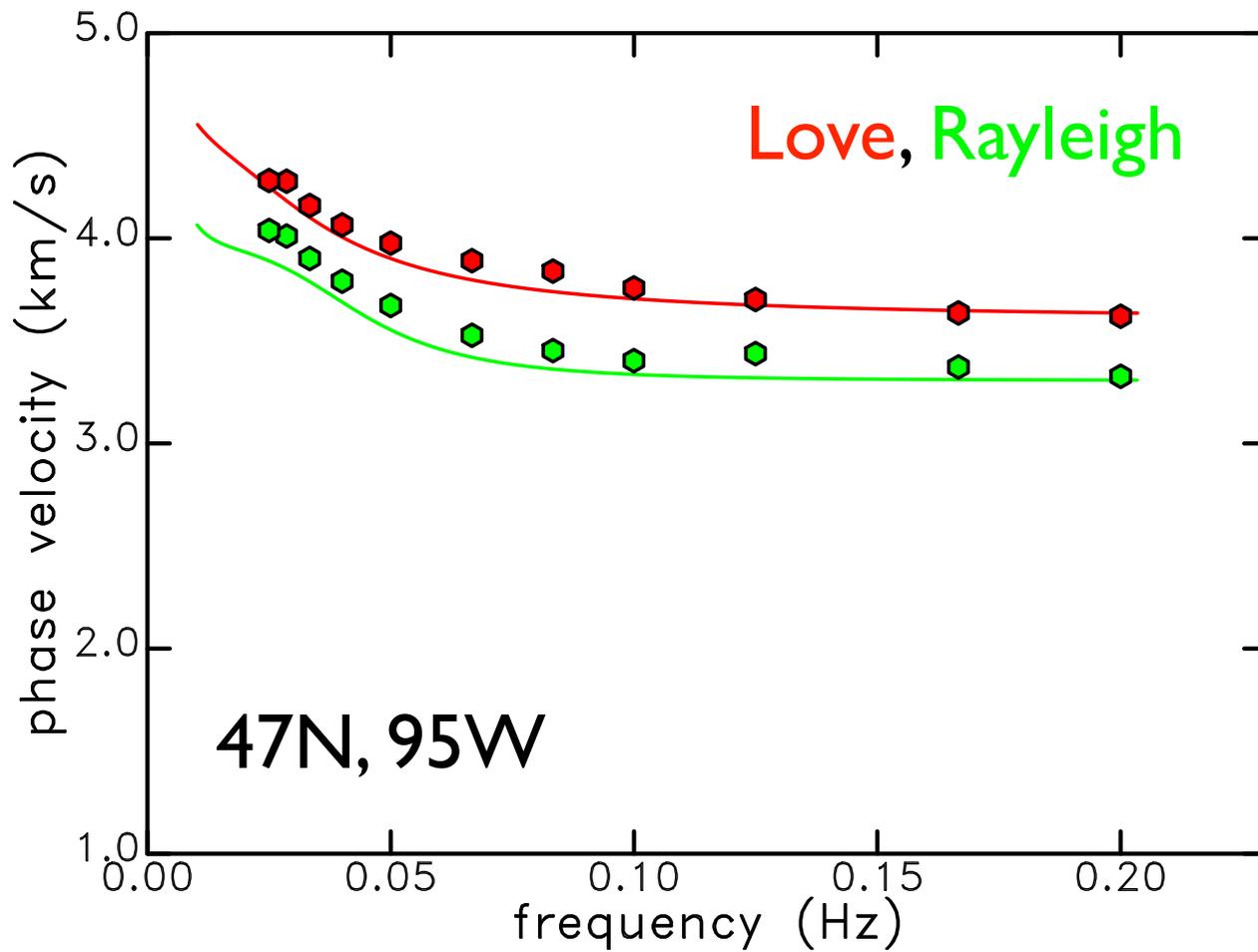
Including strong signals of slow sediments

Compare with predictions from CRUST 2.0

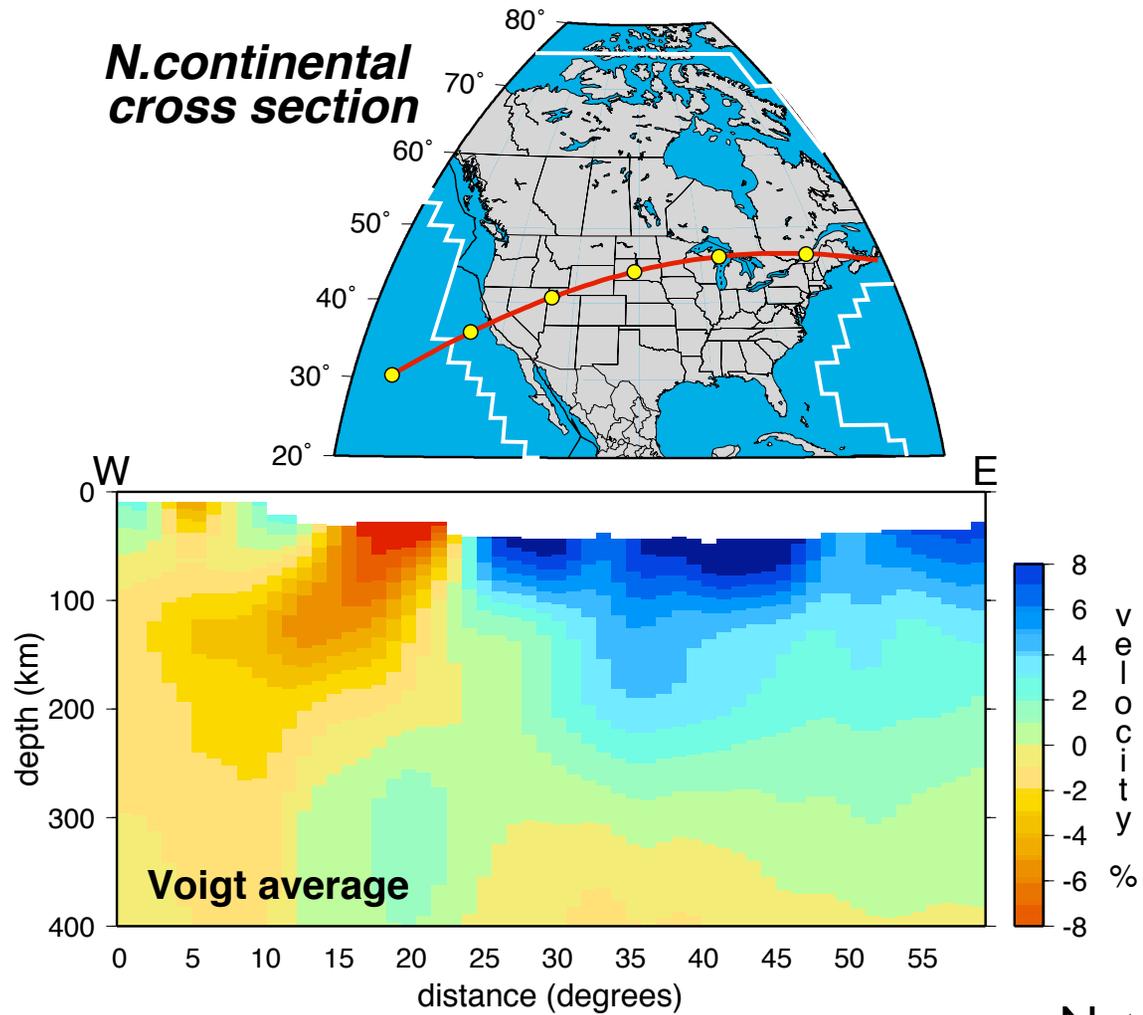
L005.1212 † bo.pix



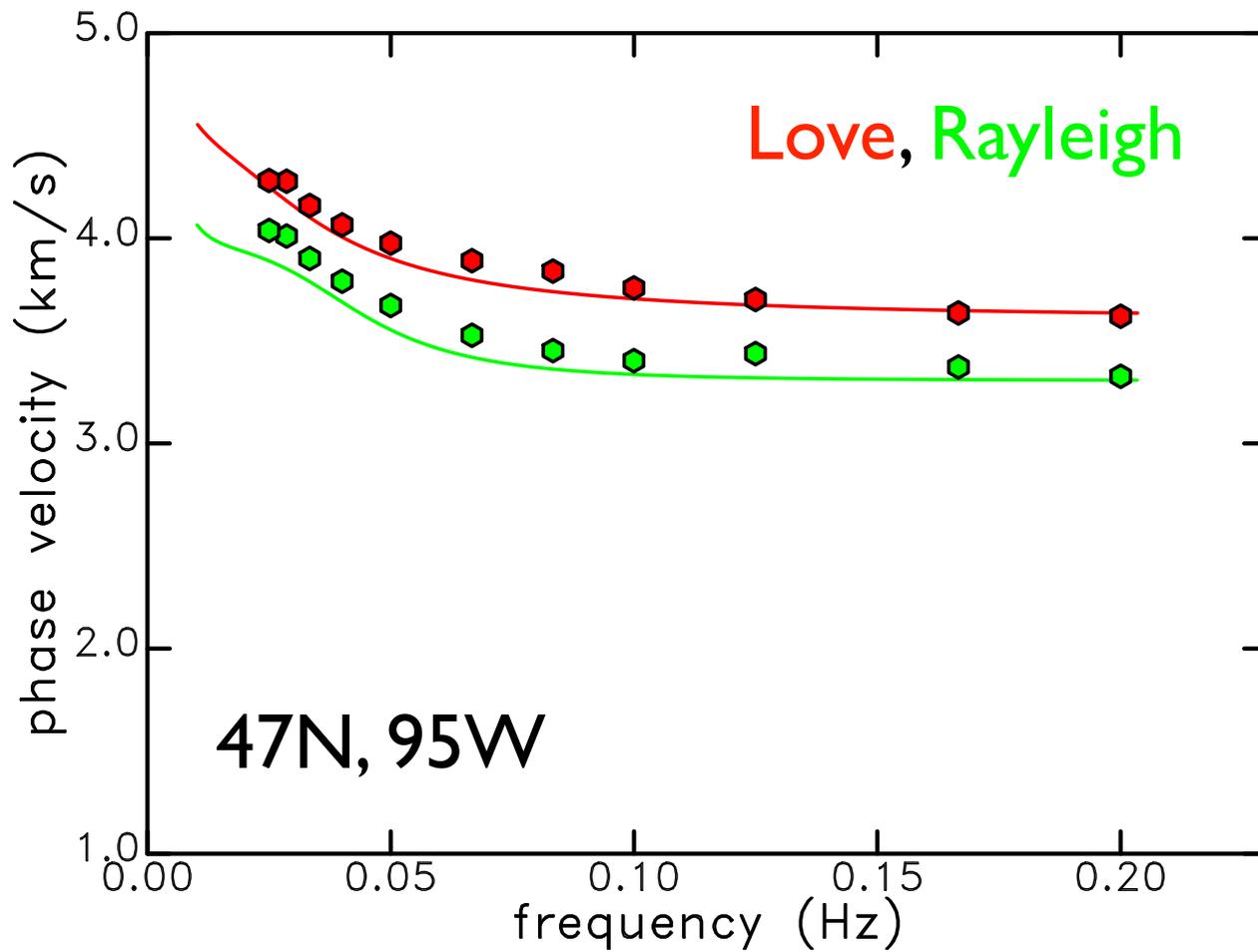
Observations and CRUST 2.0



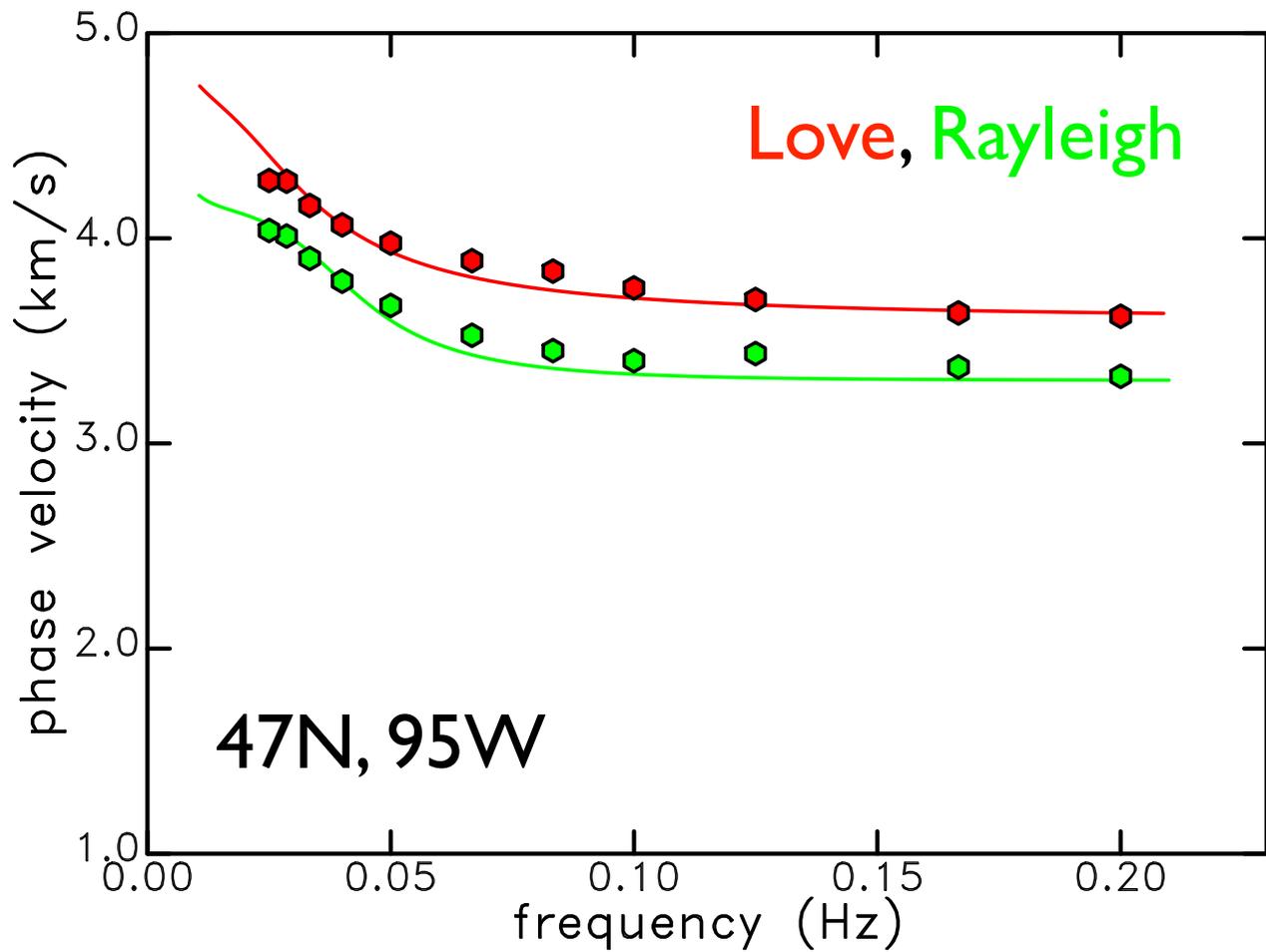
Mantle structure from surface waves



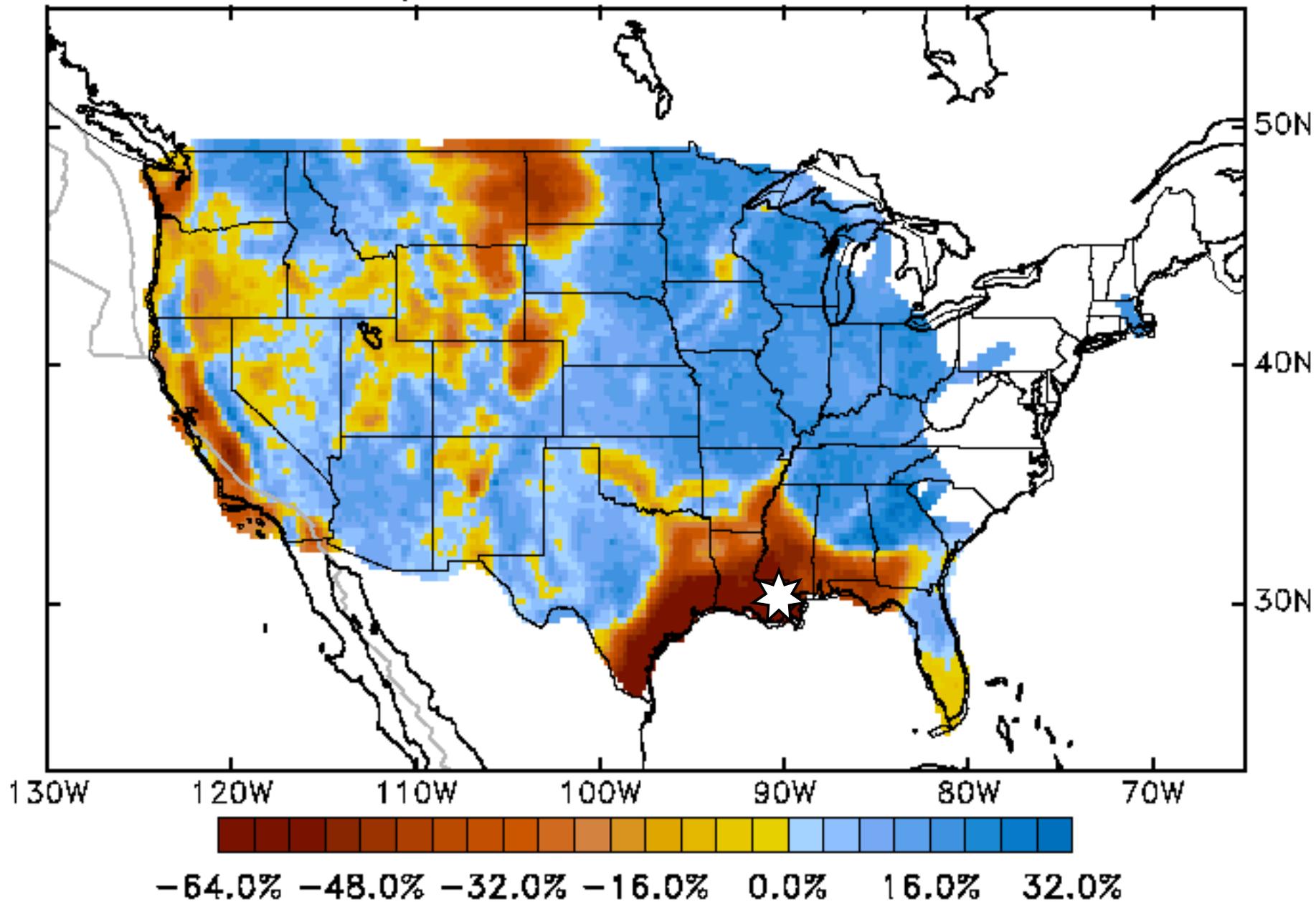
Observations and CRUST 2.0



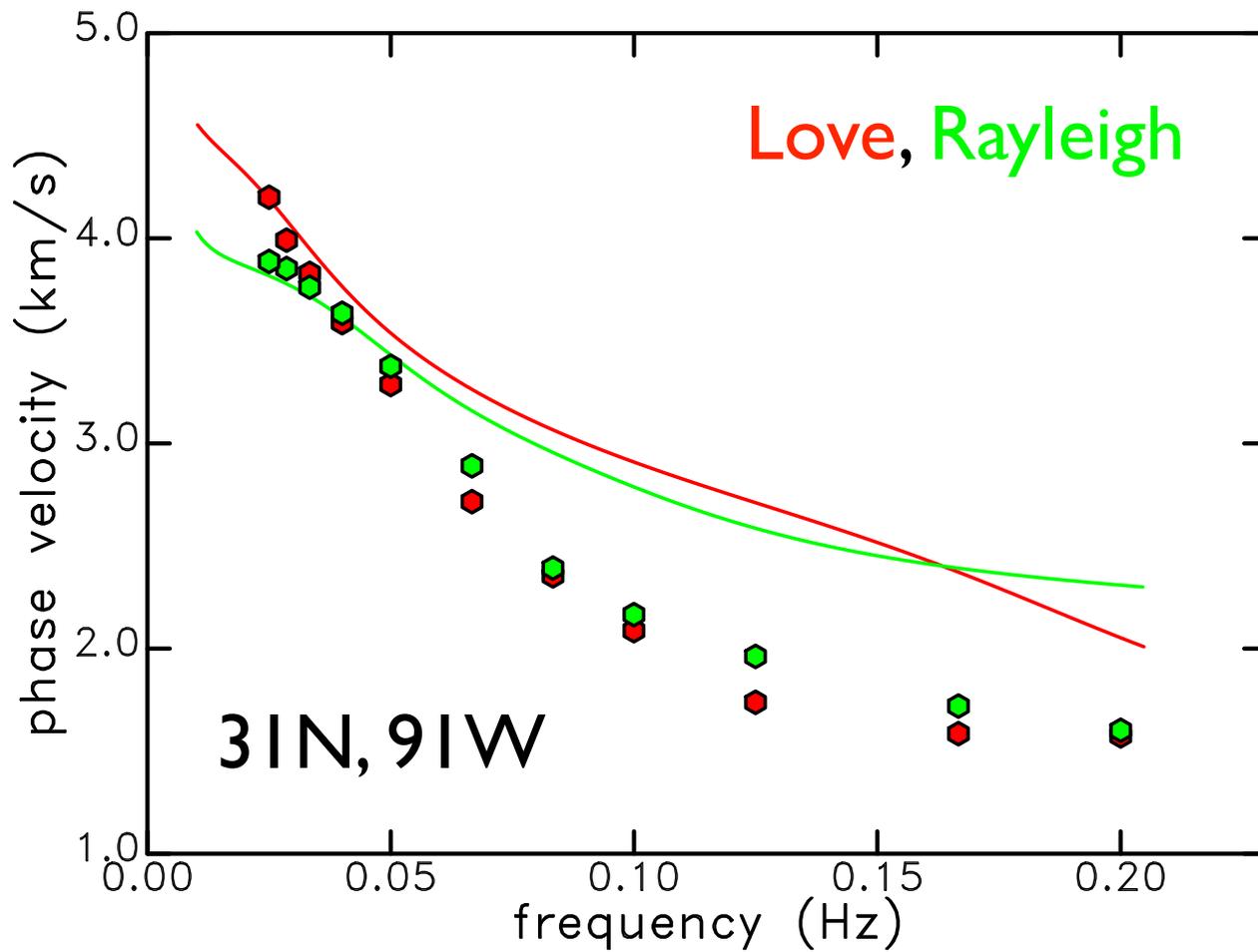
Observations and ND08



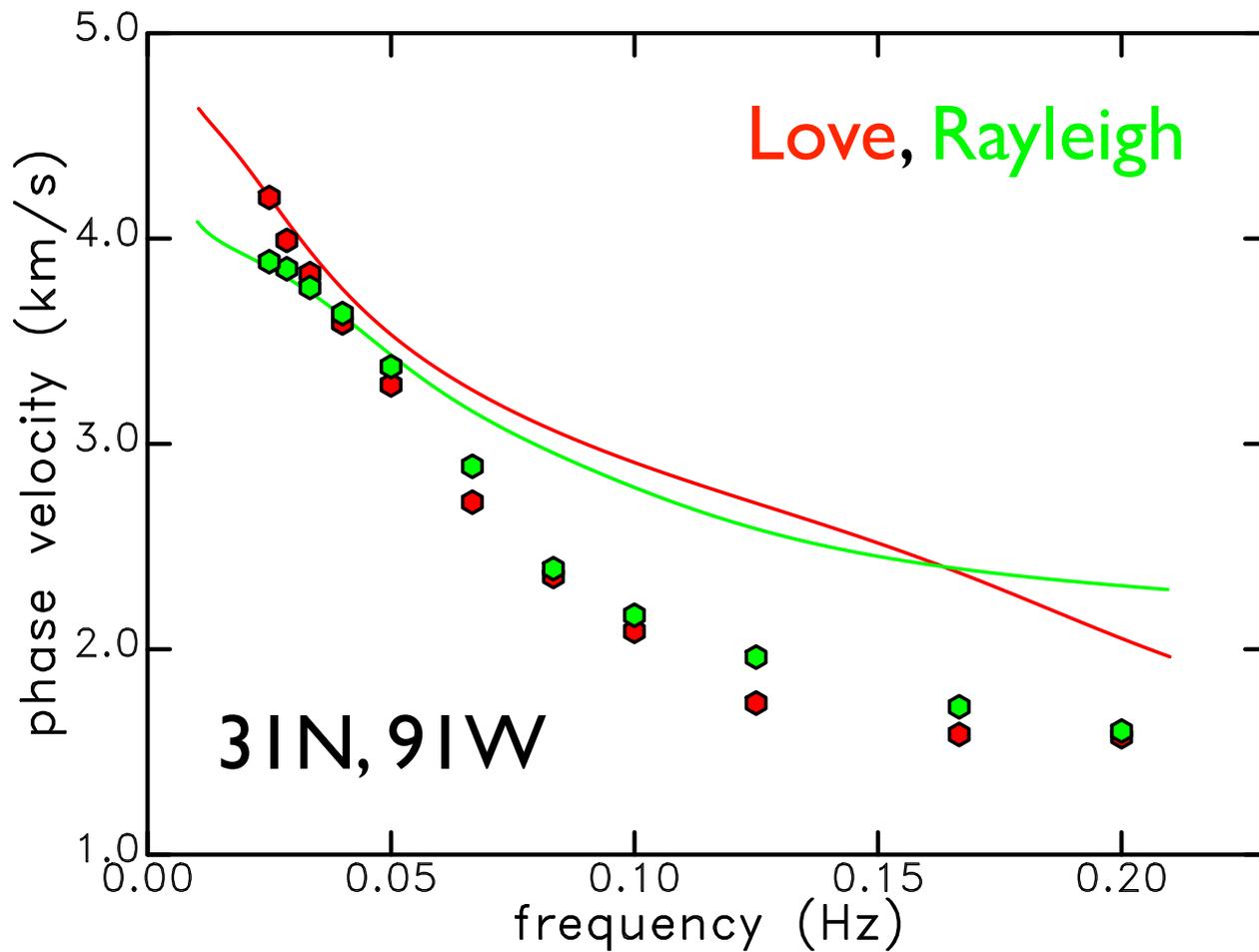
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Observations and CRUST 2.0



Observations and ND08



1. The Transportable Array of USArray allows spatially uniform mapping of surface-wave dispersion across the US using noise tomography
2. Aki's spectral approach works well for automation
3. Extremely slow Love and Rayleigh velocities along the Gulf coast (and in other areas) are not matched by current models of the crust
4. Very low V_S is needed (high V_P/V_S ratio) to explain the signals from the basins